

**GANE-1**  
**GANK-1**  
**GANT-1**

**Palynology of the boreholes GANE-1, GANK-1 and GANT-1 Nuussuaq, West Greenland**  
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## Abstract

The dinoflagellate cyst assemblages from 87 samples from the boreholes GANE#1, GANK#1 and GANT#1 on Nuussuaq, West Greenland are described. GANE#1 and GANK#1 are divided into two informal dinoflagellate cyst intervals that can be correlated and dated as Selandian to ?Early Thanetian, Late Paleocene.

GANT#1 is dated as Campanian (Late Cretaceous) to Danian (Early Paleocene) and is divided into four informal dinoflagellate cyst intervals. The Cretaceous/Tertiary boundary has not been precisely located but is supposed to occur just below the base of interval I in GANT#1. Reworked Late Cretaceous species often occur in the Paleocene species assemblages.

The palynological record from GANT#1 has made correlation possible of the Upper Cretaceous deposits from northern Nuussuaq with similar deposits on central Nuussuaq and Svartenhuk Halvø.

## Introduction

In the summer of 1995 the first commercial oil exploration drilling for almost 20 years was carried out in Greenland with the Canadian grønArctic energy inc. as operator. In May 1995 grønArctic energy inc. and Platinova A/S were granted an exclusive exploration licence for a 1692 km<sup>2</sup> large area covering western Nuussuaq (Fig. 1). grønArctic began their exploration programme with drilling of three slim-core holes (GANE#1, GANK#1 and GANT#1) to depths between 400 and 900 m.

The Geological Survey of Denmark and Greenland (GEUS) carried out a geological description at the drill sites followed by a major sampling and analytical programme. The results from these studies include detailed sedimentological, organic geochemical and biostratigraphic reports. These reports have been released from 1 April, 1997 whereas all other results obtained by grønArctic energy inc. during their exploration programme are still confidential.

GANE#1 and GANK#1 are situated in the Aaffarsuaq valley, approximately 10 and 15 km, respectively, ESE of GGU's Marraat-1 well (Fig. 1). The main target both places was a Marraat type play, i.e. a Maastrichtian–Lower Paleocene channel or turbidite sandstone reservoir with a deltaic source rock of similar age, within a structural or stratigraphic trap. Both wells are situated some kilometres SE of the area where oil has previously been documented at the surface (Christiansen *et al.*, 1995a, 1996a). Both wells were sited on hyaloclastic breccias, the base of which was penetrated at 497 m in GANE#1 and at 115 m in GANK#1. Oil was identified bleeding from the cores at several levels in the volcanic rocks. In the underlying Maastrichtian?–Tertiary sedimentary succession gas under pressure as well as oil impregnation was encountered (see details in Christiansen *et al.*, 1996c).

GANT#1 is situated in the Tunorsuaq valley (Fig. 1) where a Cretaceous–Tertiary succession similar to that known from the north coast of Nuussuaq is exposed. The main target was a structural play, with marine Cretaceous sandstones as the reservoir unit and an inferred mid-Cretaceous oil-prone source rock. The target was clearly inspired by the discovery of wet gas in a mineral exploration drill hole at Serfat (Fig. 1) on the north coast of Nuussuaq in the summer of 1994 (Dam & Nøhr-Hansen, 1995). GANT#1 penetrated a succession of Upper Cretaceous to Lower Paleocene marine sediments. The core documented a number of sandstone layers with gas under pressure (Christiansen *et al.*, 1996c).

# Palynology

## Previous palynological studies on southern and central Nuussuaq

Croxton (1978a, b) briefly described the palynomorph assemblages from seven localities on central Nuussuaq (C4-C7, C21, M19, M22; Fig. 2). The palynomorphs from Qilakitsoq (C4), Qaatunnat Ilorliit (C5), Ilugisssiq (C6) and Nallurarissat (C7) indicate a Late Cenomanian to Early Campanian age. A possible reworked Maastrichtian assemblage is recorded from the top of section C5, and dinoflagellate cysts from the topmost shale at C6 indicate a possible 'middle' Paleocene age (Croxtan. 1978a). Sections C21 and M19 represent the 'Oyster-ammonite conglomerate' from Agatdalen; according to Croxtan (1978b) palynomorphs from these sections may indicate reworked Maastrichtian floras. A few dinoflagellate cysts, probably indicating a Late Campanian age, were recorded by Croxtan (1978a) from Scaphitesnæsen (M22).

Eight sections from central Nuussuaq (M16-M23; Fig. 2) were sampled by Hansen (1976); data on the palynological content from the two sections M16 and M17 from Tunoqqu have not been published.

Hansen (1980) described the Paleocene dinoflagellate cysts from the Sonja section (M18), Turritellakløft section (M20), Qaarsutjægerdal section (M21) and Ättestupet section (M23). According to Hansen (1980) the Middle Campanian to Early Maastrichtian species *Isabelidinium microarmum* is recorded as reworked specimens in the sections M18-M21.

Piasecki *et al.* (1992) described mid-Paleocene dinoflagellate cyst assemblages from sediments interbedded with the Tertiary volcanic rocks on Disko and Nuussuaq and dated the sediments as NP 4 to NP 8.

Ehman *et al.* (1976) studied the four sections Qilakitsoq (N10), Turritellakløft (N15), Qaarsutjægerdal (N16) and Nassaar (N17) central Nuussuaq (Fig. 2). The ages given by Ehman *et al.* (1976) are Middle Cenomanian and early Danian for N10, Campanian or Maastrichtian to Paleocene for N15, Paleocene for N16 and N17. It should be noted that the ages given in the text by Ehman *et al.* (1976) are not always consistent with the ages given in their logs (see Pulvertaft, 1987, table 1).

Nøhr-Hansen (1994b; 1996) described the palynomorph assemblages from 15 sections and four subsurface sections in the Agatdalen region and dated the marine succession as Late Santonian to Middle Campanian. One sample is dated as Paleocene.

Nøhr-Hansen in Christiansen *et al.* (1995b) described the dinoflagellate cyst assemblages from four samples from the GANW#1 hole on the south coast of Nuussuaq (Fig. 2) and dated the samples as early Thanetian, now Selandian, Early Paleocene.

The upper Cretaceous spore and pollen assemblages from the Ataata Kuua section on south-western Nuussuaq have been studied by D.J. McIntyre, but data have not been published.

## **Previous palynological studies on northern Nuussuaq**

Croxton (1978a, b, 1980) briefly described palynomorph assemblages from three localities on the north coast of Nuussuaq (M25, M27, M28; Fig. 2). The palynomorphs from these three localities in the Kangilia and Annertuneq area indicate a Late Campanian to Maastrichtian age.

Hansen (1980) described the Paleocene dinoflagellate cysts and proposed a zonation for the mudstone deposited above the so-called 'basal Danian conglomerate' in the Kangilia/Annertuneq section (M25; Fig. 2).

Nøhr-Hansen (1993) described a low diversity dinoflagellate assemblage of Late Maastrichtian? to Early Paleocene age from the uppermost part of the thick turbidite succession exposed on the south-east side of the Itilli valley in Nuussuaq.

Nøhr-Hansen (1994c; 1996) described the palynomorph assemblages from five sections and three subsurface sections on the north coast of Nuussuaq. The sections make up an at least 700 m thick dark mudstone succession. The dinoflagellate cysts and pollen date the majority of the samples as Late Campanian and Maastrichtian. A few samples were dated as Coniacian-Late Santonian and Early Paleocene. The section at Annertuneq includes the Cretaceous/Tertiary boundary and the palynomorph assemblage has been described by Nøhr-Hansen & Dam (1997).

## Samples and methods

At the well site, samples were collected and canned at approximately 3 metre intervals (Bate, 1995; Dahl *et al.*, 1995a,b). Cores were wrapped in aluminium foil and packed in core boxes in order to avoid (or at least reduce) contamination. A set of 87 samples were processed for palynological studies.

### Preparation

Palynological preparation and studies were carried out at (GEUS). Palynomorphs were extracted from 20 g of sample by modified standard preparation techniques. The bulk of the minerals were dissolved by hydrochloric and hydrofluoric acids. A first slide was made after this treatment. A second slide was made of the organic residue after sieving using a 20 micron nylon mesh. A third slide was made after oxidation (3 to 10 minutes) with fuming nitric acid, followed by washing with a weak potassium hydroxide solution. The oxidation was carried out in order to clean the samples of minor amorphous kerogen particles and pyrite. Finally, palynomorphs were separated from coal particles and woody material in most samples using the separation method described by Hansen & Gudmundsson (1978) or by swirling.

After each of the steps mentioned above the organic residues were mounted in a permanent medium (Eukitt R; produced by O. Kindler, Germany).

### Recording of material and analyses

The palynological slides were studied with transmitted light using a Leitz Dialux 22 microscope (512 742/057691). All the co-ordinates in the plate captions refer to this microscope. England finder index corners: Z 75 4 = 74.6-92.3; Z 1 3 = 1.9-92.2; A 1 1 = 1.9-116.7; A 65 2 = 64.6-116.6, centre: O 38 = 38.1-103.3.

The illustrated dinoflagellate cysts are marked with a sample number, slide number, microscope co-ordinates, laser-video-record number (LVR) and data base number (MicroImage; MI) for later identification. The illustrated dinoflagellate cysts are housed at the Geological Survey of Denmark and Greenland (Copenhagen) where they are accessible for examination.

Dinoflagellate cyst and acritarch species were recorded from the sieved, oxidised or gravitation-separated slides. Counting of specimens was done on the 73 samples that revealed dinoflagellate cysts (Figs 3, 4, 5), approximately 100 specimens were counted when possible.

The dinoflagellate cyst stratigraphy proposed here for the GANE#1, GANK#1 and GANT#1 holes is based on the first and the last occurrences and the acme of stratigraphically important species.

