

Dinoflagellate cyst biostratigraphy of the Upper
Cretaceous black mudstones in central
Nuussuaq, West Greenland

Henrik Nøhr-Hansen

Open File Series 94/12



April 1994



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Abstract

Stratigraphical ranges and geographical distribution of dinoflagellate cysts and selected pollen species are described based on analysis of approximately 100 samples from 15 surface and 4 subsurface sections of Late Cretaceous age in central Nuussuaq, West Greenland. The sections make up an approximately 1100 m thick black mudstone succession, previously dated as Santonian to Maastrichtian on the basis of scattered ammonite occurrences.

The dinoflagellate cysts date the majority of the studied samples to Late Santonian to middle/Late Campanian, whereas one sample is dated as Late Paleocene. It has been possible to divide the studied succession into eight intervals based on eight distinguishable palynomorph assemblages. The diversity of the studied dinoflagellate cysts is relatively low to very low; approximately 55 species were recorded.

The presence of reworked specimen of the Early Cretaceous species *Batioladinium jaegerii* indicates, for the first time that pre-Middle Cenomanian marine deposits may have been deposited in central Nuussuaq.

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INTRODUCTION

The present study is part of the EFP-91 project ‘Sequence stratigraphic analysis of the Cretaceous sediments in West Greenland’, and aims to set up a palynostratigraphy for sequence stratigraphical analysis for the Disko–Nuussuaq–Svartenhuk Halvø area in West Greenland (Christiansen *et al.*, 1992; Christiansen, 1993).

The Upper Cretaceous–Lower Tertiary black mudstone succession on Disko, Nuussuaq and Svartenhuk Halvø represents the only marine sediments of this age exposed in the entire Labrador Sea–Baffin Bay region and studies of these are therefore essential for a detailed interpretation of the offshore geology.

The stratigraphic correlation of the sediments in the region is problematic due to the presence of interdigitating Cretaceous fluvial, deltaic and brackish to fully marine deposits (Scheiner, 1975; Pulvertaft, 1979, Pedersen & Pulvertaft, 1992).

Field work on the marine succession in the summers 1990 to 1992 has been concentrated on detailed sedimentological studies, sampling for palynological and organic geochemical studies, and mapping and structural analysis (Christiansen *et al.*, 1992).

PREVIOUS PALYNOLOGICAL STUDIES IN THE SANTONIAN TO CAMPANIAN OF GREENLAND AND ELSEWHERE

West Greenland

Previous studies of Upper Cretaceous dinoflagellate cysts from West Greenland are sparse (Croxton, 1976; 1978; 1978a; 1980; Ehman *et al.*, 1976; Lentin & Williams, 1980). Lentin & Williams mentioned (1980, p. 20) that the Campanian assemblage from West Greenland contains elements of both the offshore eastern Canadian assemblages (also called the Williams suite) and the Mackenzie Delta assemblages from Arctic Canada (the so-called McIntyre suite) described by McIntyre (1974; 1975). The Paleocene dinoflagellate cyst assemblages from West Greenland have been described by Hansen (1980).

Croxton (1978; 1978a) briefly described the palynomorph content from five localities at central Nuussuaq (C4-C7, C21, M19, M22). The palynomorphs from the localities at

Qilakitsoq (C4), Qaatunnat Ilorliit (C5), Ilugisssøq (C6) and Nallurarissat (C7) indicate a Late Cenomanian to Early Campanian age; a possible reworked Maastrichtian assemblage is recorded from the top of section C5, whereas dinoflagellate cysts from the topmost shale at C6 indicate a possible middle Paleocene age (Croxton, 1978). Section C21 and M19 represent the ‘Oyster-ammonite conglomerate’ from Agatdalen; according to Croxton (1978a) palynomorphs from this section may indicate a reworked Maastrichtian flora. Eight sections from central Nuussuaq (M16-M23) were sampled by Hansen (1976); data on the palynological content from the two sections M16 and M17 from Tunoqqu have not been published. A few dinoflagellate cysts probably indicating a Late Campanian age were recorded by Croxton (1978) from Scaphitesnæsen (M22). Hansen (1980) described the Paleocene dinoflagellate cyst content from the Sonja section (M18), Turritellakløft section (M20), Qaarsutjægerdal section (M21) and Ättestupet section (M23). According to Hansen (1980) and the present study the middle Campanian to earliest Maastrichtian species *Isabelidinium microarmum* has been recorded as reworked specimens in the sections M18-M21.

Ehman *et al.* (1976) studied the four sections Qilakitsoq (N10), Turritellakløft (N15), Qaarsutjægerdal (N16) and Nassaat (N17) in central Nuussuaq. The ages given by Ehman *et al.* (1976) are middle Cenomanian and Early Danian for N10, Campanian or Maastrichtian to Paleocene for N15, Paleocene for N16 and Paleocene for N17. It should be noted that the ages given in the text by Ehman *et al.* (1976) are not always consistent with the ages given in their logs (Pulvertaft, 1987; Table 1).

The above mentioned palynological investigations have been reviewed by Pulvertaft (1987).

Arctic Canada

Santonian to Maastrichtian Upper Cretaceous dinoflagellate cysts have been described from Arctic Canada by Manum (1963), Manum & Cookson (1964), Felix & Burbridge (1976), McIntyre (1974, 1975), Doerenkamp *et al.* (1976), Ioannides & McIntyre (1980) and Núñez-Betelu & Hills (1992). Ioannides (1986) studied the dinoflagellate cyst assemblages from the Santonian to Maastrichtian part of the Kanguk Formation and the Lower Paleocene Eureka Sound Formation on Bylot and Devon Islands. The dinoflagellate cyst assemblages described by Ioannides (1986) is very similar to the material from West

Greenland, unfortunately Ionnanides' stratigraphy is not very detailed, due to poor outcrop, and absence of macrofossils.

Western Canada, Western U.S.A.

Upper Cretaceous dinoflagellate cyst assemblages from Western Canada and Western U.S.A. have been described by Wall & Singh (1975), Harland (1973), (Harland, 1977), Sweet & McIntyre (1988), Stone (1973) and Harker *et al.* (1990).

Offshore Eastern Canada, Eastern U.S.A. and the northern hemisphere

The Upper Cretaceous dinoflagellate cysts from offshore eastern Canada are described by Barss *et al.* (1979), Bujak & Williams (1978), Williams (1975), Williams & Brideaux (1975), Williams & Bujak (1977a, 1977b), Williams *et al.* (1974) and Williams *et al.* (1990).

The stratigraphical distribution of Mesozoic and Cenozoic dinoflagellate cysts has been described by Williams & Bujak (1985) for the world and by Williams *et al.* (1993) for the northern hemisphere.

Upper Cretaceous to Paleocene dinoflagellate cyst assemblages from eastern U.S.A. are described by Benson (1976), May (1980), Tocher (1987), Aurisano & Habib (1977) who established a Campanian to lowermost Tertiary dinoflagellate cyst zonation and by Aurisano (1989) who proposed a Cenomanian to Maastrichtian dinoflagellate cyst zonation for the Atlantic Coastal Plain of New Jersey and Delaware.

Northern North Sea, clastic deposits

According to Costa & Davey (1992, pp. 105–106) dinoflagellate cyst information has not been published from this regions. However, unpublished personal observation by Lucy. I. Costa (reported in Costa & Davey, pp. 105–106) indicates assemblage affinities with the Upper Cretaceous arctic assemblages described by Vozzhennikova (1967) from Siberia, Manum & Cookson (1964) and Doerenkamp *et al.* (1976) from Arctic Canada, and McIntyre (1974) from the District of Mackenzie, Canada.

North-western Europe

The stratigraphical distribution of Upper Cretaceous dinoflagellate cysts in North-West Europe has been compiled by Foucher (1979) and by Costa & Davey (1992). Clarke & Verdier (1967) described the Cenomanian to Campanian on the Isle of Wight and made the first and only attempt to establish a dinoflagellate zonation for the British Upper Cretaceous. Hart *et al.* (1987) listed dinoflagellate cysts together with microfossils from key Upper Cretaceous sections on the Isle of Wight.

Robaszynski *et al.* (1980) described dinoflagellate cyst assemblages of Albian to Santonian age from France. Westin (1992) established a dinoflagellate cyst stratigraphy from the Albian to Santonian in the southern Sweden. The diverse assemblages described from Sweden (Westin, 1992) are dominated by North Sea and North-West European limestone facies species; however the abundance of the northern North Sea genus *Chatangiella*, especially in southern Sweden, is remarkable.

Australia, Antarctica

There are numerous papers describing Cretaceous dinoflagellate cysts from Australia. Helby *et al.* (1987) established a palynological zonation covering the entire Mesozoic of Australia.

Askin (1988) described the Campanian to Eocene palynological succession of Seymour Island and adjacent islands, Antarctica.

Mohr & Gee (1992) and Mao & Mohr (1992) described the Cenomanian to Maastrichtian dinoflagellate cyst assemblages from the ODP leg 120 in the southern Indian Ocean.

The interesting point about the Upper Cretaceous palynomorphs recorded from Australia and around Antarctica is that they are very similar at assemblage level to the material recorded from West Greenland, whereas at species level there are small but distinguishable differences between superficially similar species from the two regions, which makes direct correlation difficult.

SAMPLES AND METHODS

Samples

This study covers samples from central Nuussuaq obtained from fifteen outcrop localities and four slim cores from shallow wells drilled by GGU in 1992 (Figs 1, 2). The sections are 2 to 286 metres thick and represent an approximately 1100 m thick sandy shale succession (Fig. 3).

Preparation

Palynological preparation and studies were carried out at GGU. Palynomorphs were extracted from 20 g of sample by modified standard preparation techniques. The bulk of the minerals were dissolved by hydrochloric and hydrofluoric acids. A first slide was made after this treatment. A second slide was made of the organic residue after sieving using a 20 micron nylon mesh. A third slide was made after oxidation (3 to 10 minutes) with fuming nitric acid, followed by washing with a weak potassium hydroxide solution. The oxidation was carried out in order to clean the sample of minor amorphous kerogen particles and pyrite. Finally, palynomorphs were separated from coal particles and woody material in most samples using the method described by Hansen & Gudmundsson (1978).

After each of the steps mentioned above the organic residues were mounted in a permanent medium (Eukitt R; produced by O. Kindler, Germany).

Recording of material and analyses

The palynological slides were studied with transmitted light using a Leitz Dialux 22 microscope (512 742/057691). All the coordinates in the plate text refer to this microscope. England finder index corners: Z 75 4 = 74.6–92.3; Z 1 3 = 1.9–9220; A 1 1 = 1.9–116.7; A 65 2 = 64.6–116.6, centre: O 38 = 38.1–103.3.

The illustrated dinoflagellate cysts are marked with GGU number (sample number), slide number, microscope coordinates, laser-video-record number (LVR) and database number (MicroImage; MI) for later identification. The slides are housed at the Geological Survey of Greenland where they are accessible for examination.

Composition of the organic material and maturation

The organic material is dominated by terrestrially derived black to brownish woody material and cuticles, whereas amorphous organic material, dinoflagellate cysts, spores and pollen constitute a minor part.

The TAI (Thermal Alteration Index) evaluation was carried out on the sieved slide before oxidation. The study revealed TAI values between -2 and +3, which indicate that the organic material is thermally immature to mature with respect to oil generation.

DINOFLAGELLATE CYST STRATIGRAPHY AND ZONATION IN CENTRAL NUUSSUAQ

A dinoflagellate cyst stratigraphy has been established for the Upper Cretaceous sediments on central Nuussuaq (Figs 1 & 2). Marine palynomorphs were recorded from all the 19 studied sections. Due to the very sparse macrofossil content and the monotonous lithology, the stratigraphical correlation of the geographically widely spread 19 sections is based solely on the first and the last occurrences and acme of stratigraphically important dinoflagellate and selected pollen species.

The present study has been very time consuming and the dating and stratigraphical correlation is based on limited observations as the dinoflagellate cyst content in the sections are very low, normally a slide contains one to ten specimens. However, terrestrially dominated samples are not uncommon.

Most of the studied samples with marine palynomorphs contain one or more large specimens of the genus *Chatangiella*. According to the literature the genus *Chatangiella* ranges from the Upper Cenomanian to the Upper Maastrichtian (e.g. Costa & Davey, 1992). The genus *Chatangiella* dominates Upper Cretaceous assemblages in the Western Interior, U.S.A., western Canada, Arctic Canada and the northern North Sea. The genus *Chatangiella* is also very abundant in the southern hemisphere (especially in Australia and Antarctica), whereas it is less distributed in northwestern Europe and in the Tethyan realm (Lentin & Williams, 1980, Costa & Davey, 1992).

