# Well completion report: Ravn Pynt-1 fully cored borehole, Store Koldewey, Northeast Greenland

Contribution to Petroleum Geological Studies Services and Data in East and Northeast Greenland

Gunver K. Pedersen & Jørgen A. Bojesen-Koefoed



GEOLOGICAL SURVEY OF DENMARK AND GREENLAND MINISTRY OF ENVIRONMENT AND ENERGY

## Well completion report: Ravn Pynt-1 fully cored borehole, Store Koldewey, Northeast Greenland

## Contribution to Petroleum Geological Studies, Services and Data in East and Northeast Greenland

Gunver K. Pedersen & Jørgen A. Bojesen-Koefoed

Confidential report

Copy No.

Released 01-07-2025



© De Nationale Geologiske Undersøgelser for Danmark og Grønland (GEUS), 2018 (Geological Survey of Denmark and Greenland)

No part of this publication may be reproduced or utilised in any form or by any means, electronic or mechanical, including photocopying, recording or by any information storage and retrieval system, without prior written permission from the publisher, GEUS. Disputes on copyright and other intellectual property rights shall be governed by Danish law and subject to Danish jurisdiction.

1.	General information	4
1.1 1.2	Borehole data Borehole summary	4 5
2.	Drilling operations	8
3.	Geology	14
3.1	Description of lithology	14
4.	Further analyses	26
5.	References	26

### 1. General information

#### 1.1 Borehole data

Country Borehole number Borehole name Area Operator Drilling operator	Greenland / Denmark GGU 517009 Ravn Pynt-1 Northeast Greenland, Ravn Pynt, Store Koldewey GEUS GEUS			
Borehole Location				
Altitude:	108 m above mean sea level			
Coordinates WGS 84:	76º 08.455´N, 18º 32.745´W			
Drill rig	Sandvik DE 130			
Drilling contractor	GEUS			
Casing diameter	64/57 mm,			
Casing depth	Casing 64/57 mm to 95,5m			
Borehole diameter	56 mm			
Core diameter	42 mm			
Total depth	116.4 m			
Core recovery	95%			
Status	Abandoned, open hole			
Logistic history:				
Transportation of rig and	crew to drill site at Jurakløft	July 24 <sup>th</sup> –28 <sup>th</sup> 2017		
Establishment of field car	mp and drilling rig	July 27 <sup>th</sup> –30 <sup>th</sup> 2017		
Spud, Ravn Pynt-1		July 30 <sup>th</sup> 2017		
Drilling completed		August 2 <sup>nd</sup> 2017		
Drill rig transported to Da	neborg, Wollaston Forland	August 3 <sup>rd</sup> – 4 <sup>th</sup> 2017		
Effective drilling		4 days		
Total days on drill locatio	n	10 days		

4

#### 1.2 Borehole summary

The Ravn Pynt-1, GGU 517009, fully cored borehole was drilled to test the succession of Middle Jurassic sandstones exposed in southern Store Koldewey. The borehole was drilled during the summer of 2017 at Ravn Pynt ("Pynt" = small promontory/cape, "nose"), southern Store Koldewey (Figs 1, 2).

The borehole was drilled by GEUS as part of the ongoing GEUS-industry collaboration "Petroleum Geological Studies, Services and Data in East and Northeast Greenland", however, with participation exclusively by companies holding an exploration license offshore Northeast Greenland. Hence, the borehole was drilled on behalf of a consortium including three groups holding licenses for exploration offshore Northeast Greenland, represented by the operating companies Chevron, and ENI.

The objective was to test the Middle Jurassic sandstone succession exposed in southern Store Koldewey, represented by the Pelion Formation (Surlyk 2003 and references therein), in order to obtain data on the petroleum reservoir properties of this unit, which is expected to be present in the offhore license areas as well.

The Ravn Pynt-1 borehole was spudded on July 30<sup>th</sup> 2017 and completed on August 2<sup>nd</sup> 2017 at a total depth of 116.4 m, at which depth drilling had to be suspended since technical problems and the nature of the formation precluded further drilling. The borehole was planned to reach a depth of more than 150 m, which, however, proved impossible to reach. The total core recovery was slightly less than 95%.

The borehole was abandoned as an open hole.

For technical reasons borehole logging in the field was not attempted.