# Palynostratigraphy and thermal maturity of the borehole GANK#1

GANK#1 is situated close to the river Kuussuaq, about 10 km east of the outlet at western end of the Aaffarsuaq valley (Fig. 1). The elevation of GANK#1 is 91 m a.s.l., and the hole was drilled to a total depth of 398.98 m. A side-track hole (GANK#1A) was kicked-off at 218.55 m and terminated at a depth of 332.84 m (Dahl *et al.*, 1995a; Dam, 1996b).

The organic geochemical analyses yielded moderately low Rock Eval  $T_{max}$  values between 421 and 440°C throughout the sedimentary succession down to -400 m suggesting a maturity level corresponding to before or in the early part of the oil window (Christiansen *et al.*, 1996b).

The dinoflagellate cyst stratigraphy proposed here for the marine Paleocene section in GANK#1 is based on a study of material from 26 core samples of which one was barren.

## Palynological intervals

The species recorded from GANK#1 are listed on the range chart (Enclosure 1).

Only a few morphologically characteristic and stratigraphically important species have been recognised, and based on their first and last occurrences it has been possible to distinguish two dinoflagellate cysts intervals within the Paleocene strata. These intervals are informally described, and a correlation is suggested to the dinoflagellate cyst interval biozones described by Powell (1992) from the North Sea region (Fig. 6).

### Interval I, -120.9 to -180.8 m

Age: Selandian (Late Paleocene).

This interval is dominated by terrestrially derived black to brownish woody material, cuticles and few bisacate pollen. Dinoflagellate cysts are common. The diversity is low to moderately high (3 to 24 species). The specimens are generally well preserved. Reworking of Late Cretaceous and Early Paleocene species is common (see Enclosure 1).

The interval of GANK#1 (-130.7 to -180.8 m) may be correlated with the Selandian *Palaeoperidinium pyrophorum* (Ppy) dinoflagellate cyst Interval Biozone (Fig. 6) described from the North Sea region by Powell (1992).

*Characteristic species.* The interval is characterised by the presence of *Alisocysta margarita* (rare), *Areoligera gippingensis*, *Cerodinium diebelii*, *C. speciosum*, *C. striatum*. *Isabelidinium*

*aff. viborgense*, *Palaeocystodinium bulliforme* and *Palaeoperidinium pyrophorum* (common) (Enclosure 1).

**Discussion.** The first appearance datum (FAD) and last appearance datum (LAD) of *A. margarita* in the interval suggest a Selandian age correlating with the *P. pyrophorum* (Ppy) or the younger early Thanetian *Alisocysta margarita* (Ama) dinoflagellate cyst Interval Biozones (Fig. 6) of Powell (1992). The common occurrence of *P. pyrophorum* and the presence of *C. speciosum*, *C. striatum*, *P. bulliforme* suggest correlation with the Ppy Interval Biozone since the species all have their LAD at the top of the Ppy Interval Biozone (Powell, 1992).

The FAD of *Areoligera gippingensis* at the base of the interval I suggests correlation with the Ama Interval Biozones according to Powell *et al.* (1996). However, Mudge & Bujak (1996) have recorded *A. gippingensis* (sporadically occurring) together with common *P. pyrophorum* in the Ppy Interval Biozone of Powell (1992). The *A. gippingensis* acme described by Powell *et al.* (1996) from the lower part of the Ama Interval Biozones, has not been recorded in interval I.

The species from Nuussuaq identified as *Isabelidinium aff. viborgense* is morphological very similar to the stratigraphically important species *Isabelidinium viborgense* described from Denmark (Heilmann-Clausen, 1985). It differs, however, from the holotype of *I. viborgense* by the presence of a distinct pattern of longitudinal folds on the periphragm. *Isabelidinium viborgense* has its LAD in the lower part of the Ppy Interval Biozone and its FAD in the lower part of the older *Cerodinium speciosum* (Csp) Interval Biozone (Powell, 1992). According to Mudge & Bujak (1996) *I. viborgense* has its LAD just before the FAD of sporadic occurrences of *A. gippingensis* in the North Sea. The co-occurrence of few *I. aff. viborgense* and few *A. gippingensis* in GANK#1 may suggest that *I. aff. viborgense* have a stratigraphically higher LAD or that *A. gippingensis* have a stratigraphically lower FAD in Arctic areas.

## **Interval II, -188.9 to -395.8 m**

Age: Selandian (Late Paleocene)

This interval is dominated by terrestrially derived black to brownish woody material, cuticles and few bisacate pollen. Dinoflagellate cysts are common. The diversity is low to moderate (2 to 15 species). The specimens are generally well preserved. Reworking of Late Cretaceous and Early Paleocene species is rare (see Enclosure 1).

The lower part of GANK#1 (-188.9 to -395.8 m) may be correlated with the Selandian *Cerodinium speciosum* (Csp) dinoflagellate cyst Interval Biozone (Fig. 6) described from the North Sea region by Powell (1992).

**Characteristic species.** The interval is characterised by the presence of *Cerodinium diebelii*, *C. speciosum*, *C. striatum*, *Isabelidinium aff. viborgense*, *Palaeocystodinium bulliforme* and *Palaeoperidinium pyrophorum*. *Fibrocysta cf. capitata* is common at -379.5 m (Enclosure 1).

*Discussion.* The presence of *Cerodinium speciosum* throughout this interval suggests Selandian age correlating with the Csp dinoflagellate cyst Interval Biozones in the North Sea region (Powell, 1992).

*Isabelidinium viborgense* has its FAD in the lower part of the Csp Interval Biozone (Powell, 1992). The presence of *I. aff. viborgense* throughout the interval therefore suggests correlation with the middle and upper part of the Csp dinoflagellate cyst Interval Biozone of Powell (1992).

# Palynostratigraphy and thermal maturity of the borehole GANE#1

GANE#1 is situated at Eqalulik, about 6 km east of the outlet of the large river Kuussuaq at the western end of the Aaffarsuaq valley (Fig. 1). The elevation of GANE#1 is 114 m a.s.l., and the hole was drilled to a total depth of 641.29 m. A side-track hole (GANE#1A) was kicked-off at 533.4 and terminated at a depth of 707 m (Bate, 1995; Dam, 1996a).

The organic geochemical analyses yielded moderately low Rock Eval  $T_{max}$  values between 432 and 443 °C throughout the sedimentary succession from -503.9 to -663.5 m suggesting a maturity level corresponding to the early part of the oil window (Christiansen *et al.*, 1996b).

The dinoflagellate cyst stratigraphy proposed for the marine Paleocene in GANE#1 is based on the study of material from 25 core samples of which 9 were barren.

## Palynological intervals

Relatively low species diversity and density are recorded in GANE#1. Only a few morphologically characteristic and stratigraphically important species have been recognised, and based on their first and last occurrences it has been possible to distinguish two dinoflagellate cysts intervals within the Paleocene strata. These intervals are informally described, and a correlation is suggested to the dinoflagellate cyst interval biozones described by Powell (1992) from the North Sea region (Fig. 6).

### Interval I, -503.9 to -564.7 m

Age: Selandian (Late Paleocene)

This interval is dominated by terrestrially derived black to brownish woody material, cuticles and a few bisacate pollen. Dinoflagellate cysts are rare to common. The diversity is low to moderate (2 to 17 species). The specimens are generally well preserved. Reworking of Late Cretaceous species occurs throughout (see Enclosure 2).

The interval I of GANE#1 (-503.9 to -564.7 m) may be correlated with the Selandian *Palaeoperidinium pyrophorum* (Ppy) dinoflagellate cyst Interval Biozone described from the North Sea region by Powell (1992).

*Characteristic species.* The interval is characterised by the presence of *Areoligera gippingensis*, *Cerodinium speciosum*, *Palaeocystodinium bulliforme* and *Palaeoperidinium pyrophorum*. At the base of the interval *Alisocysta margarita* is common. A single specimen of *Isabelidinium aff. viborgense* has been recorded from the base (see Enclosure 2).

*Discussion.* The correlation of interval I with the Selandian Ppy Interval Biozone (Powell, 1992), is, like in GANK#1, based on the FAD and LAD of *A. margarita* at the base of the interval and on the FAD of *A. gippingensis*. However, the dinoflagellate cyst assemblage differs slightly from the assemblages recorded from interval I in GANK#1.

*Alisocysta margarita* and *I. aff. viborgense* only occur at the base of the interval and are characteristically not present at higher levels. *Cerodinium diebelii* is not recorded from the present interval. The species *C. speciosum*, *C. striatum* and *Palaeocystodinium bulliforme* are rare and only occur at the top of the interval.

## **Interval II, -565.5 to - 663.5 m**

Age: Selandian (Late Paleocene)

This interval is dominated by terrestrially derived black to brownish woody material, cuticles and few bisacate pollen. Dinoflagellate cysts are rare to common. The diversity is low (2 to 12 species). The specimens are generally well preserved. Reworking of Late Cretaceous species is relatively common in interval II (see Enclosure 2).

The lower part of GANE#1 (-565.5 to - 663.5 m) may be correlated with the Selandian *Cerodinium speciosum* (Csp) dinoflagellate cyst Interval Biozone described by Powell (1992) from the North Sea region (Fig. 6).

*Characteristic species.* The interval is characterised by the presence of few *Cerodinium diebelii*, *C. speciosum*, *C. striatum*, *Palaeocystodinium bulliforme* (-635.4 to -663.5 m), *Palaeoperidinium pyrophorum*, Dinocyst sp. 01 HNH 1997 (-635.4 to -649.4 m) and *Isabelidinium aff. viborgense* (common at -565.5 m but not recorded below; Enclosure 2).