The presence of *Heterosphaeridium difficile* and *Chatangiella* together with the absence of the characteristic species *Litosphaeridium siphoniphorum* and *Stephodinium coronatum* (both of which have their last occurrence in the Turonian) suggests according to

Haq *et al.* (1987) and Costa & Davey (1992) a post-Turonian to pre-Campanian age for the three sections GKP 92 1 Nall, GKP 92 V 1 Qilak. & GKP 92 V 2 Qilak. (Encl. 16, Encl. 19 and Encl. 20).

The pollen genus *Aquilapollenites* has been recorded in fourteen of the studied sections. According to Traverse (1988) *Aquilapollenites* has a sporadic occurrence from Late Turonian to Late Santonian, whereas the occurrence becomes consistent in the latest Santonian and occurs throughout to the Early Paleocene. McIntyre (1974) did not record *Aquilapollenites* species in sediments older than middle to Upper Campanian in the District of Mackenzie, Arctic Canada. Croxton (1980, p. 16) concluded “Although only a preliminary assessment has to date been made of the earliest occurrence of *Aquilapollenites* in West Greenland it is not thought to occur in strata older than Campanian in age”. The fact that Nøhr-Hansen (1994) did not record *Aquilapollenites* in the Coniacian to Upper Santonian sediments on Svartenhuk Halvø, suggests that the genus has a post-Late Santonian occurrence in West Greenland.

The presence of the species *Isabelidinium* cf. *acuminatum* and *I. microarmum* indicates a Campanian age according to McIntyre (1975) and Costa & Davey (1992).

The species list on the cumulate range chart (Encl. 1) illustrates that the assemblages are of very low diversity. Approximately 55 species were recorded. Based on the first and last occurrences, presence and absence of a few morphologically characteristic and stratigraphically important species, it has been possible to distinguish eight intervals with characteristic dinoflagellate and pollen content from the cumulate section.

Unnamed interval

The very sparse dinoflagellate cyst assemblage recorded from the unnamed interval (Encl. 1, 14, 19, 20) indicates the presence of the oldest recorded marine influenced depositional environment in central Nuussuaq.

Age. The age of the interval is most likely Late Santonian but a latest Turonian or Coniacian age cannot be excluded.

Definition. The interval is defined by containing almost no dinoflagellate cysts, its upper limit being the lowermost occurrence of the species *Heterosphaeridium difficile*.

Thickness and distribution. The entire section FGC900804/2 Kan. (Encl. 14; 244 m) and the lower part of section GKP 92 V 1 Qilak. (Encl. 19; 120 m) and GKP 92 V 2 Qilak. (Encl. 20; 96 m) are described as the unnamed interval.

Characteristic species. The interval is characterised by the presence of almost no dinoflagellate cysts; only a few specimens of the genera *Chatangiella* and *Isabelidinium* have been recorded.

Discussion. The presence of the genera *Isabelidinium* and *Chatangiella*, especially the genus *C. granulifera*, indicate an Early Coniacian to Late Campanian range according to Williams & Bujak (1985), whereas Costa & Davey (1992) reported an Early Turonian to Late Campanian range for *C. granulifera*. According to Costa & Davey's (1992) observations from the North Sea region, the absence of *Heterosphaeridium difficile* in the unnamed interval indicates an age not younger than Early Turonian. The absence of *H. difficile* in the unnamed interval may be explained by an Early Turonian age of the sediments. It is however, probably a consequence of the low content dinoflagellate cysts. The unnamed interval might correlate with part of the Late Santonian *H. difficile* Zone on Svatenhuk Halvø, which also is represented by a low diversity assemblage (Nøhr-Hansen, 1994).

***Heterosphaeridium difficile* interval**

Age. Late Santonian.

Definition. The interval is defined from the first occurrence of *Heterosphaeridium difficile* to the last occurrence of *Heterosphaeridium difficile*.

Thickness and distribution. The interval is represented by one sample in each of the sections GKP 92 V 2 Qilak. (Encl. 20) and GKP 92 1 Nall. (Encl. 16) The interval constitutes approximately 95 m of the section GKP 92 V 1 Qilak. (Encl. 19).

Characteristic species. The interval is characterised by a relatively diverse dinoflagellate cyst assemblage, whereas the abundance is low. The following species have been recorded: *Heterosphaeridium difficile*, *Laciniadinium arcticum*, *Spinidinium* aff. *echinoideum*, *Odontochitina striatoperforata*, *Chatangiella* aff. *ditissima*, *Surculosphaeridium?* *longifurcatum*, *Chatangiella granulifera*, *Desmocysta plekta*, *Xenascus* aff. *perforatus* and *Florentinia* spp.

Discussion. The last occurrences of *Heterosphaeridium difficile* within the interval indicates an age no younger than Late Santonian (Costa & Davey, 1992). The *H. difficile* interval correlates with the *H. difficile* Zone proposed for the Late Santonian on Svartenhuk Halvø (Nøhr-Hansen, 1994).

***Chatangiella* aff. *ditissima* interval**

Age. Late Santonian or ?Early Campanian

Definition. The interval is defined from the last occurrence of *Heterosphaeridium difficile* to immediately below the first occurrence of the genus *Aquilapollenites*.

Thickness and distribution. The interval is represented by one sample in the two sections GKP 92 V 1 Qilak. (Encl. 19), GKP 92 1 Nall. (Encl. 16), 59 m of well GGU 400703 (Encl. 5) and by the lower 65 m in well GGU 400704 (Encl. 6).

Characteristic species. There are no characteristic species in this interval. The species diversity is very low. The species *C. aff. ditissima* is almost the only species that has been recorded in more than one sample.

Comments. The *C. aff. ditissima* interval may be dated Late Santonian or ?Early Campanian based on the absence of *H. difficile*. The interval has not been recorded in the Late Santonian deposits on Svartenhuk Halvø (Nøhr-Hansen, 1994).

***Aquilapollenites* 1. interval**

Age. Early Campanian

Definition. The interval is defined from the first occurrence of the genus *Aquilapollenites* to immediately below the first occurrence of *Isabelidinium cf. acuminatum*.

Thickness and distribution. The interval is represented by 86 m in section GKP 91 4 Tun. (Encl. 8), the lowermost 23 m in section GKP 91 3 Tun. (Encl. 9), the lowermost 40 m of section HNH910816/1 (Encl. 13). and possible by the threee lowermost samples in section HNH910811/1 (Encl. 10).

Characteristic species. The first occurrence of the genus *Aquilapollenites* is the characteristic for the interval. A few specimens of the species *Laciadiinium arcticum* and *C. granulifera* have been recorded from section GKP 91 4 Tun. (Encl. 8). The species *Batioladinium jaegerii* and a ?hair from a leaf seem to be reworked in section GKP 91 4 Tun. (Encl. 8).

Discussion. According to Traverse (1988) the occurrence of *Aquilapollenites* becomes consistent in the latest Santonian and occurs throughout to the Early Paleocene.

Aquilapollenites species have not been recorded in sediments older than middle to Upper Campanian in the District of Mackenzie, Arctic Canada McIntyre (1974). On Svatenhuk Halvø *Aquilapollenites* has not been recorded in the Coniacian to Upper Santonian succession by Nøhr-Hansen (1994), which suggests that the genus has a post-Late Santonian occurrence in West Greenland.

According to Costa & Davey (1992) *Batioladinium jaegerii* has its last occurrence in the early Middle Cenomanian. The presence of reworked specimens of *B. jaegerii* in the

Campanian section GKP 91 4 Tun (Encl. 8) indicates that Lower Cretaceous marine influenced sediments have been deposited somewhere in central Nuussuaq. The presence of the miscellaneous ?leaf-hair also suggests reworking of Lower Cretaceous sediments. A morphologically similar ?leaf-hair has previously been recorded from Late Barremian to Early Albian in East Greenland and Late Barremian in Arctic Canada (Nøhr-Hansen, 1993).

***Isabelidinium cf. acuminatum* interval**

Age. Early Campanian

Definition. The interval is defined from the first occurrence of *Isabelidinium cf. acuminatum* to the last occurrence of *Isabelidinium cf. acuminatum*.

Thickness and distribution. The Interval is represented by the upper 15 m in section GKP 91 3 Tun. (Encl. 9), by the uppermost sample in section HNH910811/1 Tun. (Encl. 10) and by approximately 45 m in section HNH910816/1 Kan. (Encl 13).

Comments. The *I. cf. acuminatum* interval is proposed to be slightly older than the *I. microarmum* interval. The interval may, however, continue up to the *I. microarmum* interval and thereby eliminate the *Aquilapollenites* 2 interval. For further discussion see comments under the *I. microarmum* interval.

Characteristic species. The diversity in this interval is high compared to the underlying *Aquilapollenites* 1 interval. It is characterised by the presence of *Isabelidinium cf. acuminatum*, *I. aff. ?amphiatum* large, *Laciadiinium arcticum*, *Odontochitina striatoperforata*, *Chatangiella* aff. *ditissima*, *C. granulifera*, *Desmocysta plekta* and *Exocosphaeridium* sp.. Only one specimen of the following species has been recorded in the interval: *Surculosphaeridium* aff. *longifurcatum*, *Coronifera oceanica*, *Tanyosphaeridium variacalamus* and *Florentinia* aff. *mantelli*.

Discussion. The species *Isabelidinium acuminatum* has an Early to middle Campanian range in the District of Mackenzie (McIntyre, 1975). Harker *et al.* (1990) recorded *I. acuminatum* from the earliest Campanian in the Western Interior, U.S.A.. According to Costa & Davey (1992) *I. acuminatum* first occurs, or first becomes consistent in the Early Campanian. Large *Isabelidinium* species such as *I. amphiatum* have their first occurrence in the middle Campanian in the District of Mackenzie (McIntyre, 1975), whereas Costa & Davey (1992) report an Early Coniacian to Late Campanian range in and around Great Britain.

The presence of the single specimens of *Tanyosphaeridium variacalamus* and *S. aff. longifurcatum* indicate a pre-Campanian age. The single specimen occurrence suggests, however that the species may be reworked or that they have a slightly longer range in West Greenland.

***Aquilapollenites* 2. interval**

Age. Early Campanian

Definition. The interval is defined from immediately above the last occurrence of *Isabelidinium cf. acuminatum* to immediately below the first occurrence of *I. microarmum*. For further discussion see under the *I. microarmum* interval.

Thickness and distribution. The interval has been recorded above the *I. cf. acuminatum* interval in section HNH910816/1 Kan. (Encl. 13; 150 m) and below the *I. microarmum* interval in the following sections: well GGU 400702 Agat. (Encl. 4; 40 m), HNH910819/1 Aaff. (Encl. 15; 60 m) and HNH910813/1 Tun. (Encl. 11; 50 m).

Characteristic species. The interval is represented by a low diversity palynomorph assemblage. The most characteristic species is *Aquilapollenites* and the interval is only recognisable by overlying the *I. cf. acuminatum* or by underlying the *I. microarmum* interval.

***Isabelidinium microarmum* interval**

Age. Middle or ?Late Campanian

Definition. The interval is defined from the first occurrence of *Isabelidinium microarmum* to the last occurrence of *I. microarmum*.

Thickness and distribution. The interval has been recorded in the following sections: well GGU 400702 (Encl. 4; 20 m), HNH910819/1 Aaff. (Encl. 15; 37 m), HNH910813/1 Tun. (Encl. 11; 60 m), HNH910826/1 Agat. (Encl. 7; 35 m), 400701 Agat. (Encl. 3; 35 m), FGC900731/2 Agat. (Encl. 2; 50 m), GKP 92 4 Qilak. (Encl. 18; one sample), HNH910813/2 Tun. (Encl. 12; 95 m) and GKP 92 3 Qilak. (Encl. 17; one sample).

Characteristic species. The diversity in the present interval is high compared to the underlying *Aquilapollenites* 2 interval. The interval is characterised by the presence of *Isabelidinium microarmum*, *I. aff. ?amphiatum* large, *Laciniadinium arcticum*, *Odontochitina striatoperforata*, *Chatangiella* aff. *granulifera*, *Desmocysta plekta*, *Excosphaeridium* aff. *bifidum*, *Excosphaeridium* aff. *striolatum*, *Hystrichosphaeridium pulchrum* and a single specimen of *Palaeohystrichophora infusorioides*.

