

Gold content of regional stream sediment samples from South Greenland

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A newly conducted analysis of 2192 stream sediment samples shows that anomalous gold values (25-850 ppb) are distributed over much of South Greenland. The results of the analysis for Au and As are presented and briefly commented.

Sampling and analysis

The stream sediment samples were collected during a regional reconnaissance geochemical uranium exploration programme (Syduran) covering South Greenland (Armour-Brown et al., 1982). About 2300 samples were collected at an average density of 1 sample per 6 km². The samples were dry-sieved in the laboratory and the <0.1 mm grain size fraction was analysed at Risø National Laboratory, Denmark, for uranium by delayed neutron counting, and for 17 major and trace elements by radio-isotope energy dispersive X-ray fluorescence (Kunzendorf, 1979). The results have been presented as geochemical element distribution maps at a scale of 1:1,000,000 (Armour-Brown et al., 1982; Olesen, 1984). The evaluation of the results, mainly with respect to uranium, is found in Armour-Brown et al. (1983), Armour-Brown & Olesen (1984) and Steenfelt & Armour-Brown (1988).

The remains of the fine fraction of the samples have been used for instrumental neutron activation analysis at Actlabs, Canada, where gold plus 33 other trace elements were determined. The analyses were conducted on 7 gram or even on 1 gram aliquots (instead of the now recommended 30 grammes) due to very limited amount of material available. With such small amounts of material the nugget effect may seriously influence the gold determinations and the results must be considered as a rough guide only. In the evaluation of the results the emphasis should be placed on the geographical distribution of gold concentration rather than on the magnitude of individual values.

Of the 2192 samples analysed, 475 (=22.7%) yielded gold values above the detection limit of 5 ppb, and 1336 (=55.2%) yielded arsenic values above 2 ppm. The frequency distributions of the detected values are given in the table:

1850 10 1/40

Au ppb	number	As ppm nur	nber
5-25	441 of the second	2-30 12	209
26-50	14 566 5188	31-60	82
51-75	5	61-90	16
76-100	3	91-120	10
>100	12	>120	18

Summary of the geology

The geological map at 1:500,000 scale by the Geological Survey of Greenland (Allaart, 1975) enclosed with this report serves to summarize the geology of South Greenland. A description of this map is in preparation (Kalsbeek et al., in prep.). Parts of the region are covered by geological maps at 1:100,000 scale published by GGU.

The northernmost part of South Greenland is underlain by Archaean basement whereas the entire southern part belongs to the early Proterozoic Ketilidian mobile belt. The Ketilidian is thought to represent juvenile Proterozoic crust accreted to the Archaean continent (Patchett & Bridgwater, 1984; Kalsbeek & Taylor, 1985). A mid-Proterozoic alkaline igneous province (Gardar province) pierces the northern part of the Ketilidian belt and the southernmost part of the Archaean basement.

Allaart (1976) divided the Ketilidian mobile belt into 4 structural units. From north to south these are: 'border zone', 'granite zone', 'folded migmatite zone', and 'flat-lying migmatite complex'. The boundaries between these units are shown on the gold and arsenic distribution maps of this report.

The 'border zone', forming the southernmost part of the Archaean basement, consists of tonalitic to granodioritic gneisses and basic supracrustal sequences. The most important Archaean supracurstal occurrence, known as the Tartoq group, contains horizons of quartz banded iron-formation (Appel, 1984). To the east and southeast the Archaean basement is overlain by Proterozoic metasedimentary and basic metavolcanic rocks (Midternæs and Grænseland) which are progressively deformed and metamorphosed towards the south. The Proterozoic supracrustals also contain banded iron-formation (Appel, 1974).

The 'granite zone' is dominated by numerous intrusions of granitic (s.l.) batholiths with scattered inclusions of supracrustal rocks. Minor plutons of basic rocks are associated with the granites. According to isotopic age dating (summarised in Kalsbeek & Taylor, 1985) the granite intrusions were emplaced over the period c. 1850 to 1740 Ma.