*Discussion.* The presence of a few specimens of *Cerodinium speciosum*, *Palaeocystodinium bulliforme* in the interval and common *Isabelidinium aff. viborgense* at the top of interval II suggests a Selandian age correlating with the Csp or the older *Spinidinium densispinatum* Sde dinoflagellate cyst Interval Biozones described by Powell (1992) from the North Sea region (Fig. 6).

# Palynostratigraphy and thermal maturity of the borehole GANT#1

GANT#1 is situated in the Tunorsuaq valley about 11 km ESE from where the valley converges with the Itilli valley (Fig. 1). The elevation of GANT#1 is 440 m a.s.l., and the hole was drilled to a total depth of 900.68 m (Dahl *et al.*, 1995b; Dam, 1996c). The core recovery was close to 100% in all three wells.

The organic geochemical analyses yielded moderately low Rock Eval  $T_{max}$  values between 431 and 444 °C down to -794 m suggesting a maturity level corresponding to the early part of the oil. Below -794 m  $T_{max}$  values increase slightly to 448 - 467 °C suggesting peak generation (Christiansen *et al.*, 1996b).

The dinoflagellate cyst stratigraphy proposed here for the marine Lower Paleocene and Upper Cretaceous in GANT#1 is based on the study of material from 36 samples of which 4 were barren.

## Palynological intervals

The stratigraphic range of species recorded from GANT#1 are shown on the range charts (Enclosures 3 & 4). Four dinoflagellate intervals are recognised in the Early Campanian to Early Paleocene strata. These intervals are informally described, and when possible correlated to intervals described by Nøhr-Hansen (1996) from the north coast and central Nuussuaq.

### Interval I, -58.2 to -133.2 m

Age: Danian (Early Paleocene)

This interval is dominated by terrestrially derived black to brownish woody material, cuticles and few bisacate pollen. Dinoflagellate cysts are common and the diversity is moderate (up to 20 species). The specimens are generally well preserved. The Late Cretaceous dinoflagellate cyst species; *Stephodinium coronatum*, *Chatangiella* aff. *tripartita*, the genus *Isabelidinium* and the pollen genus *Aquilapollenites* are considered reworked.

The uppermost part of GANT#1 (-58.2 to -133.2 m) is correlated with the early Danian flora previously described from Annertuneq on the north coast on Nuussuaq (Nøhr-Hansen and Dam, 1997).

*Characteristic species.* The base of the interval is characterised by the FAD of *Senoniasphaera inomata* and by the FAD and abundance of the species *Trithyrodinium fragile*. The interval is also characterised by the occurrence of *Spongodinium delitiense* in the

lowermost part and by the presence of *Alisocysta circumtabulata* in the middle part (see Enclosure 4).

*Discussion.* The FAD of *Senoniasphaera inomata* and the FAD and abundance of the species *Trithyrodinium fragile* has recently been recorded from sediments just above the uppermost Maastrichtian level at Annertuneq on the north Nuussuaq (Nøhr-Hansen and Dam, 1997). The occurrence of *S. inomata* in the lower part of the interval suggests an age not younger than middle Danian (Costa & Davey, 1992) and not older than earliest Danian (Moshkovitz & Habib, 1993). The occurrence of *Spongodinium delitiense* in the lowermost part of the interval suggests an age not younger than early Danian (Williams *et al.*, 1993). This may be supported by the absence of *Cerodinium striatum* which according to Costa & Davey (1992), has a FAD in the late Danian. The species *C. striatum* has been recorded from younger sediment in the area (see above in the description of intervals from GANE#1 and GANK#1). According to Costa & Davey (1992) the presence of *Alisocysta circumtabulata* suggests an early Thanetian age (now Selandian). However, the species has also been reported from the basal Danian nannoplankton Zone N1 (Moshkovitz & Habib, 1993) and from the Late Maastrichtian (Schjølter *et al.*, 1997).

Palynological samples have not been studied from the coarse-grained sediments between -133.2 and -254.7 m (Fig. 5). This coarse-grained interval has been interpreted as representing canyon/major distributary turbidite channel deposits (Dam, 1996c) and is supposed to represent deposits of Maastrichtian and possibly earliest Danian age. The Cretaceous/Tertiary boundary is supposed to occur just below the base of Interval I.

## **Interval II, -254.7 to -401.1 m**

Age: Late Campanian.

The interval is dominated by terrestrially derived black to brownish woody material, cuticles and few bisacate pollen and dinoflagellate cysts are common. The diversity is low (6-9 species), and the specimens are generally poorly preserved.

Interval II is correlated with the Late Campanian *Isabelidinium cooksoniae* interval previously described from the north coast of Nuussuaq (Nøhr-Hansen, 1996). The interval was defined by the abundance of *I. cooksoniae* and *Palaeoperidinium pyrophorum* and its upper limit being immediately below the first occurrence of *Cerodinium diebelii*.

*Characteristic species.* The lower part of the interval is characterised by the abundance of *I. cooksoniae* and *P. pyrophorum* and the pollen genus *Aquilapollenites* is present throughout the interval.

### Interval III, -466.2 to -797.1 m

Age: Early to middle Campanian

The interval is dominated by terrestrially derived black to brownish woody material, cuticles and few bisacate pollen and dinoflagellate cysts are present. The diversity is very low (1-6 species) in the major part of the interval, and the specimens are generally poorly preserved. A questionable *Nyktericysta* species at -519.1 m may be reworked from lower Cretaceous strata.

The interval is correlated with the Late Campanian *Aquilapollenites* interval previously described from central Nuussuaq (Nøhr-Hansen, 1996). The interval was defined from the first occurrence of species of *Aquilapollenites* to the last occurrence of *Isabelidinium microarmum*.

*Characteristic species.* The lower part of the interval is almost barren with respect to dinoflagellate cysts. However, the upper part contains common dinoflagellate cysts and is characterised by the FAD of *Isabelidinium cooksoniae*, *I.* sp. 1 HNH 1997, *I.* cf. *microarmum* and *I.* *acuminatum*. The pollen genus *Aquilapollenites* is present throughout the interval.

*Discussion.* *Isabelidinium acuminatum* and *I.* cf. *microarmum* have previously only been recorded from central Nuussuaq, where they have their FAD within the *Aquilapollenites* interval (Nøhr-Hansen, 1996). This provides the first correlation between strata in northern and central Nuussuaq.

### Interval IV, -836.7 to -894.5 m

Age: Early Campanian

The interval is dominated by terrestrially derived black to brownish woody material, cuticles and few bisacate pollen and dinoflagellate cysts are common. The diversity is low (2-9 species) and the specimens are generally poorly preserved.

The interval is correlated with the Late Santonian or Early Campanian *Dinogymnium* sp. cf. *D. sibiricum* interval previously described only from Svartenhuk Halvø north of Nuussuaq (Nøhr-Hansen, 1996). The interval was defined from immediately above the last occurrence of *Heterosphaeridium difficile* up to the last occurrence of *Dinogymnium* sp. cf. *D. sibiricum*.

*Characteristic species.* The interval is characterised by the FAD, LAD and the abundance of *Isabelidinium* sp. 2 HNH 1997 and by the FAD and LAD of *Dinogymnium* sp. cf. *D. sibiricum*. The stratigraphically important species *Spinidinium echinoideum* and *Trithyrodinium*

*suspectum* also seem to have their LAD within the interval. The pollen genus *Aquilapollenites* is present in the upper part of the interval.

*Discussion.* *Dinogymnium* sp. cf. *D. sibiricum* has previously only been recorded from the *D.* sp. cf. *D. sibiricum* interval on Svartenhuk Halvø (Nøhr-Hansen, 1996). The species *Isabelidinium* sp. 2 HNH 1997 has some morphological similarities with *Isabelidinium* sp. aff. *I. acuminatum* previously only recorded from the *D.* sp. cf. *D. sibiricum* interval (Nøhr-Hansen, 1996). The species *S. echinoideum* and *T. suspectum* have also their LAD in the *D.* sp. cf. *D. sibiricum* interval. The presence of *Aquilapollenites* and the absence of *Heterosphaeridium difficile* suggests an age not older than Early Campanian (Nøhr-Hansen, 1994a, 1996, 1997). The pollen genus *Aquilapollenites* has not been recorded from the *D.* sp. cf. *D. sibiricum* interval on Svartenhuk Halvø (Nøhr-Hansen, 1996) which may suggest that the interval IV is slightly younger, probably Early Campanian.

## Conclusion

The drilled sedimentary sections in GANE#1 and GANK#1 have been dated as Selandian to ?early Thanetian. The correlation of the two boreholes is mainly based on the FAD and LAD of the dinoflagellate cyst species *Alisocysta margarita* and the FAD of *Areoligera gippingensis* which define the base of Interval I in both boreholes. Interval I has its base at -180.8 m in GANK#1 and at -564.7 m in GANE#1. It has not been possible to establish a biostratigraphic correlation between GANK#1 and GANE#1 (Selandian to ?early Thanetian) in the south and GANT#1 (Early Campanian to early Danian) in the north.

## Acknowledgements

The exploration drilling was carried out by the Canadian company grønArctic energy inc. The funding for the analytical programme was provided by the Government of Greenland, Minerals Office and the Danish State through the Mineral Administration for Greenland. The sample processing was carried out by Yvonne Desezar and Kim Villadsen, the figures were produced by Jette Halskov and reviewing was done by Martin Sønderholm, all staff at GEUS.

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## Figure captions

- Fig. 1. Map of Nuussuaq showing the drilling locations.
- Fig. 2. Map showing location of previously measured and palynologically analysed sections on Nuussuaq and Svartenhuk Halvø.
- Fig. 3. Simplified geological log of GANK#1 with position of samples that revealed dinoflagellate cysts. From Dam (1996b).
- Fig. 4. Simplified geological log of GANE#1 with position of samples that revealed dinoflagellate cysts. From Dam (1996a).
- Fig. 5. Simplified geological log of GANT#1 with position of samples that revealed dinoflagellate cysts. From Dam (1996c).
- Fig. 6. Dinoflagellate cyst interval biozones by Powell (1992) from the North Sea region. The position of the Selandian/Thanetian boundary is according to Powell *et al.* (1996) situated at the base of the *Alisocysta margarita* (Ama) dinoflagellate cyst Interval Biozones (Powell, 1992) from the North Sea region. The Selandian/Thanetian boundary is now correlated with the middle part of NP 6 (Berggren *et al.*, 1995).