Comments. The stratigraphic positions of the *I. cf. acuminatum* interval and the *I. microarmum* interval are somewhat problematic as the two species *I. cf. acuminatum* and *I. microarmum* have not been recorded in clear stratigraphical succession. They only occur together in one sample (well GGU 400702, Fig. 4, Encl. 4) or in separate sections. In the lower part of section HNH910816/1 Kan. (Fig. 3, Fig. 4, Encl. 13) the species *I. cf. acuminatum* occurs, whereas it has not been recorded in the samples from the 150 m thick section above (here named the *Aquilapollenites* 2 interval). The species *I. cf. acuminatum* has not been recorded from sediments in clear stratigraphical succession underlying sediments containing *I. microarmum* (Fig. 4).

It is here proposed that the *I. cf. acuminatum* interval is older than the *I. microarmum* interval, but whether the *Aquilapollenites* 2 interval should be defined as an interval or included in the upper part of the *I. cf. acuminatum* interval is still debatable.

It should be noted that the species *I. microarmum* has been observed reworked in a Paleocene dominated assemblage (Hansen, 1980). However, when reworked there are never more than one to three specimens.

Discussion. The species *Isabelidinium microarmum* has an Early Campanian to middle Maastrichtian range in the District of Mackenzie (McIntyre, 1975).

Ioannides (1986) recorded *I. microarmum* from questionable Maastrichtian strata in Arctic Canada. According to Costa & Davey (1992) *I. microarmum* does not seem to persisted beyond the end of the Campanian in the North Sea region. According to McIntyre (1975) large *Isabelidinium* species as such *I. amphiatum* have their first occurrence in the middle Campanian in the District of Mackenzie, whereas Costa & Davey (1992) report an Early Coniacian to Late Campanian range from the North Sea region.

According to Clarke & Verdier (1967) the species *Odontochitina striatoperforata* is a taxonomic junior synonym of *O. costata*, which has its last occurrence in the earliest Maastrichtian (Costa & Davey, 1992).

The presence of the species *Palaeohystrichophora infusorioides* and *Chatangiella* aff. *granulifera* indicates a pre-Maastrichtian age (Costa & Davey, 1992). A middle Campanian age is proposed for the *I. microarmum* interval on the basis of Costa & Davey's (1992, p. 126) comments that *Isabelidinium* species are common and diverse in the Early Campanian but much less in the Late Campanian.

***Cerodinium speciosum* interval**

Age. Late Paleocene

Definition. The interval is defined from the first occurrence of *Cerodinium speciosum* subsp. ?*glabrum*.

Thickness and distribution. The interval is only recorded in the uppermost sample of section GKP 92 3 Qilak. (Encl. 17). The lowermost sample in the section represents the

Early or ?middle Campanian *I. miroarmum* interval, and the approximately 200 metres of strata between the two samples is barren of dinoflagellate cysts.

Characteristic species. The interval is characterised by the presence of the species *Cerodinium speciosum* subsp. *?glabrum*, *Phelodinium kozlowskii* and *Glaphyrocysta* sp.. The presence of the pollen species *Wodehousia spinata* indicates reworking.

Discussion. The species *Cerodinium speciosum* dates the sample to a Late Paleocene age. In and around Great Britain the range of the species correlates with NP5 (Powell, 1992). The pollen species *Wodehousia spinata* has a rather restricted stratigraphical range from the latest Maastrichtian to the earliest Paleocene (Nichols & Brown, 1992) which indicates that the species may be reworked in the present sediments.

It is unlikely that the 200 metres of strata between the sample containing *I. microarmum* and the sample with *C. speciosum* represents the entire Upper Campanian to Upper Paleocene stratigraphic column and an unconformity is therefore proposed somewhere in the strata between the two samples.

COMPARISON WITH PREVIOUSLY REPORTED MACROFOSSIL AGES

The presence of ammonites of Santonian, Campanian and Maastrichtian ages in central Nuussuaq (Fig. 4) was recorded by Birkelund (1965; Fig. 2, Table 1) in her monograph on Upper Cretaceous ammonites from West Greenland. The ammonite record was correlated with other macrofossil records by Rosenkrantz & Pulvertaft (1969) in their review of Cretaceous–Tertiary stratigraphy and tectonics in West Greenland.

Birkelund (1965) recorded ammonites indicating the presence of Santonian deposits at the Nordre Baculiteskløft locality in Agatdalen. This locality is very close to the well GGU 400704; in the present study a late Santonian age (the *Chatangiella* aff. *ditissima* interval) is proposed. From the Scaphitesnæsen locality Birkelund (1965) recorded ammonites of Early Campanian age. The section FGC900731/2 Agat. covers the same locality and the dinoflagellate cysts assemblage from Scaphitesnæsen, suggests an Early or ?middle Campanian age.

The ammonites recorded from Ilugissoq and Tunoqqu indicate an Early Santonian age (Birkelund, 1965). The dinoflagellate cyst assemblage from the section GKP 92 V 2 Qilak., which is very close to Birkelund's (1965) Ilugissoq section, suggests a Late Santonian age. The studied sections from Tunoqqu contain dinoflagellate cysts suggesting an Early or ?middle Campanian age.

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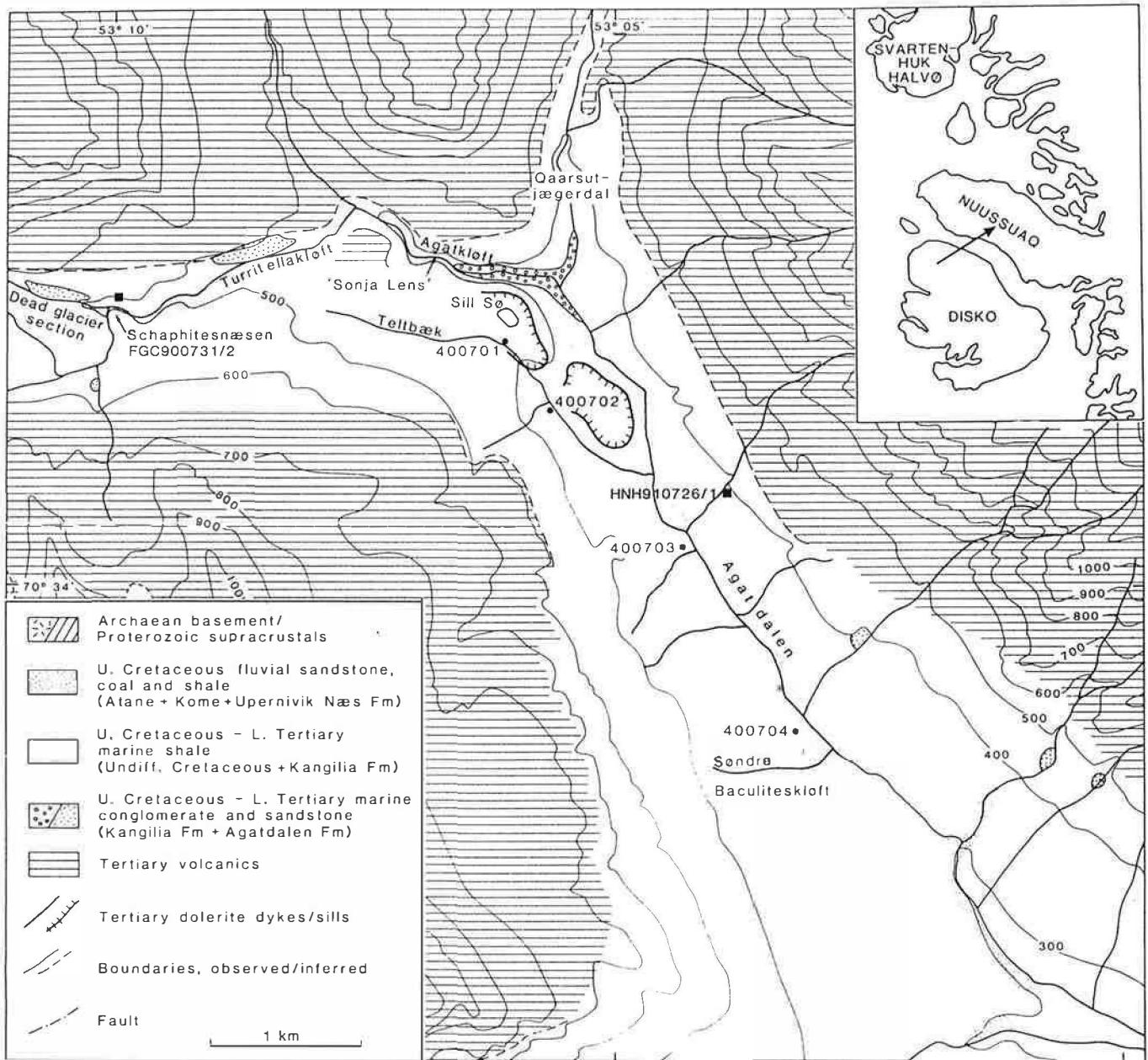


Fig. 1. Location of the examined outcrops and subsurface sections in the Agatdalen area that yielded dinoflagellate cysts.

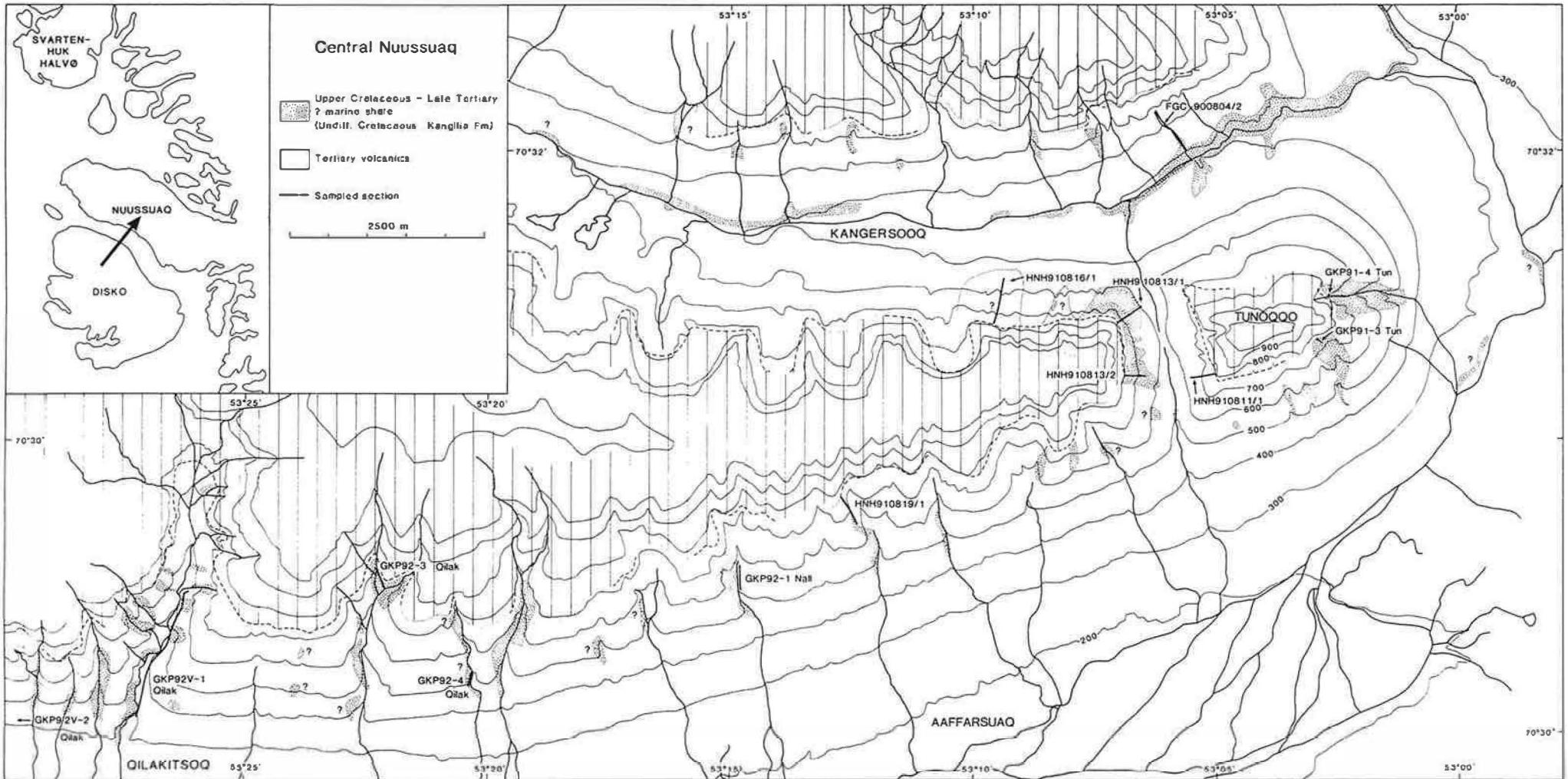


Fig. 2. Location of the examined outcrops and subsurface sections south of the Agatdalen area that yielded dinoflagellate cysts.