Fig. 1. Overview map of Northeast Greenland. Yellow star: Basecamp in Daneborg. Red star: Drill-site Jurakøft-1, 2 boreholes. Blue star: Drill site Ravn Pynt-1 borehole. Orange triangle: fuel depot at Twinotter strip near Haystack.Red frame: detailed map, Fig. 2.



Fig. 2 Geological map of Store Koldewey (from Piasecki et al 2012). Blue star: Drill-site Ravn Pynt-1 borehole. Orange triangle: Twinotter strip used for demobilisation.

### 2. Drilling operations

Drilling team:

Driller Jan Varup Driller Tonny Arndt Driller's assistant Kathrine Hedegaard Driller's assistant Kristine Balslev Jørgensen Fasting Driller's assistant Christopher Trinder Nielsen Driller's assistant Rune Hende Assistant (and expedition leader) Jørgen A. Bojesen-Koefoed

Logistics were handled by GEUS.

Helicopter from Air Greenland (Eurocopter AS-350B3) was chartered for a full field season.

Drilling was carried out by a Sandvik DE 130 wire line rig, with inner and outer casing diameter of 57 mm and 67 mm, respectively. The borehole diameter is 56 mm, whereas the core diameter is 42 mm. Core barrels are 1.5m and 3m long.

Most field equipment, helicopter fuel and the drill rig was transported late summer 2016 by ship to Daneborg at Wollaston Foreland, the headquarters of the Sirius dog sledge Patrol. Equipment was left in sealed containers, to be opened when operations were initiated during the summer of 2017. Daneborg served as basecamp for the field campaign, which included both shallow borehole drilling and traditional fieldwork.

Rig, equipment and personnel were mobilised directly from the previous drill site at Jurakløft (Wollaston Forland) by helicopter to the drill site at 76° 08.455 N, 18° 32.745 W, 108 m above sea level (Figs 1, 2). Due to the long distances and heavy sling loads, mobilisation took place in several stages using the Twinotter strip at Haystack as fuel depot and reloading station. In addition, due to limited helicopter capacity, partly caused by constant, strong catabatic winds, drilling was carried out with less manpower than usual, leaving two persons behind at the previous drill site.

The objective was to test the Middle Jurassic Pelion Formation sandstones exposed at Ravn Pynt, in order to carry out analyses of petroleum reservoir properties of the unit. A secondary objective was to penetrate the "base-Cretaceous unconformity" (BCU), which can be observed in the area as a Cretaceous-age valley-incision into the Middle Jurassic succession, filled with Lower Cretaceous mudstones of the Stratumbjerg Formation (Figs 3, 4, 5).

Placing the borehole on a narrow ridge protruding towards the East from the coastal cliffs seemed to allow both objectives to be fulfilled. However, the ridge proved much too narrow to accommodate the drill rig, and an alternative drill site further to the west had to be chosen. This was placed on a narrow ridge protruding towards the East between two gullies, and outcrops on the East-facing front of the ridge prognosed a succession including a few

metres of overburden followed by a few metres of Barremian shale, before encountering the Pelion Formation. This drill site was thus expected also to allow penetration of the BCU.

The Ravn Pynt-1 borehole was spudded on Sunday, July 30<sup>th</sup> at 1200, location 76<sup>o</sup> 08.455 N, 18<sup>o</sup> 32.745 W, 108 m above sea level. Water was pumped to a reservoir by the rig from a creek situated approximately 1.2 km to the NW.

Eventually, despite a distance of less than 50 metres between the outcrops and the drill site, no Stratumbjerg Formation mudstones were encountered, the overburden resting directly on the Pelion Formation sandstones. Hence, the BCU represented by Cretaceous shale on Jurassic sands was not penetrated. The Pelion Formation proved to consist of alternations between highly indurated sandstones, and perfectly soft and unlithified sands. The nature of the deposit was highly challenging for core drilling, and casing to great depth was necessary.

At 0530 a.m. on Tuesday, August 1<sup>st</sup>, permafrost caused the drill stem to become stuck in the hole just before commencing the 95-98 m run. Re-establishing circulation was impossible, but after 4 hours of flushing with hot water, the casing could be rotated. Casing was then drilled from 79.1 m to 95.5 m, thereby releasing the drill stem and core-barrel, which was pulled out of the hole. All drill rods were then checked for damage arising from the violent attempts to release the drill stem and from the process of drilling casing around the immobile drill stem. A considerable length of drill stem had to be discarded. Shortly before midnight, drilling resumed, initially drilling/melting ice in the hole.