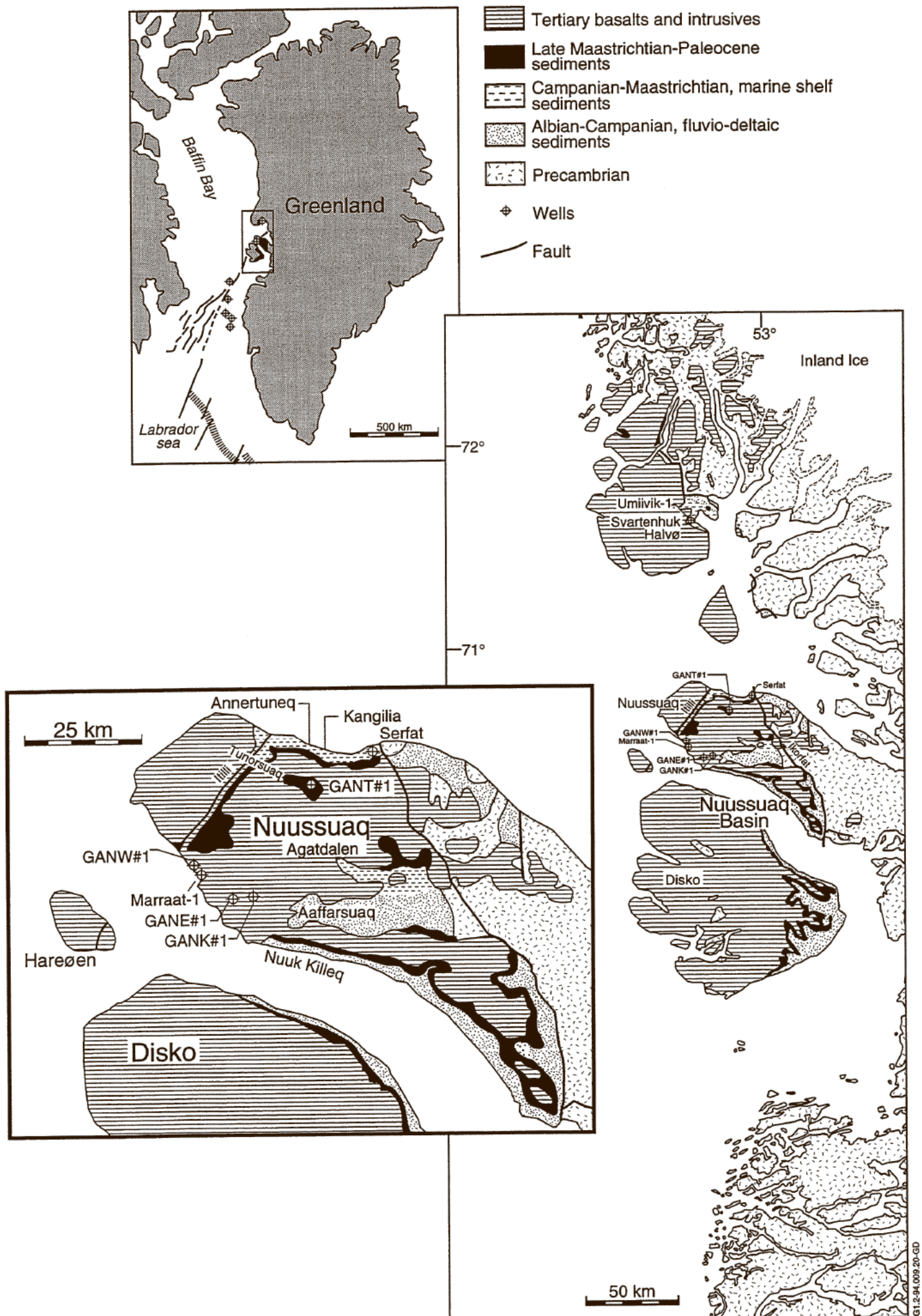


Fig. 1

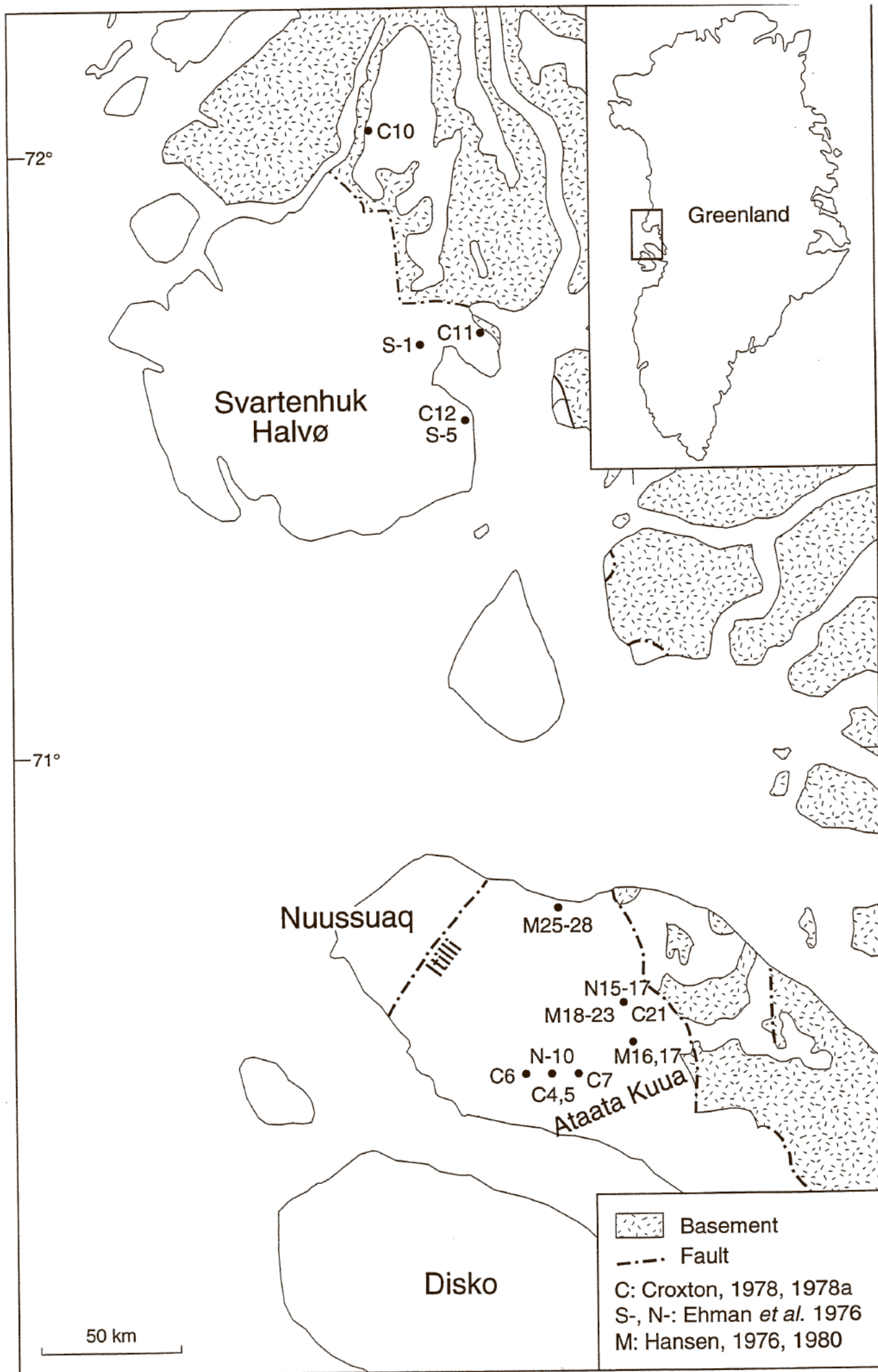


Fig.2.