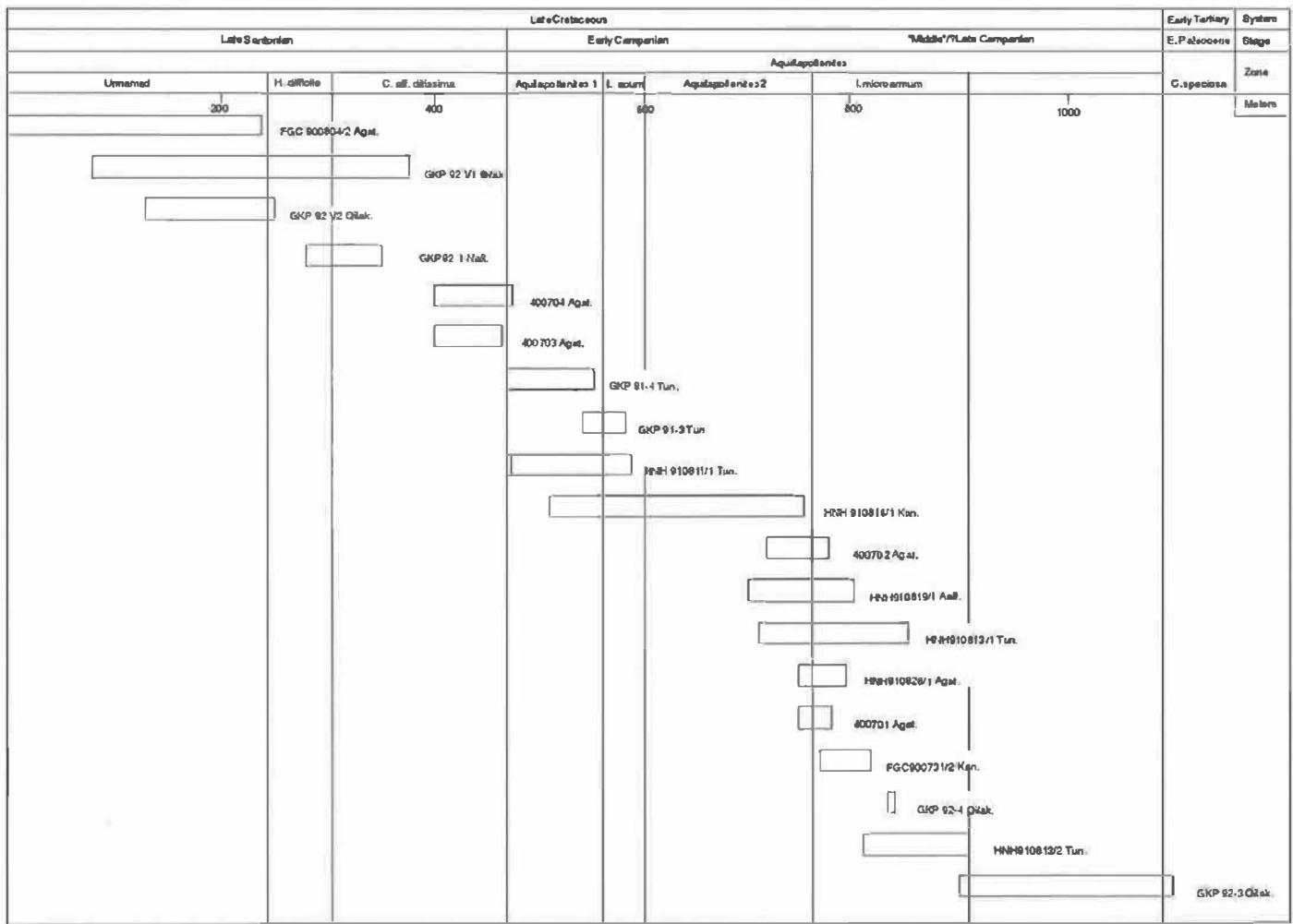


Fig. 3. Stratigraphical correlation of the sections that yielded dinoflagellate cysts . Kan.=Kangerlussuaq, Agat.=Agatdal, Qilak.=Qilakitsoq, Nall.=Nalluarissat, Tun.=Tunnoqqo, Aaff.=Aaffarsuaq.

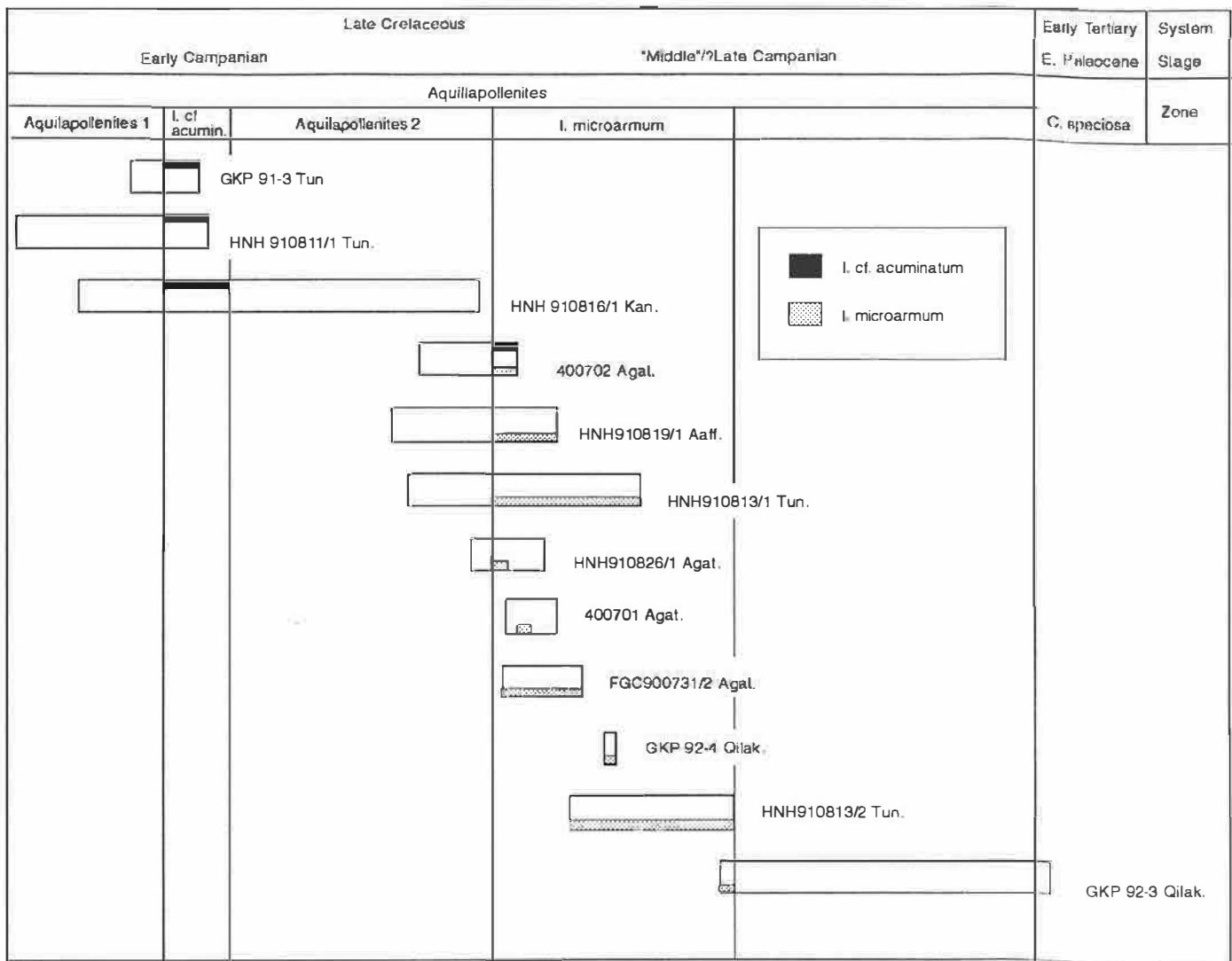


Fig. 4 Stratigraphic occurrence between the species *I. cf. acuminatum* and *I. microarmatum*.

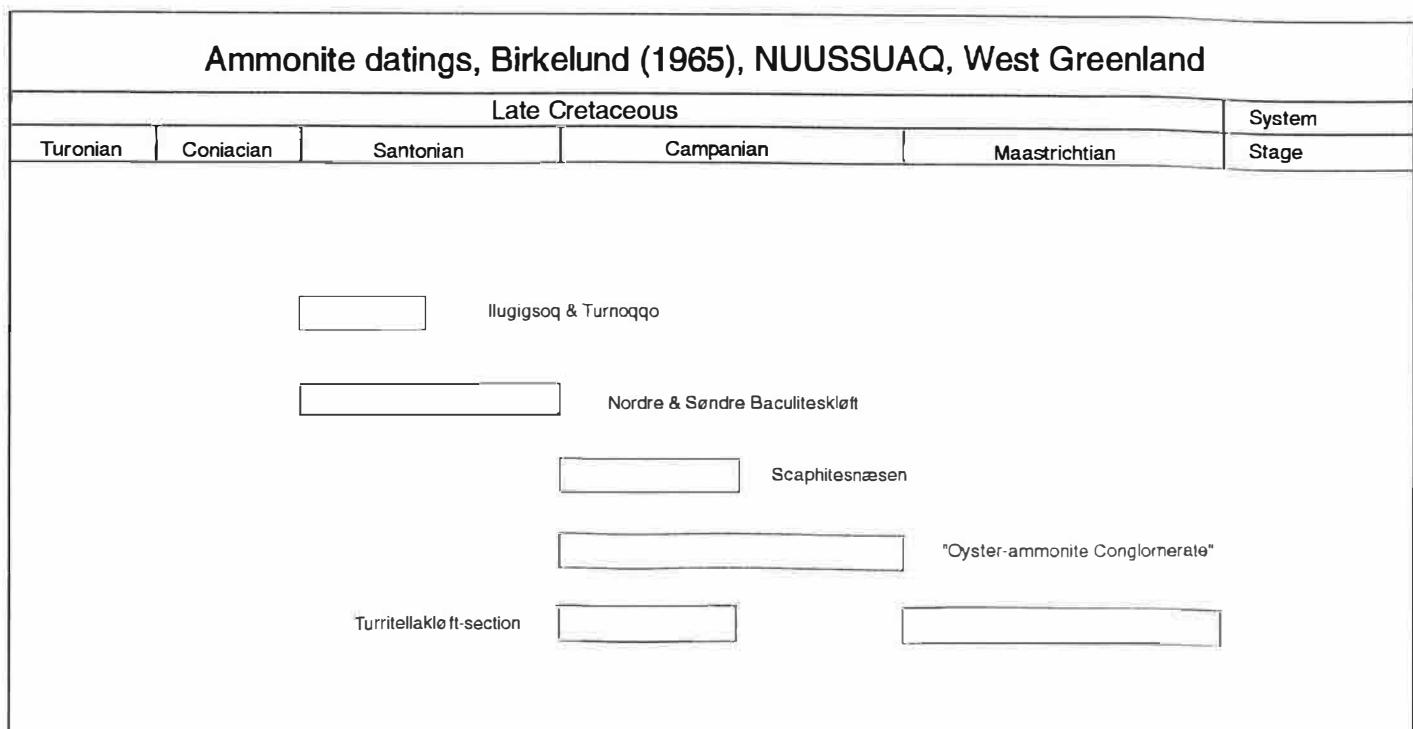


Fig. 5. Previous ammonite datings obtained from Birkelund (1965).