Drilling of the Ravn Pynt-1 borehole was eventually terminated shortly after 0400 a.m. on Wednesday August 2<sup>nd</sup> with a frozen and unrecoverable drill stem. The total depth reached was 116.4 metres, and except for the overburden, core recovery was close to 100% with excellent quality core, even in unconsolidated portions of the succession. TD was roughly 8 metres below sea level, and a significant proportion of the penetrated succession is unknown. Given the extremely complicated drilling conditions, this result is satisfactory, despite the planned depth of >150 m was not realised.

The drill stem broke close to the base of the borehole, allowing most of it to be recovered. Logging of the hole was deemed too risky and was not attempted.

Demobilisation started later on the same day by lifting equipment and personnel by helicopter to a Twinotter-strip approximately 7 kilometres SW of the drill site (Fig. 2), from where everything was carried by Twinotter (chartered from Norlandair, Akureyri, Iceland) to Daneborg, allowing demobilisation to be complete by Friday, August 4<sup>th</sup>



Fig. 3 Simplified profile of the study area (from Piasecki et al. 2012). The Middle Jurassic rests on Caledonian basement and is unconformably overlain by Lower Cretaceous marine mudstones of the Stratumbjerg Formation. Where documented, the contact between the Middle Jurassic and the overlying Lower Cretaceous successions is highly irregular and steeply inclined. For legend, see fig. 2.





Fig. 4. Drill site viewed from the NE towards the SW. Upper part: view of the exposures showing both the preferred drill sitre, which proved unsuitable, and the actual drillsite located higher and slightly more westerly. Lower part: dimmed image showing geological units, the steeply inclined BCU and the position of profiles reported in Piasecki et al. (2012). (Photomosaic by Peter Alsen).



Fig. 5 The preferred drill site would have been a few metres left of the hammer, which is placed at the BCU (also marked by white line). However, the narrow ridge was unable to accommodate the drill rig. (Photo by Peter Alsen).



Fig. 6. Drill site after drilling resumed around midnight on August  $2^{nd}$  2017. View towards the NE.



Fig. 7. Sedimentary log of the drilled succession. Minor adjustments may be added to the final version of the log. Legend, see fig 8.

Litho	logy	Sedir	mentary structures
	Mudstone		Low angle cross-stratification
	Sandy mudstone		Low-angle cross-bedding
-	Muddy sandstone		Trough cross-bedding
	Sandstone	<u> </u>	Trough cross-bedding
	Sandstone		Planar cross-bedding
	Heterolithic mudstone/sandstone		Planar lamination/bedding
	Conglomerate		Hummocky cross-stratification
	Sand/sandstone		Hummocky cross-stratification
00000	Granules, pebbles	$\sim$	Swaley cross-stratification
	Sandy silt/silty sand		Sigmoidal cross-stratification
	Sandy silt/silty sand	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	Ripple cross-lamination
	Silty mudstone, siltstone		Combined flow ripple cross-stratification
	Clayey mudstone		Wave-ripple cross-lamination
	Coal	APTC	Climbing ripple cross-stratification
	Coaly mudstone		Herringbone cross-stratification
	Coal clasts	200 20	Mud-draped foresets
0	Mudstone clasts		• Erosive sandstone bed
•••	Sandstone clasts	$\sim$	Wavy bedding
(×)	Concretion	~	Flaser bedding
	Concretion		Lenticular bedding
7,7,7	Calcite cement		Heterolithic bedding (sand dominated)
* Py	Pyrite		Sand streak
1-1	Crystalline basement		Thin sand streak
S	Sulphur		Weak lamination
SS	Slickenside		Indistinct lamination/bedding (often mottled heterolith)
			Structureless
Trace	Fossils		Structureless
Å	Chondrites	1188	Contorted bedding
** \$t	Ophiomorpha isp	22	Loading
***		6	Load structure
6	Palaeophycus tubularis	K	Water escape structure
-0	Planolites		
<u>)</u>	Rhizocorallium isp.		
$\checkmark$ Y	Roots	Pioturb	ation dorroo
J	Skolithos isp.		
0000	Teichichnus isp.		61–100 % intense
Ū	Vertical shaft	5	1-5 % weak