# GANK#1



# GANK#1A

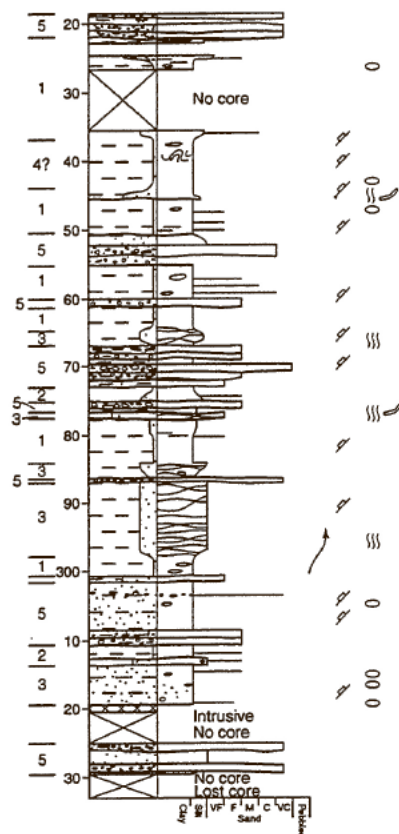


Fig. 3

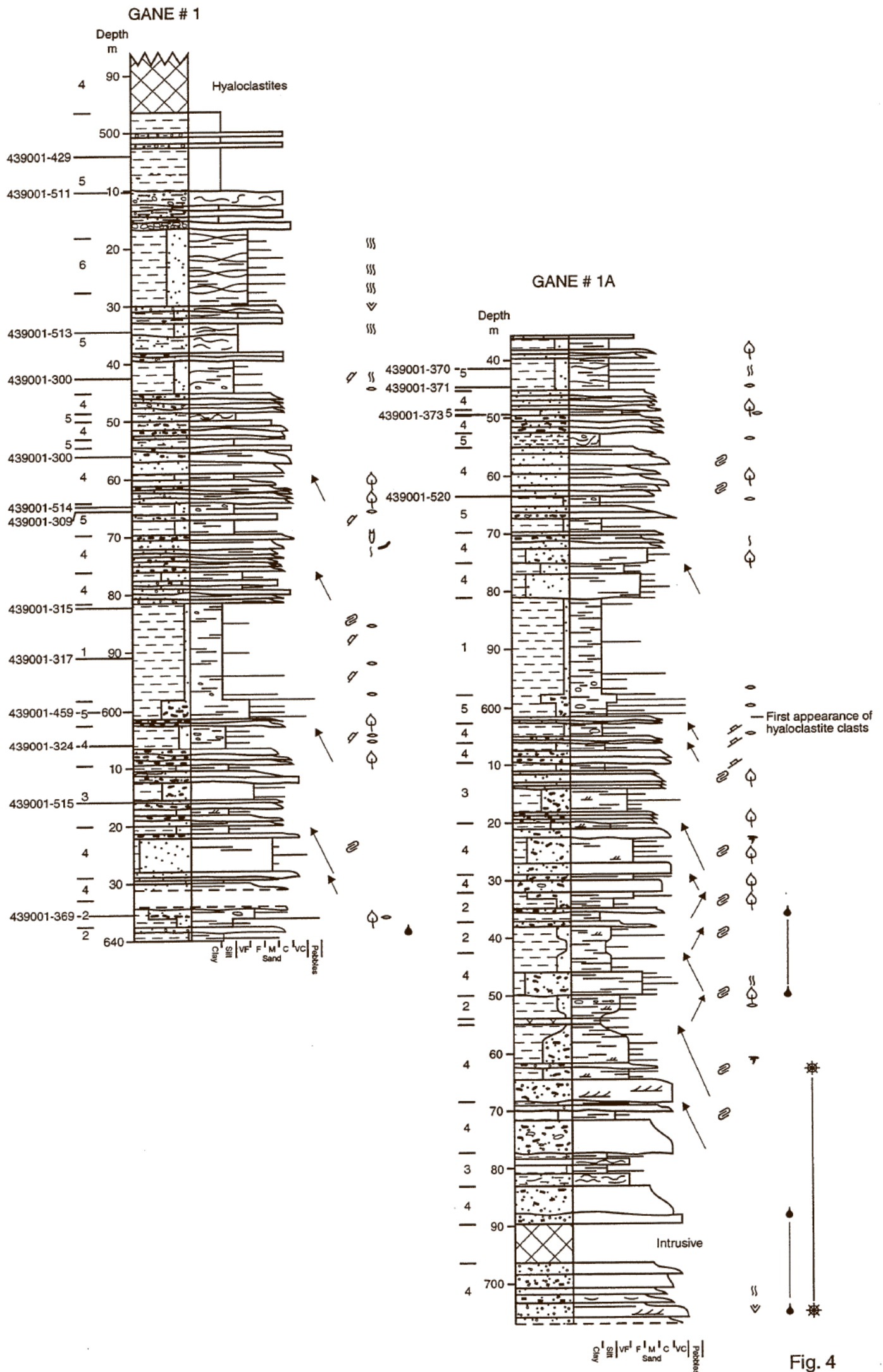


Fig. 4

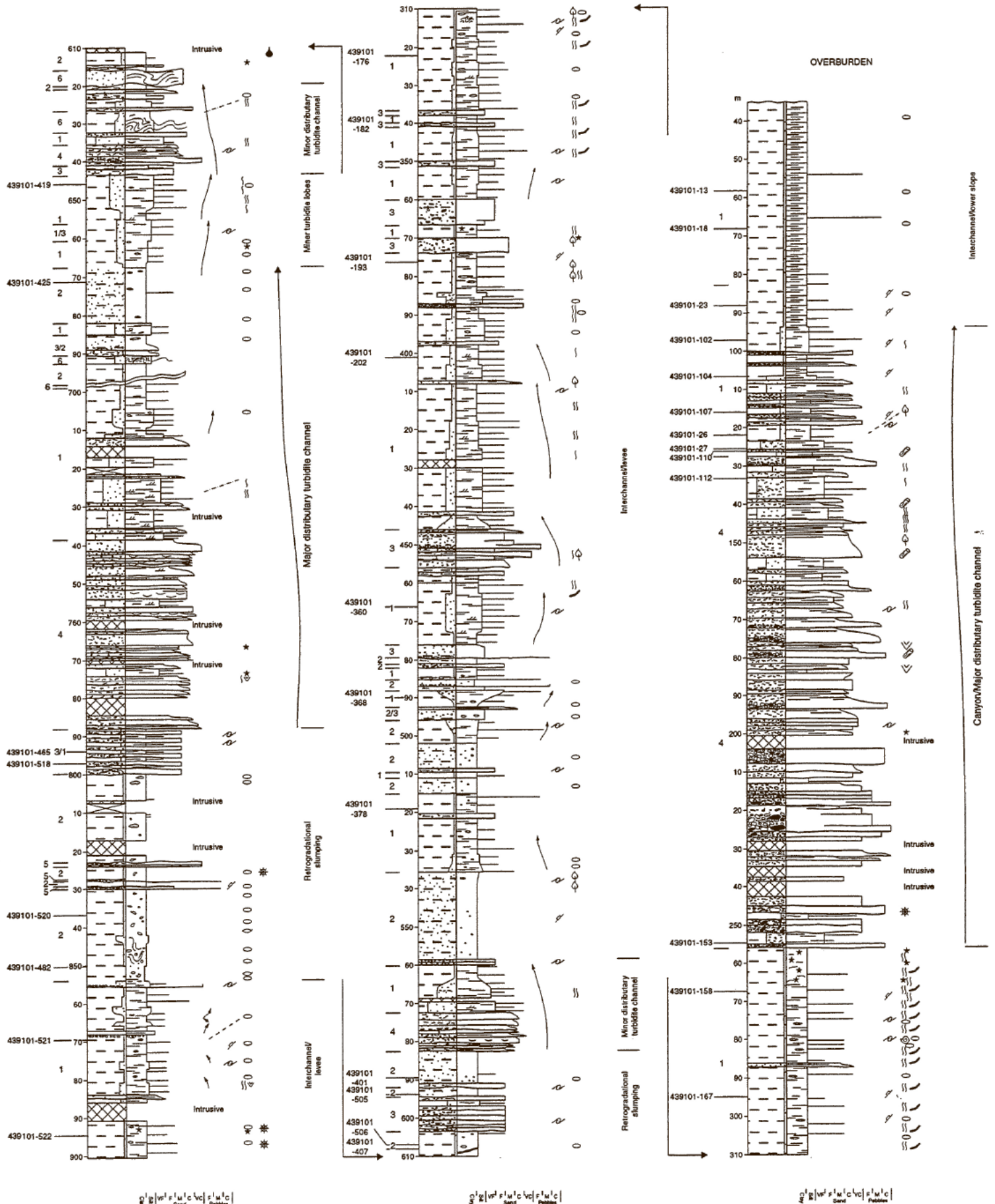


Fig. 5

## LEGEND

### Facies associations

- 1 Thinly interbedded sandstone and mudstone
- 2 Massive sandy mudstone
- 3 Massive muddy sandstone
- 4 Amalgamated sandstones and conglomerates
- 5 Single graded sandstone beds
- 6 Slumped beds



Volcanic sills



Clay and siltstone



Muddy sandstone/sandy mudstone



Sandstone



Sandstone with pebbles and mudstone clasts



Parallel lamination



Slumping



Disturbed bedding



Cross-lamination



Bioturbation



Concretions



Ammonite



Bivalve



Plant and wood fragments



Logs



Weakly bioturbated



Moderately bioturbated



Heavily bioturbated



*Planolites* isp.



*Helminthopsis horizontalis*



Escape burrows



Fractures



Pyrite



Oil



Gas



CU-succession



FU-succession

Fig. 5 continued

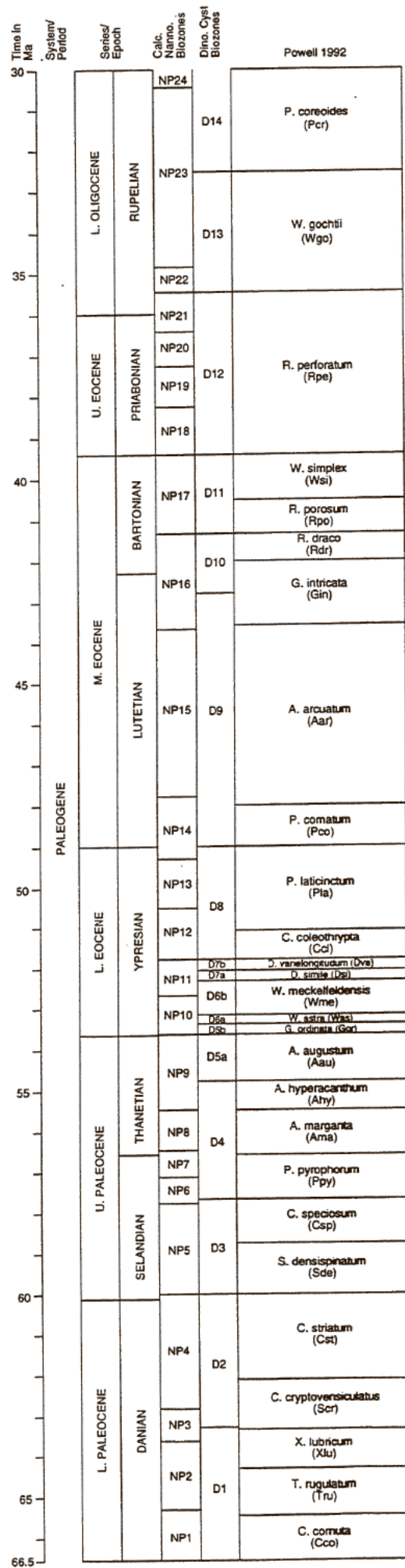


Fig. 6.