S I S B / F R K E

- 66 -

ENCL: B GKP 91-4 TUNORQO (735 - 649M)

ANALYSTS: 1994 HN-H
DATE: 22/3

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SIS BJERKE

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ANALYSTS: HN-H
DATE: 22/3

SIS BURKE

ENCL: 11 HNH 91-8-13/1 TU (770 - 6350)

ANALYSIS: 1994 HN-H
DATE: 3/3

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S I S B U F F K E

—SIS BURKE

S I S B JÆRKE

-EGU-

ENCL:	14	FGC 90-8-04/2 KA (671 - 427m)	ANALYSTS: 1994 HN-H DATE: 9/3
CRETACEOUS		S Y S T E M	
?SANTONIAN		STAGE	
UNAMED		Z O N E	
		D E P T H	L I T H O L O G Y
		REF:	S A M P L E S
			Chatangiella sp. Chonet cyst Biscaccate pollen
	+366568	700	
	+366566	650	■
	+366566	550	
	+366563	500	
	+366561	450	
	+366560	400	

ENCL: 15 HNH 91-8-19/1 AA (844 - 744m)

ANALYSTS: 1994 HN-H
DATE: 1/3

(66)

ANALYSTS: 1994 HN-H
DATE: 24/3

-66-

S I S B JØRKE

ENCL: 17 GKP 92-3 QLAKIT (870 - 66AM)

ANALYSTS: 1994 HN-H
DATE: 4/3

-GGU-

S I S B U J E K T

ENCL: 18 GKP 92-4 QILAKIT (580 - 580m)

**ANALYSTS: 1994 HN-H
DATE: 4/3**

-66-

S I S BJAERKE

ENGL: 19

GKP 92V 1 QTI AKT (570 - 284m)

ANALYSTS: 1994 HN-H
DATE: 7/3

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S I S BJÆRKE

ENCL: 20

ENEL: 20 GKP 92 V 2 QILAK (484 - 361m)

ANALYSTS: 1994 HN-H
DATE: 22/3

-68-

- 515 BWERKE -

Index of Figured Species

Dinoflagellate cysts

<i>Albertia</i> sp.	Pl. 1;	Fig.	1.
<i>Batioladinium jaegerii</i>	Pl. 5;	Figs	8-9.
<i>Chatangiella</i> aff. <i>spectabilis</i>	Pl. 3;	Fig.	10-12.
<i>Chatangiella</i> sp. with short spines	Pl. 2;	Fig.	5.
<i>Chatangiella granulifera</i>	Pl. 1;	Figs	8-9.
<i>Chatangiella ditissima</i>	Pl. 2;	Figs	1-4.
<i>Circulodinium distinctum</i>	Pl. 6;	Fig.	1.
<i>Circulodinium</i> cf. <i>?distinctum</i>	Pl. 6;	Figs	2-5.
Chorat cyst	Pl. 7;	Figs	10.
<i>Cribroperidinium?</i> sp.	Pl. 6;	Figs	11-12.
<i>Desmocysta plekta</i>	Pl. 5;	Figs	10-12.
Dinocyst sp. 10 HNH	Pl. 7;	Figs	1-2.
<i>Diconodinium?</i> sp. HNH	Pl. 2;	Figs	6.
<i>Dinogymnium?</i> sp.	Pl. 8;	Fig.	2.
<i>Exochosphaeridium</i> aff. <i>bifidum</i>	Pl. 7;	Figs	3-6.
<i>Exochosphaeridium</i> aff. <i>striolatum</i>	Pl. 7;	Figs	7-8.
<i>Fromea nicosia</i>	Pl. 8;	Fig.	3.
<i>Heterosphaeridium difficile</i>	Pl. 7;	Fig.	9.
<i>Hystrichosphaeridium?</i> sp.	Pl. 7;	Fig.	11.
Hyalin sphaeromorph	Pl. 6;	Fig.	7.
<i>Isabelidinium</i> aff. <i>acuminatum</i>	Pl. 4;	Figs	1-12.
<i>Isabelidinium</i> <i>microarmum</i>	Pl. 3;	Figs	1-5.
<i>Isabelidinium</i> aff. <i>?amphiatum</i> large	Pl. 1;	Figs	1-12.
<i>Isabelidinium</i> sp. 14 HNH	Pl. 3;	Figs	7-9.
<i>Laciniadinium arcticum</i>	Pl. 2;	Figs	7-9.
<i>Odonthochitina striatoperforata</i>	Pl. 5;	Figs	1-4.
<i>Oligosphaeridium</i> sp.	Pl. 7;	Fig.	12.
<i>Palaeocystodinium</i> aff. <i>golzowenze</i>	Pl. 5;	Fig.	7.
<i>Spinidinium echinoideum</i>	Pl. 2;	Figs	10-12.
<i>Tanyosphaeridium variecalamus</i>	Pl. 8;	Fig.	1.
<i>Trithyrodinium</i> aff. <i>suspectum</i>	Pl. 6;	Figs	9-10.
<i>Xenascus</i> aff. <i>perforatus</i>	Pl. 5;	Figs	5-6.
Acritarchs			
Acritarch sp. 1 HNH	Pl. 8;	Fig.	5.
<i>Schizocystia</i> sp.	Pl. 8;	Fig.	6.
<i>Tasmanites</i> sp.	Pl. 8;	Fig.	4.
Spores			
Monolete spore	Pl. 8;	Fig.	7.
Pollen			
<i>Aquilapollenites</i>	Pl. 8;	Figs	8-9.
Fungal			
Fungal sp. 1 HNH	Pl. 8;	Fig.	10-12.

Plate 1 Central Nuussuaq

Fig. 1. *Isabelidinium* aff. *?amphiatum* large, GGU 366523-3, FGC900731/2, 47.5-92.9; LVR 1.1889; MI 1333

Fig. 2. *Isabelidinium* aff. *?amphiatum* large, GGU 366523-3, FGC900731/2, 47.5-92.9; LVR 1.1890; MI 1333

Fig. 3. *Isabelidinium* aff. *?amphiatum* large, GGU 360732-6, HNH910819/1, 33.9-103.3; LVR 1.1970; MI 1400

Fig. 4. *Isabelidinium* aff. *?amphiatum* large, GGU 360718-4, HNH910816/1, 25.0-101.9; LVR 1.3091; MI 2338

Fig. 5. *Isabelidinium* aff. *?amphiatum* large, GGU 360717-5, HNH910816/1, 51.3-113.7; LVR 1.1950; MI 1385

Fig. 6. *Isabelidinium* aff. *?amphiatum* large, GGU 360717-5, HNH910816/1, 22.3-98.8; LVR 1.1944; MI 1379

Fig. 7. *Isabelidinium* aff. *?amphiatum* large, GGU 366523-6, FGC900731/2, 19.4-103.1; LVR 1.1884; MI 1330

Fig. 8. *Isabelidinium* aff. *?amphiatum* large, GGU 360717-9, HNH910816/1, 39.0-99.6; LVR 1.1945; MI 1380

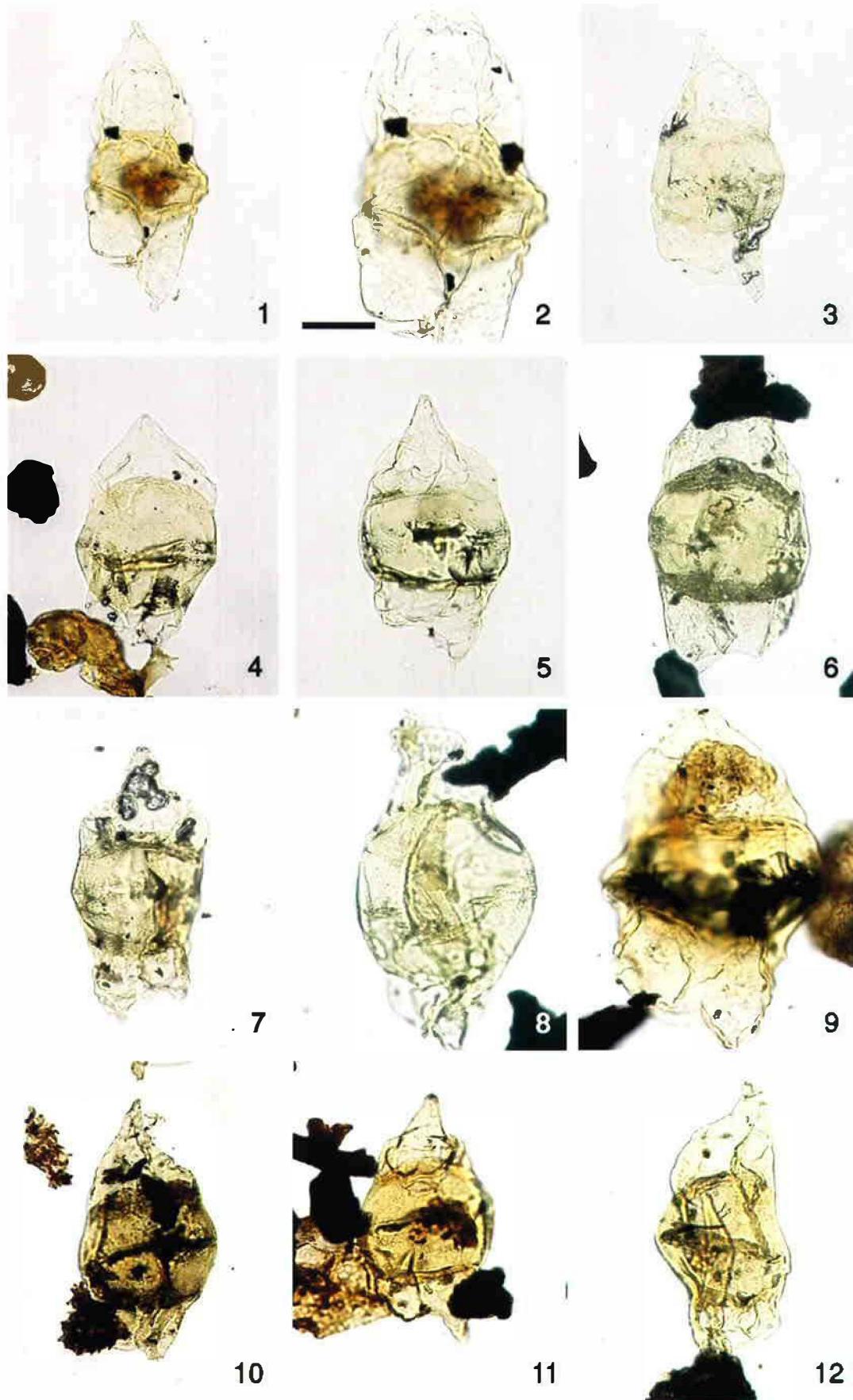
Fig. 9. *Isabelidinium* aff. *?amphiatum* large, GGU 366523-4, FGC900731/2, 51.0-110.5; LVR 1.1892; MI 1334

Fig. 10. *Isabelidinium* aff. *?amphiatum* large, GGU 400702-12-3, 49.7-98.7; LVR 1.3088; MI 2336

Fig. 11. *Isabelidinium* aff. *?amphiatum* large, GGU 400702-12-9, 53.6-106.4; LVR 1.3090; MI 2337

Fig. 12. *Isabelidinium* aff. *?amphiatum* large, GGU 400702-12-3, 51.4-107.7; LVR 1.3087; MI 2335

CENTRAL NUUSSUAQ



20 μm

Plate 2 Central Nuussuaq

Fig. 1. *Chatangiella ditissima*, GGU 400585-8, GKP 92-V1 Qilak., 39.9-105.4; LVR 1.1988; MI 1415

Fig. 2. *Chatangiella ditissima*, GGU 360718-8, HNH910816/1, 52.6-105.9; LVR 1.3102; MI 2349

Fig. 3. *Chatangiella ditissima*, GGU 351822-3, GKP 91-3 Tun., 35.1-106.0; LVR 1.3107; MI 2354

Fig. 4. *Chatangiella ditissima*, GGU 351822-3, GKP 91-3 Tun., 38.7-103.6; LVR 1.3108; MI 2355

Fig. 5. *Chatangiella* sp. with short spines, GGU 366523-3, FGC900731/2, 48.2-107.4; LVR 1.1893; MI 1335

Fig. 6. *Diconodinium?* sp. HNH, GGU 360732-4, HNH910819/1, 52.1-109.3; LVR 1.1972; MI 1401

Fig. 7. *Laciniadinium arcticum*, GGU 369287-4, HNH910813/1, 24.9-101.4; LVR 1.1906; MI 1347

Fig. 8. *Laciniadinium arcticum*, GGU 369287-4, HNH910813/1, 26.5-100.5; LVR 1.1907; MI 1348

Fig. 9. *Laciniadinium arcticum*, GGU 351828-3, GKP 91-4 Tun., 53.1-107.3; LVR 1.3111; MI 2357

Fig. 10. *Spinidinium echinoideum*, GGU 366523-3, FGC900731/1, 20.1-102.8; LVR 1.1874; MI 1324

Fig. 11. *Spinidinium echinoideum*, GGU 369287-4, HNH910813/1, 37.3-113.2; LVR 1.1914; MI 1354

Fig. 12. *Spinidinium echinoideum*, GGU 369287-4, HNH910813/1, 43.4-113.1; LVR 1.1915; MI 1355