#### Sedimentary structures

37	Synaeresis	crack	ς

- → Ptygmatic fold
- S Slump fold
- Slump fold
- Soft sediment deformation
- \_\_\_\_\_ Flame structure
- √ Gutter cast
- w\_\_\_\_\_\_\_Stylolite
- @ Imbrication
- A Biomottled

#### Fossils

- 😪 🎳 Fragment
- A Belemnite
- L L Rootlets
- φ Plant fragments
- Ammonite
- Logs
- Ø Bivalve
- Oyster

#### Miscellaneous >\_< Joints 0 Vertical fracture 4 Intrusion, intrusive sandstone Erosional boundary -Sharp boundary ---Gradational boundary Palaeocurrent direction \* Current lineation ~ Wave ripple crest orientation

#### Fig 8. Legend for sedimentary log

## 3. Geology

Lithostratigraphy of Store Koldewey (Table 1):

			Lithostratigraphy		Main depositional
					environment
	Lower	Barremian, upper	Stratumbjerg Fm		Upper slope channel?
			Sorte Kløft Fm		Marine slope w gravity flows
		Barremian, Lower			(sst)
		Hauterivian	Delastekee	Down Dwnt Mh	
sno		Valanginian Bjerg Fm	Rierg Em		Slope channel
tace			Бјегд ЕШ	Midter Gneisnes	Dobris flows in slope shapped
Cret			Mb	Debris nows in slope channel	
		Byazanian	Lindemans Bugt Fm		Marine slope w gravity flows
		Nyazaman			(sst)
		Volgian, middle-upper			
		Volgian, lowermost	Bernhierg Em		
	Upper	Kimmeridgian, upper	bennbjerg i m		
		Kimmeridgian, lower	Payer Dal Fm		
ssic		Oxfordian, upper			
Jura		Oxfordian, lower			
	id	Callovian	Dolion Em	Spath Plateau Mb	
	Σ	Bathonian		undivided	

Table 1. The table is based on the chapter "Depositional evolution of Store Koldewey" by Hovikoski, Bjerager & Piasecki (2012) in Piasecki et al. (2012).

The Middle Jurassic succession at Store Koldewey is referred to the Pelion Formation of Surlyk (2003). The Pelion Fm comprises a lower part, of Bathonian age, which is not referred to members, and an upper part referred to the Spath Plateau Member of Callovian age (Vosgerau et al. 2004; Piasecki et al. 2004). The lower boundary of the Spath Plateau Member is tentatively located at c. 30 m in the Ravn Pynt-1 core. The final identification of the boundary awaits the results of the biostratigraphic analysis. The Spath Plateau Mb. overlies an erosional unconformity, which covers a longer period at Ravn Kløft than in outcrops further south (Piasecki et al 2004: their Fig. 3).

The Upper Jurassic succession on Store Koldewey is Oxfordian – Kimmeridgian, and is referred to the Payer Dal Formation of Alsgaard et al. (2003). It is overlain by the upper Kimmeridgian Bernbjerg Formation (Piasecki et al. 2004: figs: 5, 6).

#### 3.1 Description of lithology

The Ravn Pynt-1 core recovered a 116.4 m thick succession, representing the upper part of the Pelion Formation. The uppermost c. 30 m are referred to the Spath Plateau Member,

outlined in blue in Table 2. The succession below is referred to the undivided, lower part of the Pelion Fm. The base of the Pelion Formation was not reached. Five lithologies are distinguished in the preliminary description pf the core. Three of these, the medium-grained sandstone; the heterolithic, fine-grained sandstone; and the bioturbated muddy sandstone recur several times through the core, as shown in Table 2. Two lithologies, laminated mudstone and muddy sandstone are restricted to single intervals. Brief descriptions of the lithologies are found below. The trace fossil assemblage awaits a more detailed examination. The lithological description starts from the top of the core.