## Enclosures

- Enclosure 1. Range chart of the recorded palynomorphs, with division of GANK#1 into two palynological intervals.
- Enclosure 2. Range chart of the recorded palynomorphs, with division of GANE#1 into two palynological intervals.
- Enclosure 3. Range chart of the recorded palynomorphs, with division of GANT#1 into four palynological intervals.
- Enclosure 4. Detailed range chart of the recorded palynomorphs from the upper part of GANT#1.

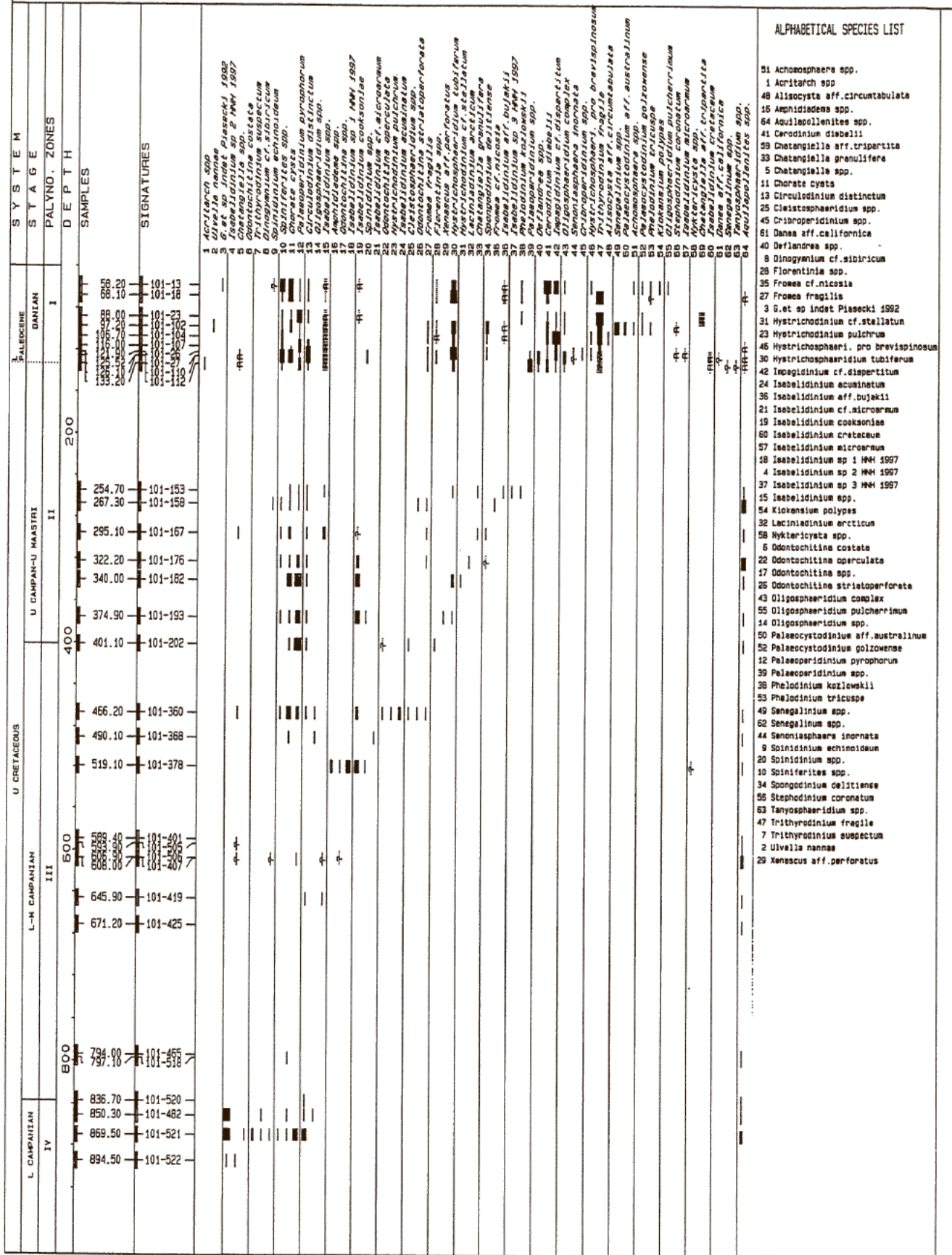
ALPHABETICAL SPECIES LIST

[illegible]

439001 GANE#1 ( 503 - 683m)

[illegible]

439101 GANT#1 ( 58 - 894m)



ENCLOSURE

SHEET: 1

Groups: A D S P E

4

SCALE: 1: 1000

Complete spp. set

GEUS

439101 GANT#1 ( 58 - 267m)

SYSTEM	STAGE	PALYNO. ZONES	DEPTH	SAMPLES	SIGNATURES
L PALEOCENE	DANIAN	1	50		
			58.20	101-13	
			68.10	101-18	
			88.00	101-23	
			97.20	101-102	
			106.70	101-104	
			116.00	101-107	
			121.90	101-26	
			128.50	101-27	
			133.20	101-112	
			150		
			200		
			250		
			254.70	101-153	

ALPHABETICAL SPECIES LIST
33 Achromosphaera spp. 1 Acritarch spp. 30 Alisocysta aff. circumtabulata 45 Aquilepollenites spp. 29 Cerodinium diebelii 40 Chatangiella aff. tripartita 15 Chatangiella granulifera 6 Chorata cysts 8 Circulodinium distinctum 27 Criloperidinium spp. 22 Deflandrea spp. 13 Florentinia spp. 17 Fromea cf. nicosia 12 Fromea fragilis 3 G. et sp indet Piasceki 1992 28 Hystrichosphaeridium pro brevispinosum 14 Hystrichosphaeridium tubiferum 24 Imagidinium cf. dispersitum 18 Isabelidium aff. bujakii 41 Isabelidium cretaceum 39 Isabelidium microarum 19 Isabelidium sp 3 MNH 1997 9 Isabelidium spp. 36 Kiokansium polypes 11 Odontochitina striatoperforata 25 Oligosphaeridium complex 37 Oligosphaeridium pulcherrimum 34 Palaeocystodinium aff. australinum 7 Palaeoperidinium pyrrophorum 21 Palaeoperidinium spp. 20 Phelodinium kozlowskii 35 Phelodinium tricuspe 31 Senegalinium spp. 43 Senegalinium spp. 26 Senoniasphaera inornata 4 Spinidinium echinoideum 10 Spinidinium spp. 5 Spiniferites spp. 16 Spongodinium delitense 38 Stephodinium coronatum 44 Tanyosphaeridium spp. 29 Trithyrodinium fragile 2 Ulvella nannae

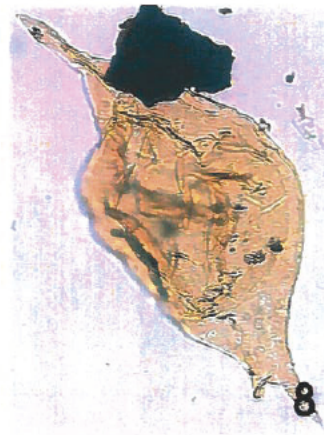
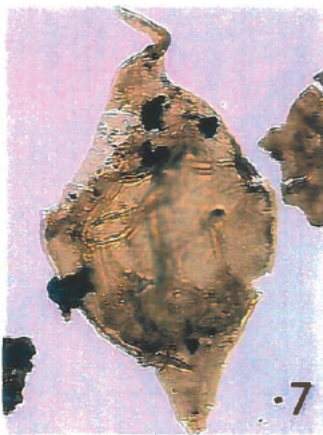
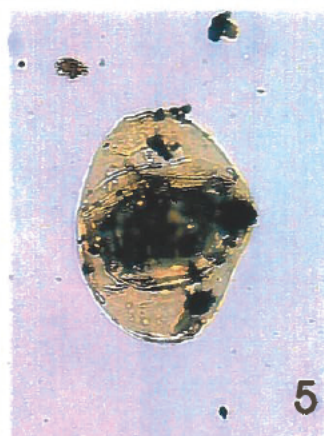
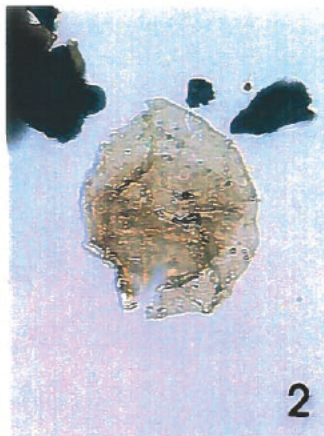
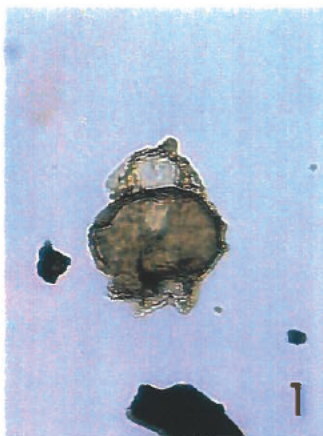
# ALPHABETICAL SPECIES LIST