CENTRAL NUUSSUAQ

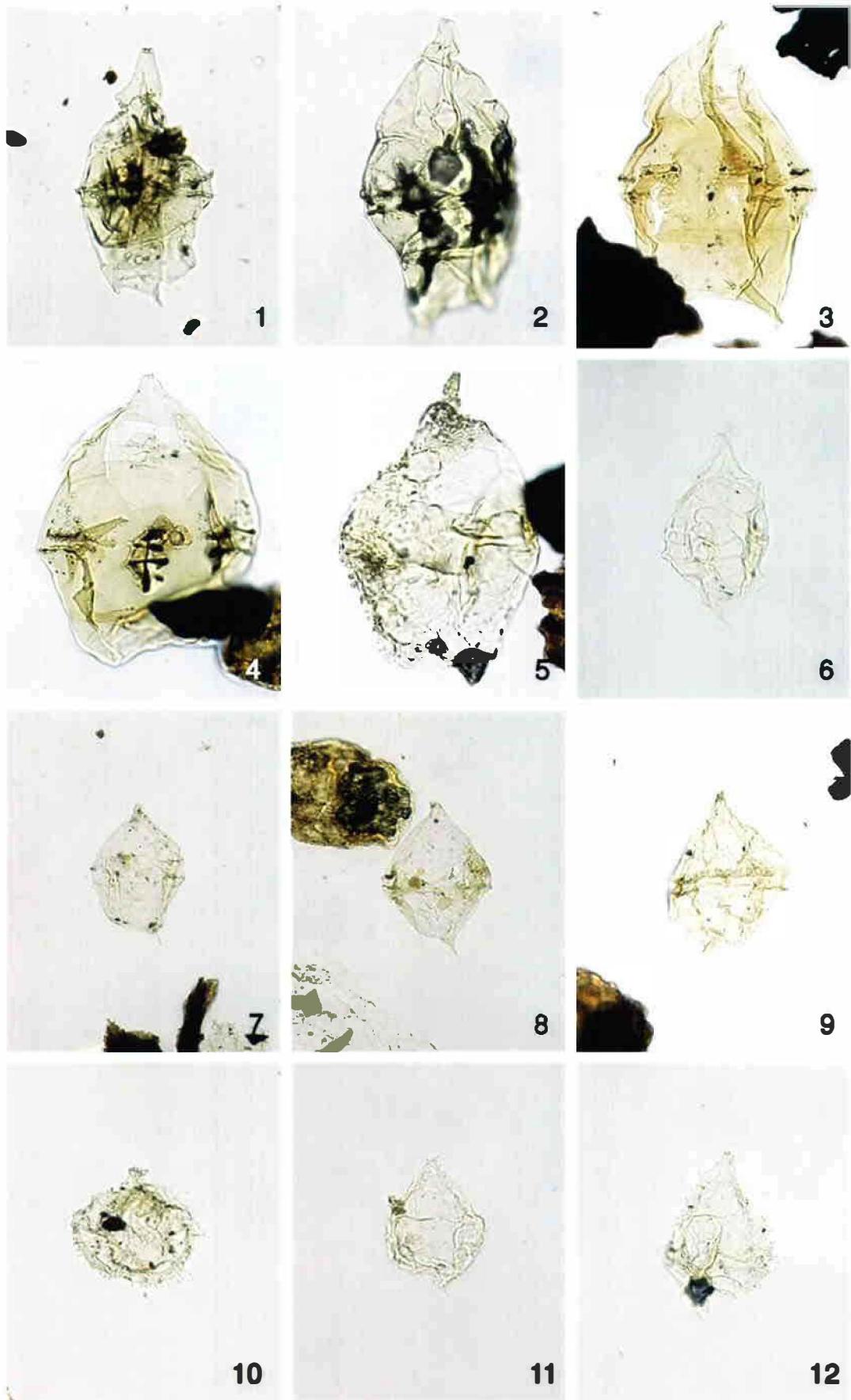


Plate 3 Central Nuussuaq

Fig. 1. *Isabelidinium microarmum*, GGU 360731-10, HNH910819/1, 48.7-107.4; LVR 1.1962; MI 1394

Fig. 2. *Isabelidinium microarmum*, GGU 360732-9, HNH910819/1, 25.5-101.8; LVR 1.1966; MI 1396

Fig. 3. *Isabelidinium microarmum*, GGU 360732-10, HNH910819/1, 29.6-103.2; LVR 1.1968; MI 1398

Fig. 4. *Isabelidinium microarmum*, GGU 360732-6, HNH910819/1, 33.3-114.3; LVR 1.1969; MI 1399

Fig. 5. *Isabelidinium microarmum*, GGU 360717-4, HNH910816/1, 38.6-97.7; LVR 1.3084; MI 2332

Fig. 6. *Albertia* sp., GGU 369287-4, HNH910813/1, 30.7-99.8; LVR 1.1911; MI 1351

Fig. 7. *Isabelidinium?* sp. 14 HNH, GGU 369284-4, HNH910813/1, 21.5-98.2; LVR 1.1899; MI 1340

Fig. 8. *Isabelidinium?* sp. 14 HNH, GGU 369284-4, HNH910813/1, 50.4-113.6; LVR 1.1902; MI 1343

Fig. 9. *Isabelidinium?* sp. 14 HNH, GGU 360731-7, HNH910819/1, 52.7-107.9; LVR 1.1961; MI 1393

Fig. 10. *Chatangiella* aff. *spectabilis*, GGU 400577-4, GKP , 27.6-96.8; LVR 1.1973; MI 1402

Fig. 11. *Chatangiella* aff. *spectabilis*, GGU 400577-4, GKP , 27.6-96.8; LVR 1.1974; MI 1402

Fig. 12. *Chatangiella* aff. *spectabilis*, GGU 400577-4, GKP , 43.4-105.8; LVR 1.1975; MI 1404

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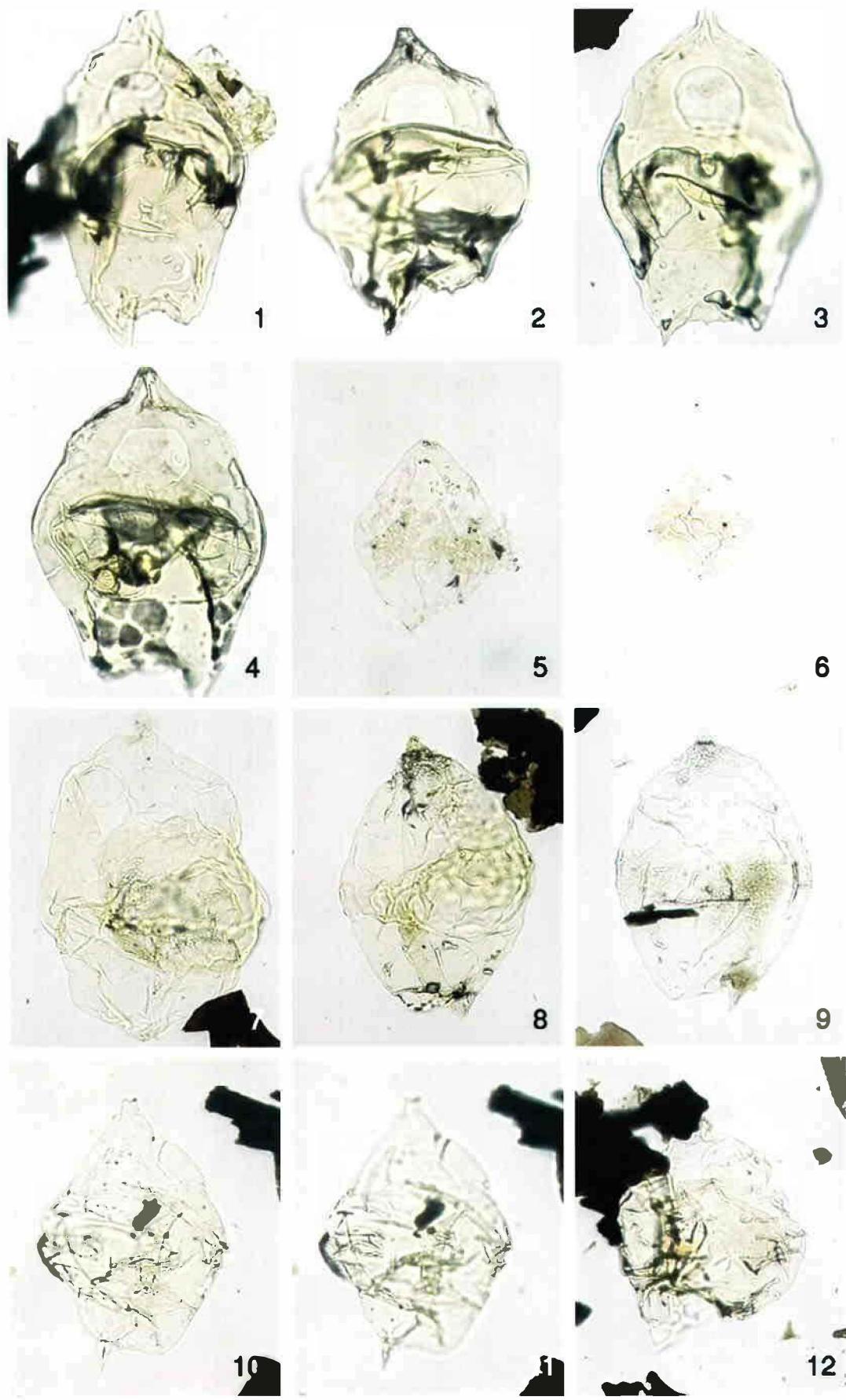
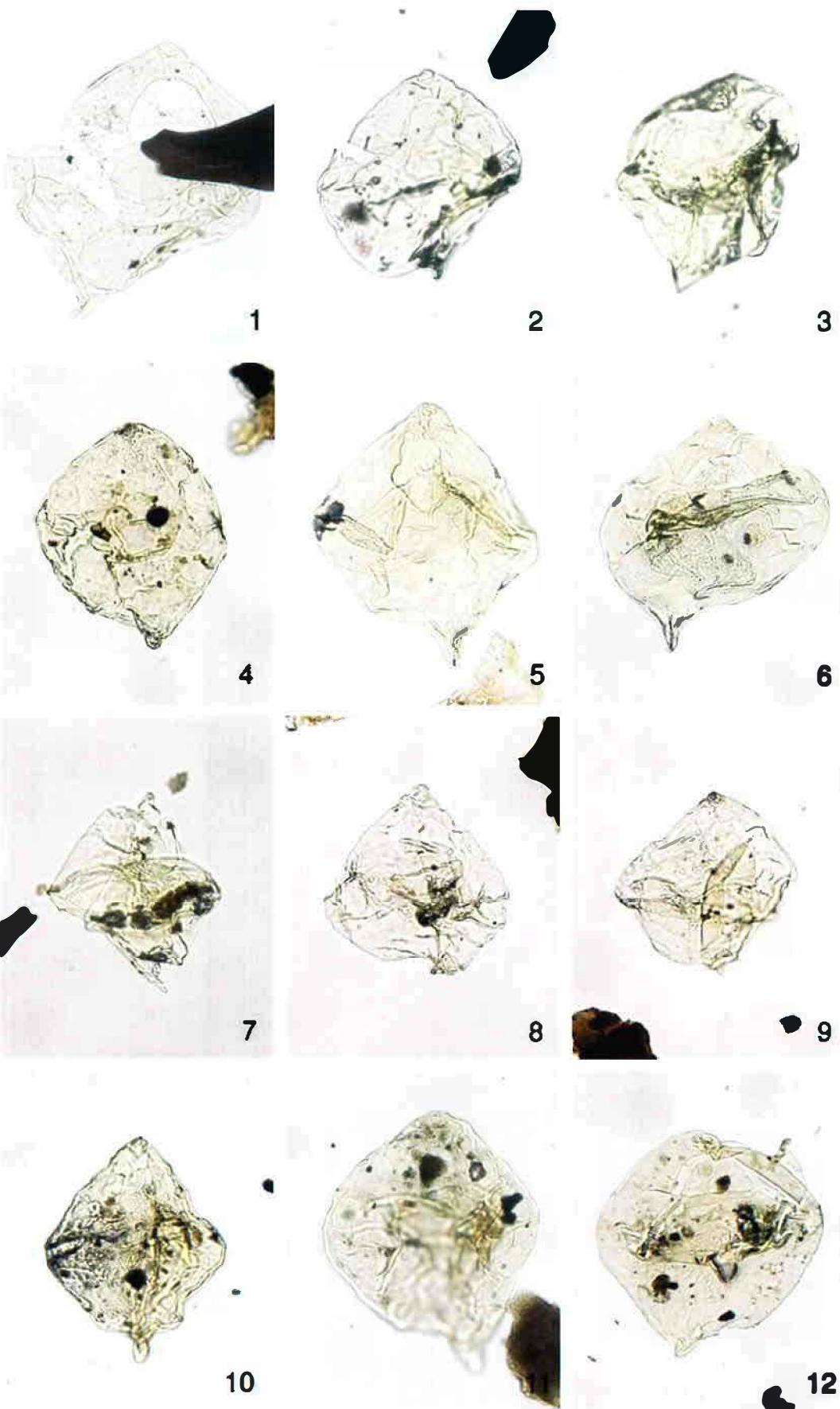