Strata		Lithology	Donth (m)	Thickness	
		Lithology	Depth (m)	(m)	
	SP Mb		5–15.6 m		
		Medium-grained, well-sorted, friable	47.5–52.7 m	c. 49 m	
		sandstone.	58.5–79.5m		
		Locally interbedded with the fine-grained	83–88 m		
		heterolithic sandstone	102.5–107 m		
			113.5–116.4 m		
	SP Mb		15.6–18.9 m		
		Heterolithic fine-grained sandstone with	29.5–43.3 m	c. 30 m	
		moderate to strong bioturbation.	52.7–53.5 m		
		Locally interbedded with the medium-	60.7–63.6 m		
Pelion Fm		grained sandstone, or the bioturbated	79.4–83 m		
		very fine-grained sandstone	88–94 m		
			101.3–102.4		
	SP Mb		18.9–29.5 m		
		Bioturbated, very fine-grained muddy	53.5–58.5 m		
		sandstone.	74–74.8 m	0.06 m	
		Locally interbedded with the fine-grained	89.5–92 m	C. 20 III	
		heterolithic sandstone	96.5–101.3 m		
			107–108 m		
		Laminated and weakly laminated mud-	108 100 7 m	c 2 m	
		stone	100-109.7 11	0.2111	
		Medium-grained, muddy sandstone	43.3–47.5 m	c. 4 m	

Table 2. Five lithologies characterize the Ravn Pynt-1 core at St. Koldewey. The Spath Plateau Member is indicated in blue. The lower c. 30 m of the member constitute the upper part of the core. The dominant lithology is the medium-grained, well-sorted, friable sand-stone, with a cumulative thickness of c. 49 m.

#### Medium-grained, well-sorted, friable sandstone

#### 5–15.6 m, and deeper intervals in Ravn Pynt-1 (Table 2)

Greyish brown sandstone, cross-bedded, medium- to coarse-grained, well-sorted with few trace fossils, including *Ophiomorpha* isp. Coal- and mud-clasts occur locally. The sand-stone is generally friable, but includes cemented intervals, which may represent concre-

tions. The sandstone may have been deposited in a fairly high-energy environment possibly within the middle to upper shoreface.

Comparable sandstones occur at greater depth in the Ravn Pynt-1 core.

#### Heterolithic, fine-grained sandstone with moderate to strong bioturbation of the Pelion Fm

#### 15.6–18.9 m, and deeper intervals in Ravn Pynt-1 (Table 2)

The heterolithic sandstone is fine-grained and contains small amounts of mudstone as mud-drapes and mud-clasts. The sand is well-sorted, and the sandstone is locally cemented. The sandstone is moderately bioturbated, the trace fossils awaits further study. Fragments of belemnites and coalified wood occur locally.

The belemnites indicate a marine environment, and the heterolithic sandstone suggests deposition at depths around wave-base. The core shows examples of transitions between the heterolithic fine-grained sandstone and the medium- to coarse-grained, well-sorted sandstone (described above) or transitions to the very fine-grained muddy sandstone (described below).

Comparable sandstones occur at greater depth in the Ravn Pynt-1 core.

#### Bioturbated, very fine-grained muddy sandstone

#### 18.9–32 m, and deeper intervals in Ravn Pynt-1 (Table 2)

This sandstone is brownish grey and very fine-grained, locally silty. The sandstone is thoroughly bioturbated, and physical sedimentary structures are not preserved. An overall fining-upwards trend is seen from 21–32 m and an upwards coarsening trend characterize the upper part of the interval and continues into the heterolithic fine-grained sandstone (21–16 m). Similar trends are seen in the interval 52.7–56 m.

Comparable sandstones occur at greater depth in the Ravn Pynt-1 core.

#### Bioturbated, very fine-grained muddy sandstone.

#### 18.9-30.5 m, and deeper intervals

Medium grey, very fine-grained sandstone to siltstone, which is strongly bioturbated. The trace fossils have not yet been examined, but they resemble traces, which have been interpreted as produced by deposit feeders. In some intervals (56–58.5 m; 110.3–112.3 m), the very fine-grained sandstone is interbedded with sharp-based, graded somewhat coarser sandstone beds, which tentatively are interpreted as storm sand layers. This suggests that the muddy, very fine-grained sandstone was deposited at depths between fair-weather wave-base and storm wave-base.