- Achromosphaera spp.
- Acritarch spp.
- Alisocysta aff. circumtabulata
- Aquilapollenites spp.
- Cerodinium diebelii
- Chatangiella aff. tripartita
- Chatangiella granulifera
- Chorate cysts
- Circulodinium distinctum
- Cribrerodinium spp.
- Denea aff. californica
- Deflandrea spp.
- Florentinia spp.
- Fromea cf. nicosia
- Fromea fragilis
- G.et sp indet Piasceki 1992
- Hystrichosphaeri. pro brevispinosus
- Hystrichosphaeridium tubiferum
- Imagidinium cf. dispersitum
- Isabelidium aff. bujakii
- Isabelidium cretaceum
- Isabelidium microarum
- Isabelidium sp 3 MNH 1997
- Isabelidium spp.
- Kiokansium polypea
- Odontochitina striatoperforata
- Oligosphaeridium complex
- Oligosphaeridium pulcherrimum
- Palaeocystodinium aff. australinum
- Palaeocystodinium golzowense
- Palaeoperidinium pyrrophorum
- Palaeoperidinium spp.
- Phelodinium kozlowskii
- Phelodinium tricuspe
- Senegalinium spp.
- Senegalinium spp.
- Senoniasphaera inornata
- Spinidinium echinoideum
- Spinidinium spp.
- Spiniferites spp.
- Spongodinium delitense
- Stephodinium corenatum
- Tanyosphaeridium spp.
- Trithyrodinium fragile
- Ulvella nannae

## Plate 1, GANK#1.

- Fig. 1. *Isabelidinium* aff. *viborgense*, GGU 439201-38-4, 24.9–108.0; LVR 1.8863; MI 6396.
- Fig. 2. *Isabelidinium* aff. *viborgense*, GGU 439201-38-4, 30.3–94.4; LVR 1.8864; MI 6397.
- Fig. 3. *Isabelidinium* aff. *acuminatum*, GGU 439201-50-3, 25.9–100.0; LVR 1.8888; MI 6415.
- Fig. 4. *Isabelidinium* *cretaceum*, GGU 439201-50-7, 42.3–99.5; LVR 1.8890; MI 6416.
- Fig. 5. *Isabelidinium* *cretaceum*, GGU 439201-50-7, 47.3–102.0; LVR 1.8891; MI 6417.
- Fig. 6. *Isabelidinium* *cretaceum*, GGU 439201-50-7, 42.8–112.3; LVR 1.8892; MI 6418.
- Fig. 7. *Palaeocystodinium* *bulliforme*, GGU 439201-37-7, 41.0–97.6; LVR 1.8875; MI 6404.
- Fig. 8. *Palaeocystodinium* *bulliforme*, GGU 439201-48-9, 36.9–111.6; LVR 1.8884; MI 6411.
- Fig. 9. *Palaeocystodinium* *bulliforme*, (length: 259  $\mu$ m) GGU 439201-44.8, 52.8–104.2; LVR 1.8881; MI 6409.
- Fig. 10. *Palaeocystodinium* *golzowense*, GGU 439201-48.8, 33.7–108.0; LVR 1.8886; MI 6413.
- Fig. 11. *Cerodinium* *speciosum*, GGU 439201-40.5, 42.0–103.5; LVR 1.8878; MI 6407.
- Fig. 12. *Cerodinium* *speciosum* supsp. *glabrum*, GGU 439201-40-5, 43.5–111.5 LVR 1.8879; MI 6408.

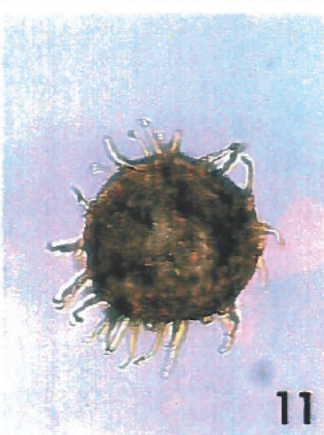
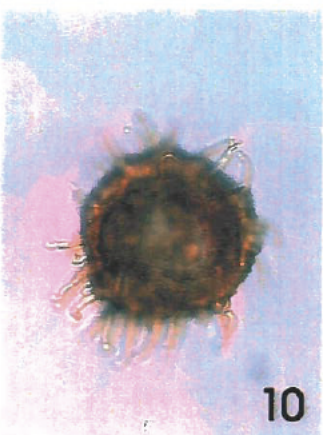
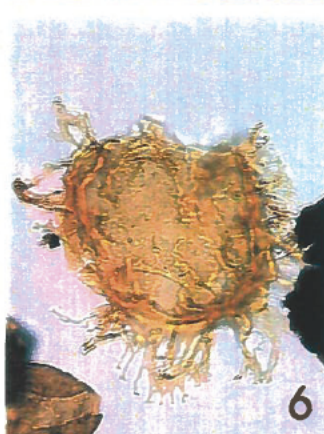
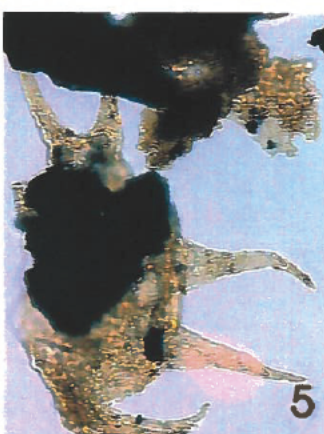
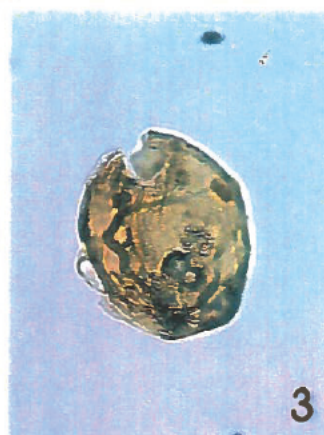
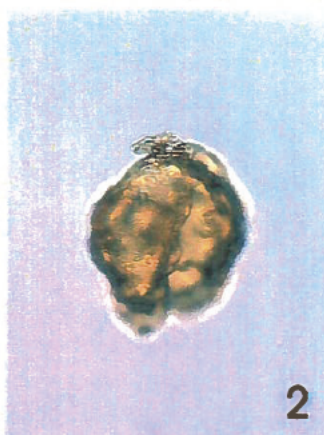
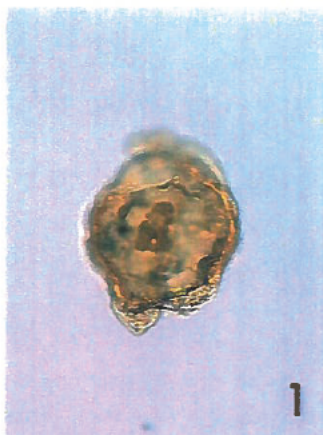
20  $\mu\text{m}$



## Plate 2, GANK#1.

- Fig. 1. Dinocyst 01 HNH, 1997, GGU 439201-38-4, 21.3–96.8; LVR 1.8867; MI 6399.
- Fig. 2. Dinocyst 01 HNH, 1997, GGU 439201-38-4, 21.3–96.8; LVR 1.8868; MI 6399.
- Fig. 3. Dinocyst 01 HNH, 1997, GGU 439201-38-4, 21.0–101.1; LVR 1.8870; MI 6400.
- Fig. 4. *Gonyaulacysta?* sp., GGU 439201-38-4, 23.3–102.8; LVR 1.8858; MI 6394.
- Fig. 5. Gen. et sp. Indet. Piasecki 1992, GGU 439201-38-7, 28.9–104.1; LVR 1.8874; MI 6403.
- Fig. 6. *Glaphyrocysta gippingensis*, GGU 439201-40-5, 29.9–111.0; LVR 1.8876; MI 6405.
- Fig. 7. *Impagidinium* cf. *dispertitum*, GGU 439201-50-4, 33.2–96.4; LVR 1.8893; MI 6419.
- Fig. 8. *Impagidinium* cf. *dispertitum*, GGU 439201-50.-, 33.2–96.4; LVR 1.8894; MI 6419.
- Fig. 9. *Fibroysta* cf. *capitata*, GGU 439201-118-4, 28.7–99.0; LVR 1.8898; MI 6422.
- Fig. 10. *Fibroysta* cf. *capitata*, GGU 439201-118-4, 26.7–106.7; LVR 1.8896; MI 6421.
- Fig. 11. *Fibroysta* cf. *capitata*, GGU 439201-118-4, 26.7–106.7; LVR 1.8897; MI 6421.
- Fig. 12. *Fibroysta* cf. *capitata*, GGU 439201-118-4, 31.0–98.3; LVR 1.8899; MI 6423.

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### Plate 3, GANE#1.

- Fig. 1. *Alisocysta margarita*, GGU 439001-514-4, 18.8–113.7; LVR 1.8839; MI 6383.
- Fig. 2. *Alisocysta margarita*, GGU 439001-514-4, 18.8–113.7; LVR 1.8840; MI 6383.
- Fig. 3. *Alisocysta margarita*, GGU 439001-514-6, 47.3–104.9; LVR 1.8841; MI 6384.
- Fig. 4. Dinocyst 01 HNH, 1997, GGU 439001-370-4, 23.0–94.3; LVR 1.8847; MI 6388.
- Fig. 5. Dinocyst 01 HNH, 1997, GGU 439001-370-4, 26.7–95.2; LVR 1.8849; MI 6389.
- Fig. 6. Dinocyst 01 HNH, 1997, GGU 439001-369-4, 23.5–98.3; LVR 1.8844; MI 6387.
- Fig. 7. Dinocyst 01 HNH, 1997, GGU 439001-370-4, 25.0–100.5; LVR 1.8854; MI 6392.
- Fig. 8. Dinocyst 01 HNH, 1997, GGU 439001-370-4, 36.5–102.8; LVR 1.8851; MI 6390.
- Fig. 9. *Trithyrodinium* sp., GGU 439001-370-3, 20.6–108.8; LVR 1.8853; MI 6391
- Fig. 10. *Areoligera gippingensis*, GGU 439001-514-4, 24.3–103.6; LVR 1.8842; MI 6385.
- Fig. 11. *Cerodinium striatum*, GGU 439001-515-4, 33.6–102.6; LVR 1.8843; MI 6386.
- Fig. 12. *Paleocystodinium bulliforme*, GGU 439001-371-4, 50.9–105.2; LVR 1.8856; MI 6393.