Plate 4 Central Nuussuaq

- Fig. 1. *Isabelidinium cf. acuminatum*, GGU 360717-4, HNH910816/1, 25.5-95.7; LVR 1.1946; MI 1381
- Fig. 2. *Isabelidinium cf. acuminatum*, GGU 360717-9, HNH910816/1, 44.5-100.3; LVR 1.1947; MI 1382
- Fig. 3. *Isabelidinium cf. acuminatum*, GGU 360717-9, HNH910816/1, 37.0-96.2; LVR 1.1948; MI 1383
- Fig. 4. *Isabelidinium cf. acuminatum*, GGU 400702-12-9, 38.9-103.0; LVR 1.3081; MI 2329
- Fig. 5. *Isabelidinium cf. acuminatum*, GGU 351824-3 GKP 91-3 Tun., 20.1-102.3; LVR 1.3082; MI 2330
- Fig. 6. *Isabelidinium cf. acuminatum*, GGU 351824-3 GKP 91-3 Tun., 27.8-115.0; LVR 1.3083; MI 2331
- Fig. 7. *Isabelidinium cf. acuminatum*, GGU 369276-4, HNH910811/1, 53.1-100.3; LVR 1.3085; MI 2333
- Fig. 8. *Isabelidinium cf. acuminatum*, GGU 400702-10-7, 41.0-103.0; LVR 1.3076; MI 2324
- Fig. 9. *Isabelidinium cf. acuminatum*, GGU 400702-10-5, 32.7-103.7; LVR 1.3077; MI 2325
- Fig. 10. *Isabelidinium cf. acuminatum*, GGU 400702-12-9, 24.5-109.3; LVR 1.3079; MI 2327
- Fig. 11. *Isabelidinium cf. acuminatum*, GGU 400702-12-9, 23.8-108.2; LVR 1.3078; MI 2326
- Fig. 12. *Isabelidinium cf. acuminatum*, GGU 400702-12-9, 33.0-110.4; LVR 1.3080; MI 2328

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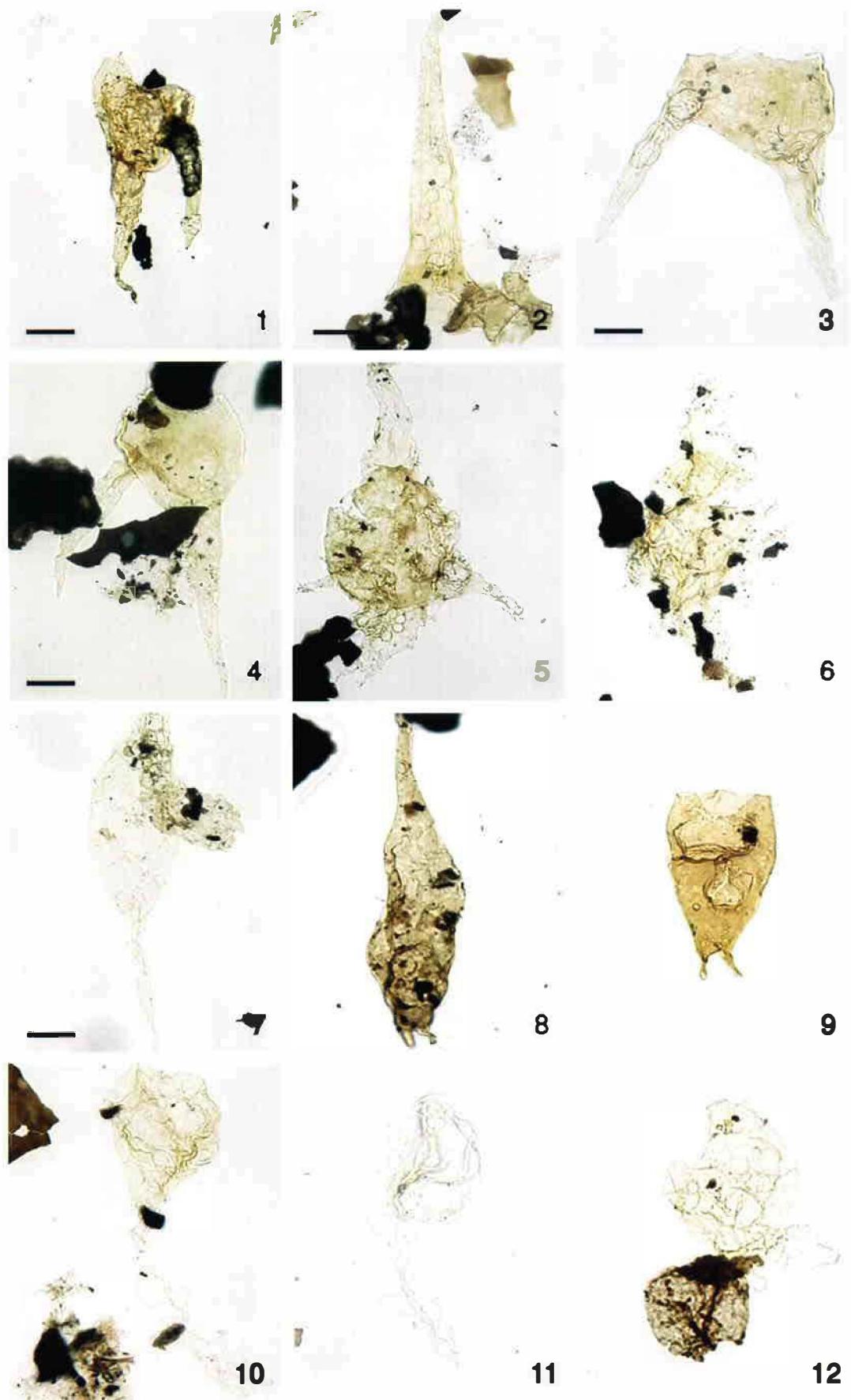


20 μm

Plate 5 Central Nuussuaq

- Fig. 1. *Odontochitina striatoperforata*, GGU 366523-4 FGC900731/2, 48.6-108.0; LVR 1.1885; MI 1331
- Fig. 2. *Odontochitina striatoperforata* GGU 369287-4, HNH910813/1, 55.5-101.8; LVR 1.1918; MI 1358
- Fig. 3. *Odontochitina striatoperforata* GGU 369287-6, HNH910813/1, 55.1-106.0; LVR 1.1927; MI 1365
- Fig. 4. *Odontochitina striatoperforata* GGU 360729-10, HNH910819/1; 25.4-110.2; LVR 1.1954; MI 1389
- Fig. 5. *Xenascus* aff. *perforatus* GGU 400585-6, GKP 92 V1 Qilak., 33.2-104.0; LVR 1.1985; MI 1413
- Fig. 6. *Xenascus* aff. *perforatus* GGU 400585-4, GKP 92 V1 Qilak., 40.6-107.0; LVR 1.1984; MI 1412
- Fig. 7. *Palaeocystodinium* aff. *golzowenze* GGU 400701-4-9, 31.8-93.0; LVR 1.3096; MI 2343
- Fig. 8. *Batioladinium jaegerii* GGU 351828-2, GKP 91 4 Tun., 43.6-95.5; LVR 1.3109; MI 2356
- Fig. 9. *Batioladinium jaegerii* GGU 351828-3, GKP 91 4 Tun., 23.8-113.5; LVR 1.3110; MI 2357
- Fig. 10. *Desmocysta plekta*, GGU 269284-4, ? 51.0-103.7; LVR 1.1903; MI 1344
- Fig. 11. *Desmocysta plekta* GGU 400585-4, GKP 92 V1 Qilak., 32.8-94.0; LVR 1.1980; MI 1408
- Fig. 12. *Desmocysta plekta* GGU 360718-8, HNH910816/1, 27.0-111.0; LVR 1.3104; MI 2351

CENTRAL NUUSSUAQ



20 μm

Plate 6 Central Nuussuaq

- Fig. 1. *Circulodinium distinctum*, GGU 360717-9, HNH910816/1, 55.3-97.8; LVR 1.1943; MI 1378
- Fig. 2. *Circulodinium* cf. *?distinctum*, GGU 400702-12-3, 34.7-104.2; LVR 1.3092; MI 2339
- Fig. 3. *Circulodinium* cf. *?distinctum*, GGU 400702-12-7, 34.1-111.2; LVR 1.3093; MI 2340
- Fig. 4. *Circulodinium* cf. *?distinctum*, GGU 400702-12-9, 35.0-108.8; LVR 1.3094; MI 2341
- Fig. 5. *Circulodinium* cf. *?distinctum*, GGU 400702-12-9, 25.5-105.8; LVR 1.3095; MI 2342
- Fig. 6. Hyalin sphaeromorph GGU 369287-4, HNH910813/1, 36.2-103.0; LVR 1.1904; MI 1345
- Fig. 7. Hyalin sphaeromorph GGU 369287-4, HNH910813/1, 45.1-109.0; LVR 1.1905; MI 1346
- Fig. 8. *Trityrodinium* aff. *suspectum* GGU 400585-8, GKP 92 V1 Qilak., 35.5-100.2; LVR 1.1989; MI 1416
- Fig. 9. *Trityrodinium* aff. *suspectum* GGU 400585-8, GKP 92 V1 Qilak., 35.5-100.2; LVR 1.1990; MI 1416
- Fig. 10. *Trityrodinium* aff. *suspectum*, GGU 366523-3 FGC900731/2, 42.8-102.0; LVR 1.1887; MI 1332
- Fig. 11. *Cribroperidinium?* sp., GGU 351822-3 GKP 91-3 Tun., 22.6-94.5; LVR 1.3105; MI 2352
- Fig. 12. *Cribroperidinium?* sp., GGU 351822-3 GKP 91-3 Tun., 47.8-96.6; LVR 1.3106; MI 2353

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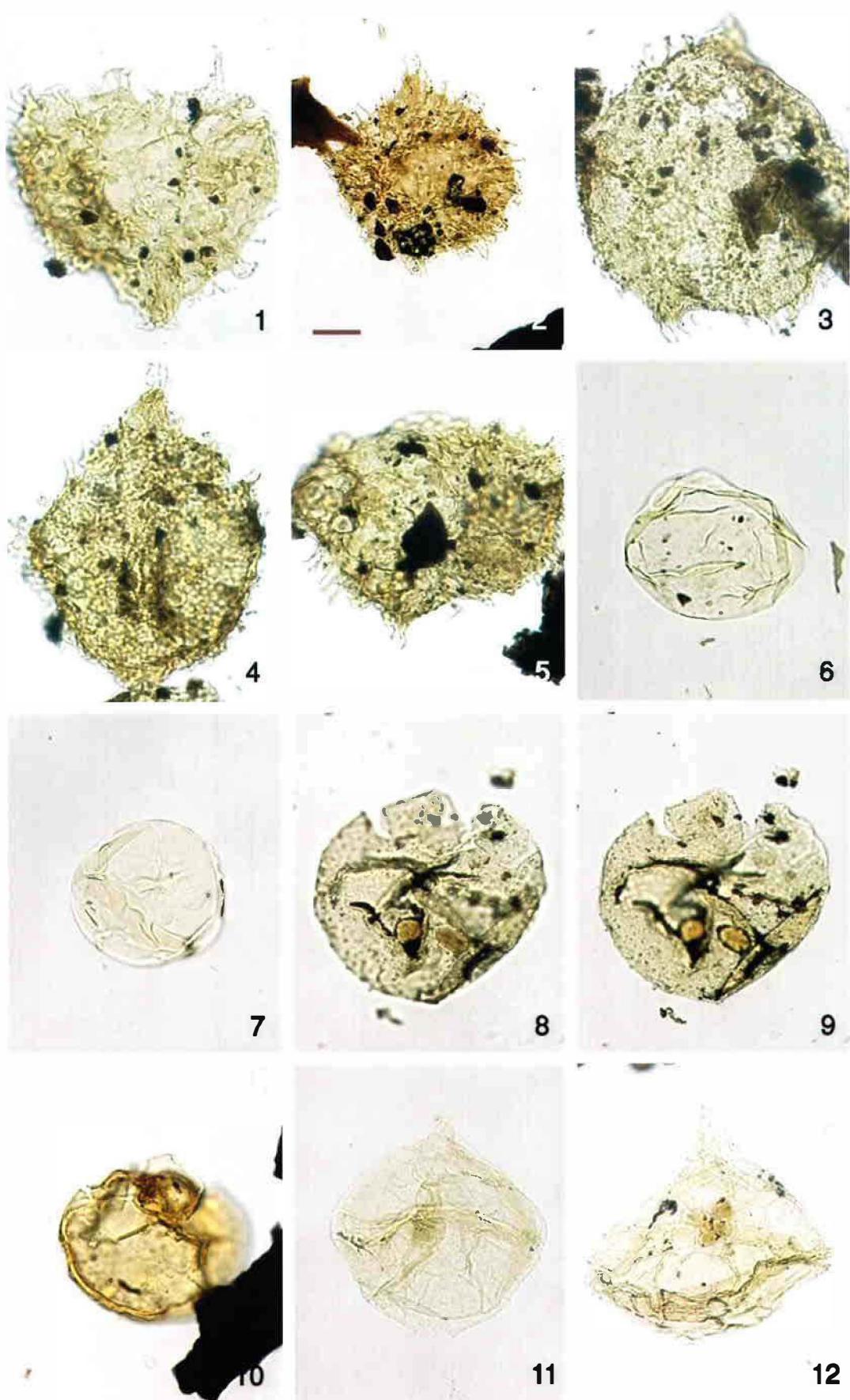