## Laminated and weakly laminated mudstone

#### 108–109.7 m

The dark grey, mudstone is restricted to a single occurrence. The mudstone is laminated or weakly laminated, and locally sand-streaked. It is tentatively interpreted as deposied below storm wave-base in a shelf environment, and the mudstone may possibly contain a maximum flooding surface.

# Medium-grained, muddy sandstone 43.3–47.5 m

Grey, "dirty", medium-grained sandstone, which appears to have a muddy matrix and scattered mud-clasts. This lithology appears to be restricted to one interval. The type and amount of matrix is not yet documented.

Figs 9 through 25 (this and following pages) show photos of cores. The photos were taken at the drill site, and core-depths are preliminary, related to markers used as datums for measurements, which may occasionally cause a small apparent overlap of cores. Slight revisions of depths will take place during later core description and laboratory handling of the cores.



Fig. 9. Core photo from drill site. Box 1: 3.10 - 12.83 m



Fig. 10. Core photo from drill site. Box 2: 12.90 - 19.80 m



Fig. 11. Core photo from drill site. Box 3: 19.85 - 26.61 m



Fig. 12. Core photo from drill site. Box 4: 26.52 - 33.42 m



Fig. 13. Core photo from drill site. Box 5: 33.50 - 40.40 m



Fig. 14. Core photo from drill site. Box 6: 40.49 - 47.20 m



Fig. 15. Core photo from drill site. Box 7: 47.28 - 54.08 m



Fig. 16. Core photo from drill site. Box 8: 54.12 - 60.52 m



Fig. 17. Core photo from drill site. Box 9: 60.65 - 67.48 m



Fig. 18. Core photo from drill site. Box 10: 67.43 - 74.27 m



Fig. 19. Core photo from drill site. Box 11: 74.15 – 81.77 m



Fig. 20. Core photo from drill site. Box 12: 81.76 - 88.40 m



Fig. 21. Core photo from drill site. Box 13: 88.46 - 95.22 m



Fig. 22. Core photo from drill site. Box 14: 95.22 - 101.99 m



Fig. 23. Core photo from drill site. Box 15: 101.99 - 108.80 m



Fig. 24. Core photo from drill site. Box 16: 108.84 - 115.36 m



Fig. 25. Core photo from drill site. Box 16: 115.41 - 116.40 m

#### 4. Further analyses

- Detailed sedimentological analysis and core description
- High resolution photography in both UV and white light
- Core scanning to produce a spectral GR-log
- Biostratigraphic dating
- Systematic porosity and permeability analysis
- Diagenesis analyses

Data from such studies will be included in the final reporting.

#### 5. References

Alsgaard, P.C., Felt, V.L., Vosgerau, H. & Surlyk, F. 2003 : The Jurassic of Kuhn Ø, North-East Greenland. Geological Survey of Denmark and Greenland Bulletin 1, 865–892.

Piasecki, S., Alsen, P., Anthonsen, K. Bjerager, M., Bojesen-Koefoed, J.A., Fyhn, MBW, Japsen, P., Hovikoski, J., Kjøller, C., Morigi, C., Nielsen, L.H., Nøhr-Hansen, H., Sheldon, E., Wiebel, R. & Kazerouni, A.M. 2012: The geology of Store Koldewey, North-East Greenland, 76° - 76.45°N Implications for offshore petroleum geology. GEUS rapport 2012/33 20pp. *Confidential report and GIS compilation available to companies participating in the GEUS-industry collaboration project "Petroleum studies, services and data in East and Northeast Greenland"* 

Piasecki, S., Callomon, J.H. & Stemmerik, L. 2004: Jurassic dinoflagellate cyst stratigraphy of Store Koldewey, North-East Greenland. Geological Survey of Denmark and Greenland Bulletin 5, 99–112.

Surlyk, F. 2003: The Jurassic of East Greenland: a sedimentary record of thermal subsidence, onset and culmination of rifting. In: Ineson, J, R. and Surlyk F. (eds.): The Jurassic of Denmark and Greenland. Geological Survey of Denmark and Greenland Bulletin 1, 659-722.

Vosgerau, H., Larsen, M., Piasecki, S. & Therkelsen, J. 2004 : A new Middle-Upper Jurassic succession on Hold with Hope, North-East Greenland. Geological Survey of Denmark and Greenland Bulletin 5, 51–71.