The 'folded migmatite zone' comprises large occurrences of pelitic to semipelitic sediments and basic volcanic sequences which are metamorphosed up to granulite facies and are often strongly migmatised. They are intruded by granites of the same general age as those in the 'granite zone'.

The 'flat-lying migmatite complex' consists of flat-lying, high-grade migmatites intruded by large sheets and bodies of granites, including major Rapakivi complexes.

The Gardar Province comprises mid-Proterozoic sediments and lavas deposited in a continental rift zone as well as 10 major alkaline intrusive complexes and numerous dykes emplaced between c. 1300 and 1120 Ma. The latest review of the province is given by Upton & Emeleus (1987). Extensive hydrothermal activity accompanying the emplacement of the magmas (Berrangé, 1966) caused mineralisation in the margins of two intrusive complexes (Nb, Ta, U, Th, Pb, Mo) (Sørensen et al., 1974; Tukiainen, 1988) as well as in fracture zones in the surrounding granite-gneiss (U) (Armour-Brown et al., 1982; Nyegaard et al., 1986; Nyegaard & Armour-Brown, 1986).

Previous gold exploration

During reconnaissance exploration in the Archaean Tartoq group of the 'border zone' gold bearing sulphide mineralisation in greenstone sequences was encountered (Appel & Secher, 1984). Three types were distinguished: 1) Siliceous strata-bound arsenopyrite-pyrite layers, 2) tennantite-bearing quartz veins associated with these layers, and 3) scattered chalcopyritebearing quartz veins and lenses. Analyses of 5 mineralised grab samples yielded from 2 to 17 ppm Au.

At the abandoned copper mine (Josva mine) on the south coast of Kobberminebugten Harry & Oen (1964) found grains of gold in bornite veins hosted by Proterozoic supracrustal rocks. Later investigations (Ghisler, 1968; Secher & Kalvig, 1987; Secher, unpublished data) confirmed the presence of small amounts of gold (2-3ppm) associated with the bornite mineralisation. A soil survey with samples collected in a 100 by 100 m grid over the supracrustals of the south coast of Kobberminebugten did not indicate any important gold mineralisation (gold values up to 16 ppb) other than the one at Josva mine associated with bornite (A. Steenfelt, unpublished data).

10

Reconnaissance by Kryolitselskabet Øresund A/S (Nielsen, 1976) and the Geological Survey of Greenland (Secher & Kalvig, 1987 and unpublished data) over the Proterozoic supracrustal sequences of the 'border zone' resulted in the discovery of gold bearing sulphide mineralisations (2-5 ppm in grab samples) at other sites in the area surrounding Kobberminebugten.

In the 'folded migmatite zone' gold has been detected in both stream sediments (fine fraction) and panned concentrates. In 1987 a batch of 238 stream sediment samples of the Geological Survey of Greenland's Syduran collection was analysed for gold by instrumental neutron activation analysis (Steenfelt, 1987). The result showed that gold anomalies are frequent in the metavolcanic as well as in the metasedimentary sequences. Gold has been

reported to occur in small sulphide mineralised ultramafic bodies which have intruded the volcano-sedimentary sequences (Berrangé, 1970). However, the frequency and size of the mafic intrusions are not adequate to explain the number of stream sediment gold anomalies, and the source of gold is more likely associated with the supracrustal rocks. There has been limited field work in the area, and so far mineralisations have not been reported.

Following the discovery of gold in panned concentrates in the Nanortalik district a joint-venture of Nanortalik Minerals A/S and Greenex A/S evaluated the potential for placer type gold in Quaternary alluvial deposits in two large valleys north of Nanortalik (Christensen, 1989). A large number of 0.6 m³ samples were collected along lines across the valleys from the uppermost 1 to 4 metres of the alluvial sediment, preconcentrated in a Denver Goldsaver, and two fractions (2-10 mm and <2 mm) were analysed. The gold contents varied typically between 0.001 g/m³ and 0.02 g/m³, and this was regarded unfavourable and the testing was discontinued.