Plate 7 Central Nuussuaq

Fig. 1. Dinocyst sp. 10 HNH, GGU 366523-3 FGC900731/2, 29.2-122.6; LVR 1.1876; MI 1326

Fig. 2. Dinocyst sp. 10 HNH, GGU 366523-3 FGC900731/2, 29.2-122.6; LVR 1.1877; MI 1326

Fig. 3. *Exochosphaeridium* aff. *bifidum*, GGU 369287-4, HNH910813/1, 33.2-105.0; LVR 1.1913; MI 1352

Fig. 4. *Exochosphaeridium* aff. *bifidum*, GGU 369284-4, HNH910813/1, 22.8-106.4; LVR 1.1896; MI 1338

Fig. 5. *Exochosphaeridium* aff. *bifidum*, GGU 369284-6, HNH910813/1, 55.9-100.5; LVR 1.1898; MI 1339

Fig. 6. *Exochosphaeridium* aff. *bifidum*, GGU 369284-4, HNH910813/1, 18.9-108.9; LVR 1.1895; MI 1337

Fig. 7. *Exochosphaeridium* aff. *striolatum*, GGU 369287-6, HNH910813/1, 45.8-96.8; LVR 1.1930; MI 1368

Fig. 8. *Exochosphaeridium* aff. *striolatum*, GGU 369287-9, HNH910813/1, 47.4-96.9; LVR 1.1931; MI 1369

Fig. 9. *Heterosphaeridium difficile* GGU 400577-7, GKP 92 V1 Qilak., 30.0-112.3; LVR 1.1978; MI 1406

Fig. 10. Chorat cyst GGU 400577-9, GKP 92 V1 Qilak., 30.5-104.6; LVR 1.1979; MI 1407

Fig. 11. *Hystrichosphaeridium?* sp., GGU 351824-3 GKP 91-3 Tun., 49.4-103.7; LVR 1.3098; MI 2345

Fig. 12. *Oligosphaeridium* sp., GGU 360718-6, HNH910816/1, 36.4-110.8; LVR 1.3100; MI 2347

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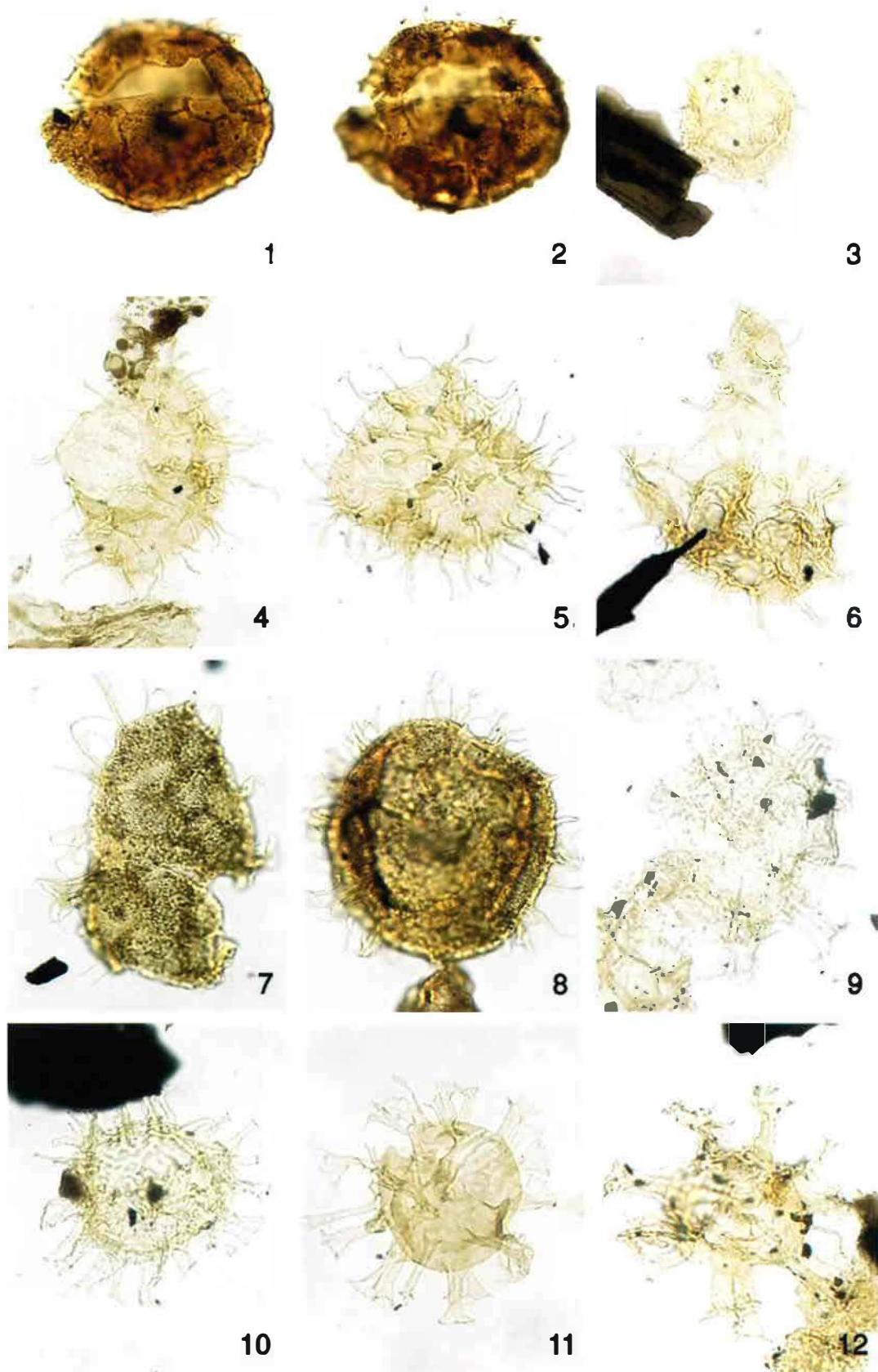
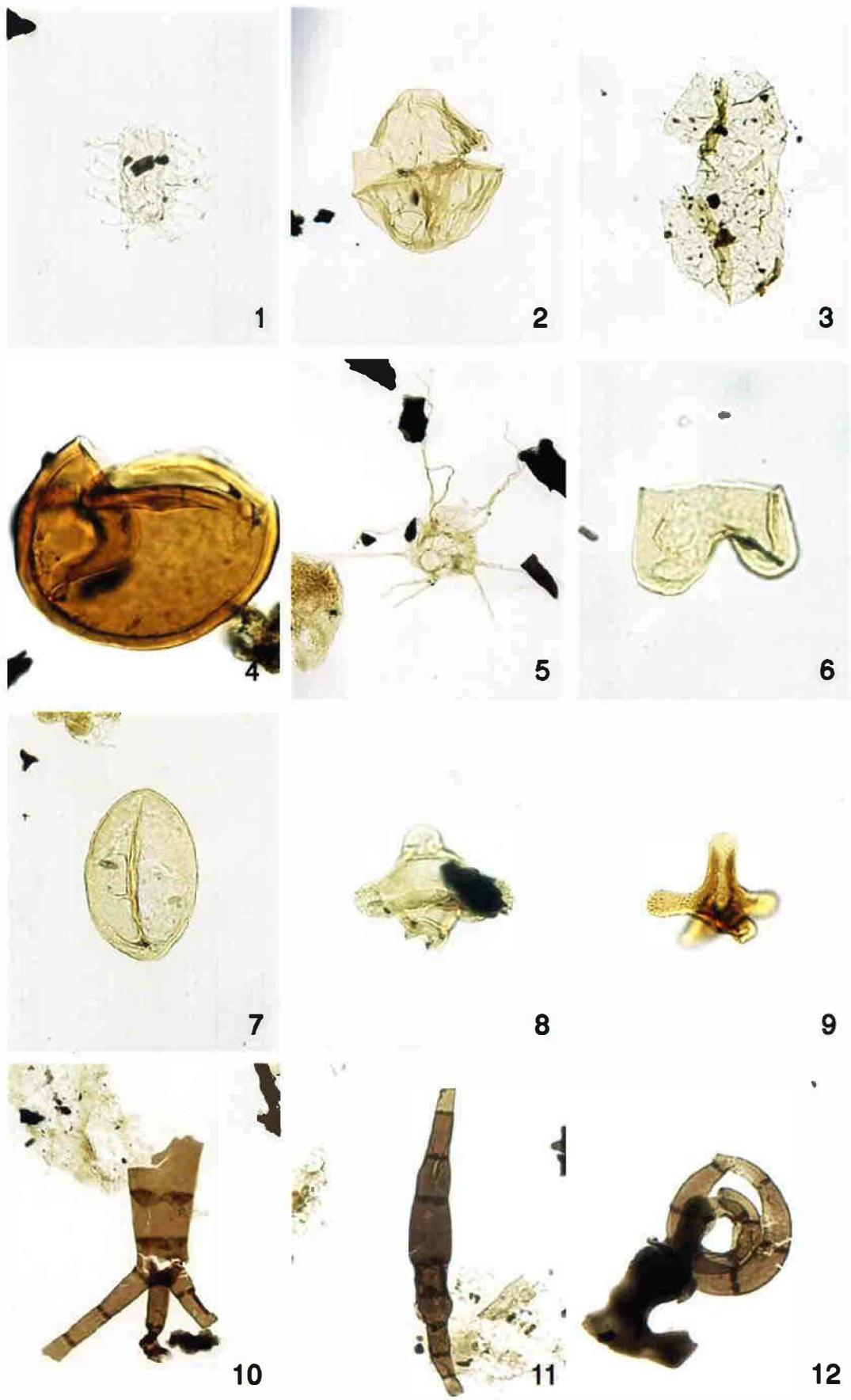


Plate 8 Central Nuussuaq

- Fig. 1. *Tanyosphaeridium variecalamus*, GGU 360729-4, HNH910819/1, 28.0-100.2; LVR 1.1953; MI 1388
- Fig. 2. *Dinogymnium?* sp., GGU 369287-4, HNH910813/1, 45.1-104.9; LVR 1.1916; MI 1356
- Fig. 3. *Fromea nicosia*, GGU 400702-10-7, 42.9-99.7; LVR 1.3075; MI 2323
- Fig. 4. *Tasmanites* sp., GGU 366523-6, FGC900731/2, 40.0-106.0; LVR 1.1894; MI 1336
- Fig. 5. Acritarch sp. 1 HNH, GGU 369287-4, HNH910813/1, 25.5-107.0; LVR 1.1908; MI 1349
- Fig. 6. *Schizocystia* sp., GGU 369276-4, HNH910811/1, 37.9-113.4; LVR 1.3099; MI 2346
- Fig. 7. Monolete spore, GGU 369287-4, HNH910813/1, 55.7-93.6; LVR 1.1920; MI 1360
- Fig. 8. *Aquilapollenites* sp., GGU 360717-5, HNH910816/1, 41.3-103.7; LVR 1.1949; MI 1384
- Fig. 9. *Aquilapollenites* sp., GGU 369284-4, HNH910813/1, 28.2-112.8; LVR 1.1900; MI 1341
- Fig. 10. Fungal sp. HNH, GGU 369287-4, HNH910813/1, 37.5-103.5; LVR 1.1912; MI 1353
- Fig. 11. Fungal sp. HNH, GGU 360717-4, HNH910816/1, 22.0-103.0; LVR 1.1935; MI 1372
- Fig. 12. Fungal sp. HNH, GGU 360717-4, HNH910816/1, 38.6-109.5; LVR 1.1936; MI 1373

CENTRAL NUUSSUAQ



20 μm

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