Samples of panned concentrate were collected in the drainage systems above the gold-bearing alluvium (Christensen, 1989). Gold (100 ppb) and arsenic (up to 4600 ppb) were obtained in samples draining mafic metavolcanic rocks, and it is therefore assumed that the source of gold is associated with these rocks. However, gold mineralisation has not yet been found *in situ* and the c. 100 rock analyses made were all below 88 ppb gold.

Distribution of gold values in the present study

In the evaluation of exploration data the 98th percentile is commonly used use as a threshold above which values are defined as anomalous. In the present to mile case the use of this principle would result in a threshold of 25 ppb Au and 34 anomalies. However, taking into consideration that the samples used in this survey were not collected specifically for gold exploration and that the amounts analysed are very small, the actual numbers may not be very significant. Consequently, low values (but above the detection limit) should and not be disregarded and higher values should not be used to rank the importance of anomalous sites. The most significant aspect of these regional data are the abundance and spatial distribution of the gold-bearing samples in relation to the geological setting.

When the enclosed gold distribution map is compared with the geological map, the first impression is that the gold containing samples are scattered area. They appear not to be ploysed area.

controlled by any lithological or structural features. However, the distribution of arsenic is a great help in interpreting the distribution pattern of gold. This pattern is best described using the structural zones of Allaart (1976).

<u>In the 'border zone'</u> all the elevated arsenic values as well as almost all the gold values above the detection limit may be related to occurrences of basic supracrustal rocks. The few gold values situated in areas of gneiss do not coincide with elevated arsenic and cannot be explained with our present knowledge. The gold-arsenic data confirm the existence of sulphide associated gold mineralisation both in the Archaean (Tartoq group), and the Proterozoic supracrustals (Midternæs, Grænseland and the island of Arsuk) which is in agreement with the results of most of the previous geological reconnaissance mentioned above.

Over the 'granite zone', gold anomalies are scattered with some tendency to clustering around Igaliko Fjord. These anomalies are located in gneiss or granite and do not coincide with areas of elevated arsenic. In fact, the absence of arsenic in the 'granite zone' is remarkable. The few arsenic values above 2 ppm are located in the Gardar alkaline intrusion Ilímaussaq, east of Narssaq. There is a scarcity of supracrustal rocks within the 'granite zone' and it is unlikely that the anomalies are related to unmapped basic volcanics. At present a documented explanation of the anomalies in this zone cannot be offered. One can only guess at a possible relation to shear and fracture zones which are known to occur quite frequently in the 'granite zone' and in the Igaliko Fjord area in particular (Nyegaard & Armour-Brown, 1986; Conradsen et al., 1985). The wallrock of many of the faults and fractures is affected by extensive secondary alteration (hematitisation, chloritisation, silicification) accompanied by small amounts of fluorite, calcite, iron and copper sulphides (Berrangé, 1966). One sample of altered rock contained 0.2 ppm gold. Nyegaard & Armour-Brown (1986) provided further evidence for hydrothermal alteration in fracture zones in the form of albitisation, carbonatisation, and fluoritisation associated with uranium mineralisation.

<u>In the 'folded migmatite zone'</u> there is a concentration of high arsenic values. The boundary between the 'granite zone' and the 'folded migmatite zone' nearly coincides with the northern boundary of this arsenic province. The high arsenic values appear to be related to the Proterozoic metasediments and basic metavolcanics. Compared with the 'border zone' the supracrustals of the 'folded migmatite zone' are considerably higher in arsenic.

The 'folded migmatite zone' also contains the greatest density of elevated gold values. It should be noted that 238 of the Syduran samples from this

zone have been analysed for gold once before, as mentioned above, and little material (80 samples) was left to be included in the present analysis. Therefore, there are more anomalies in the Nanortalik region (see Steenfelt, 1987) than shown in the present gold distribution map.

The result of the present analyses strongly supports previous indications (anomalies in stream sediments and gold plus arsenic in panned concentrates) of the existence of gold mineralisation associated with the Proterozoic supracrustal sequences in the 'folded migmatite zone'. As mentioned before, this area has not been satisfactorily explored for mineral occurrences, and primary gold mineralisation has not been so far reported.

In the 'flat-lying migmatite complex' the level of arsenic is low. The few high gold values are scattered and cannot be related to any particular geological setting.

Conclusion

The new analytical data provide the first regional information on the abundance and distribution of gold in South Greenland, and they show that anomalies occur over large areas.

While previous gold exploration has mainly been concerned with greenstones in the 'border zone', the present data strongly suggest that the 'folded migmatite zone' and the 'granite zone' should also be considered as target areas for intensified gold exploration.

In the 'folded migmatite zone' the gold anomalies occur in arsenic rich supracrustals which suggests that the gold is associated with sulphide mineralised rocks. By contrast, the gold anomalies in the 'granite zone' occur in granitic rocks devoid of arsenic. This indicates a completely different type of mineralisation which is tentatively related to hydrothermal activity associated with Gardar magmatism.

The frequency and magnitude of gold concentrations encountered in South Greenland is surprisingly high considering the fact that the samples and fractions used for analysis were not ideal from a gold exploration point of view. Moreover when the gold results are compared with that of other regional surveys the values obtained in South Greenland seem high. For example, in the northern part of Fennoscandia the 95th percentile for gold content in the fine fraction of till was 9.5 ppb and maximum 436 ppb (Bølviken et al.,1986).

Acknowledgments

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Enclosures

Geological map of Greenland 1:500 000 sheet 1, South Greenland (Allaart, 1975).

Semi-transparent map 1:500 000 of gold in the <0.1 mm grain size fraction of stream sediments. The radius of the circles is linear proportional to the gold concentration of the sample (see scaling factor on the map). Maximum radius (0.5 cm) equals 100 ppb. Gold concentrations above 100 ppb are written in the circles.

Semi-transparent map 1:500 000 of arsenic in the same samples and presented in the same way as gold. Maximum radius equals 100 ppm. Higher concentrations are not quantified on the map.

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SHEET 2 FREDERIKSHÅBS ISBLINK - SØNDRE STRØMFJORD

SHEET 1 Kortal h SYDGRØNLAND

Topography based on maps of the Geodetic Institute, Denmark, and reproduced with permission A. 649/72.





General compilation by J. H. Allaart.

Geological mapping by GGU teams on scales of 1:40 000 or smaller 1966-1970, and by R. H. Wallis with the St. Andrews University South Greenland Expedition 1960 and the Birmingham University Expedition to South Greenland 1961.

GEOLOGICAL MAP OF GREENLAND SHEET 1 SYDGRØNLAND

1:500 000

 Implicit
 Implicit

 5
 0
 10
 20
 30
 40 statute miles
 Højdeforskellen mellem kurverne 200 m Contour interval 200 m

------ BOUNDARY, established BOUNDARY, inferred STRIKE AND DIP of lithological layering of any origin 🖌 💉 STRIKE AND DIP of foliation ----- FAULT AND MYLONITE THRUST ----- UNCONFORMITY CONGLOMERATE AUGEN TEXTURE ◇ PORPHYRITIC TEXTURE 🗞 🛛 MINE, abandoned





~ MIGMATISATION



SOUTH GREENLAND STREAM SEDIMENTS GRAIN SIZE FRACTION < 100 MICRON ELEMENT UNIT SC.FACT MAX MIN AU PPB 0.005 0.5 0.02 .

SCALE 1:500000

PROJECTION: LAMBERT'S CONICAL ORTHOMORPHIC. ELLIPSOID: INTERNATIONAL. A = 6378388M. F = 1/297. STANDARD PARALLEL: 606 45' 00 N



SCALE 1:500000

PROJECTION: LAMBERT'S CONICAL ORTHOMORPHIC. ELLIPSOID: INTERNATIONAL. A = 6378388M. F = 1/297. STANDARD PARALLEL: 60G 45' 00 ' N