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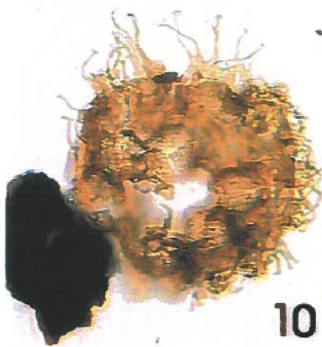
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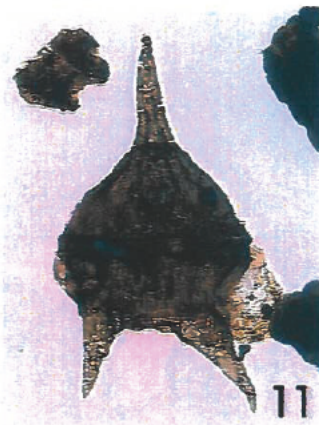
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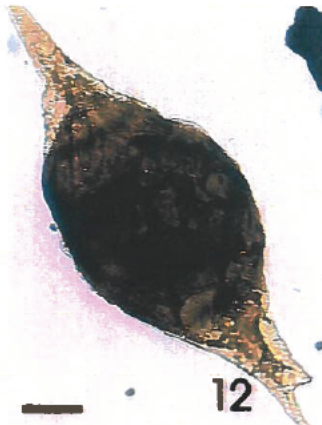
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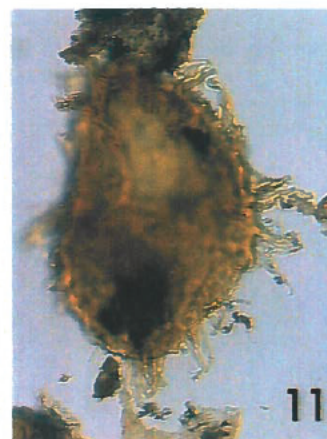
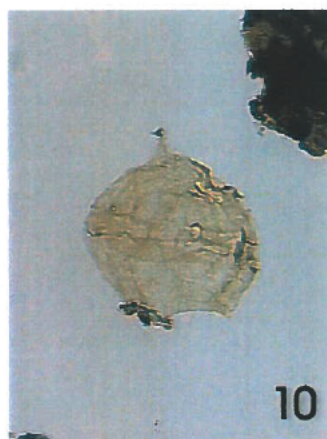
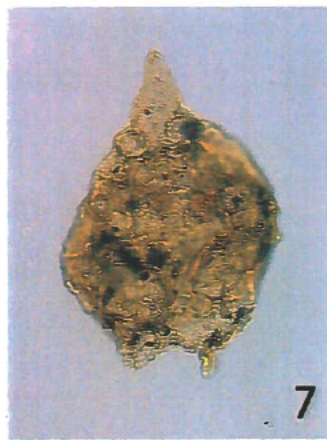
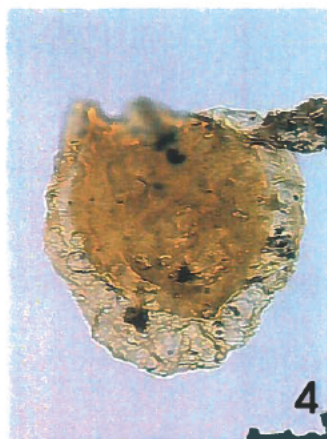
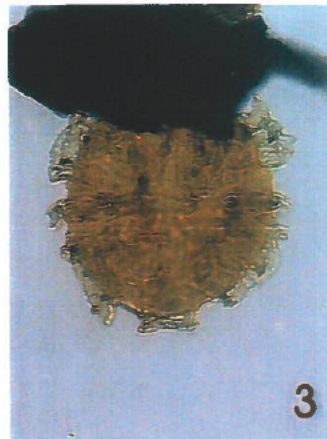


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## Plate 4, GANT#1.

- Fig. 1 *Trithyrodinium fragile*, GGU 439101-23-2, 46.5–103.9; LVR 1.9123; MI 6578.
- Fig. 2. ?*Trithyrodinium fragile*, GGU 439101-112-10, 33.1–111.7; LVR 1.9122; MI 6577.
- Fig. 3. *Alisocysta circumtabulata*, GGU 439101-104-9, 53.5–99.4; LVR 1.9125; MI 6579.
- Fig. 4. *Senoniasphaera inornatum*, GGU 439101-27-9, 39.8–97.4; LVR 1.9126; MI 6580.
- Fig. 5. *Senoniasphaera inornatum*, GGU 439101-27-10, 38.5–107.7; LVR 1.9127; MI 6581.
- Fig. 6. *Senoniasphaera inornatum*, GGU 439101-27-10, 38.5–107.7; LVR 1.9128; MI 6581.
- Fig. 7. *Deflandrea* sp., GGU 439101-27-9, 47.9–106.4; LVR 1.9129; MI 6582.
- Fig. 8. *Deflandrea* sp., GGU 439101-27-9, 54.2–106.1; LVR 1.9130; MI 6583.
- Fig. 9. *Paleoperidinium* sp., GGU 439101-112-6, 23.4–109.8; LVR 1.9131; MI 6584.
- Fig. 10. *Paleoperidinium* sp., GGU 439101-110-9, 32.3–108.2; LVR 1.9132; MI 6585.
- Fig. 11. ?*Danea* aff. *californica*, GGU 439101-110-69, 43.9–112.9; LVR 1.9133 MI 6586.
- Fig. 12. ?*Danea* aff. *californica*, GGU 439101-110-69, 43.9–112.9; LVR 1.9134 MI 6586.

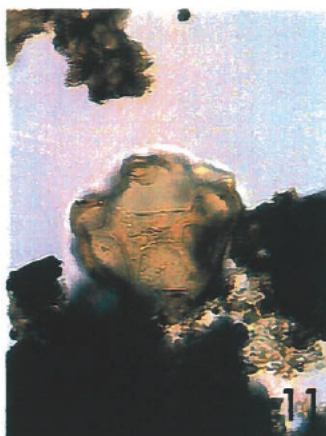
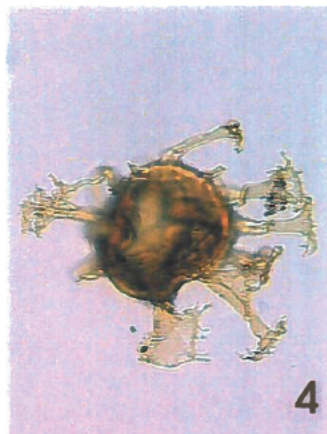
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## Plate 5, GANT#1.

- Fig. 1. Gen. et sp. Indet. Piasecki 1992, GGU 439101-13-9, 36.9–103.2; LVR 1.8900; MI 6424.
- Fig. 2. *Chatangiella* aff. *tripartita*, GGU 439101-23-4, 48.4–104.0; LVR 1.8901; MI 6425.
- Fig. 3. *Chatangiella* aff. *tripartita*, GGU 439101-23-4, 20.6–104.4; LVR 1.8902; MI 6426.
- Fig. 4. *Hystrichosphaeridium tubiferum*, GGU 439101-102-4, 28.3–111.3; LVR 1.8903; MI 6427.
- Fig. 5. *Palaeocystodinium* aff. *australinum*, 439101-102-4, 26.8–97.2; LVR 1.8904; MI 6428.
- Fig. 6. *Palaeocystodinium* aff. *australinum*, 439101-102-4, 30.1–95.5; LVR 1.8906; MI 6429.
- Fig. 7. *Phelodinium tricuspe*, 439101-102-4, 41.5–95.8; LVR 1.8911; MI 6434.
- Fig. 8. *Stephodinium coronatum*, 439101-26-4, 42.8–101.2; LVR 1.8916; MI 6437.
- Fig. 9. *Dinoymnium* cf. *sibiricum*, 439101-482-3, 41.6–109.6; LVR 1.8932; MI 6450.
- Fig. 10. *Trithyrodinium fragile*, GGU 439101-102-4, 36.7–99.1; LVR 1.9112; MI 6436.
- Fig. 11. *Trithyrodinium suspectum*, GGU 439101-521-9, 39.6–98.6; LVR 1.9138; MI 6455.
- Fig. 12. *Trithyrodinium suspectum*, GGU 439101-521-9, 34.2–108.0; LVR 1.9140; MI 6456.

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## Plate 6, GANT#1.

- Fig. 1. *Isabelidinium* aff. *bujakii*, GGU 439101-102-4, 44.6–109.6; LVR 1.8909; MI 6432
- Fig. 2. *Isabelidinium* sp. 3 HNH 1997, GGU 439101-102-4, 40.2–106.9; LVR 1.8917; MI 6438.
- Fig. 3. *Isabelidinium* cf. *acuminatum*, GGU 439101-360-3, 16.2–112.9; LVR 1.8920; MI 6440.
- Fig. 4. *Isabelidinium* cf. *microarmum*, GGU 439101-368-3, 44.2–102.5; LVR 1.8921; MI 6441.
- Fig. 5. *Isabelidinium* *cooksoniae*, GGU 439101-378-7, 54.0–101.4; LVR 1.8922; MI 6442.
- Fig. 6. *Isabelidinium* *cooksoniae*, GGU 439101-378-4, 38.0–112.5; LVR 1.8925; MI 6444.
- Fig. 7. *Isabelidinium* sp. 1 HNH 1997, GGU 439101-378-4, 27.2–104.3; LVR 1.8923; MI 6443.
- Fig. 8. *Isabelidinium* sp. 1 HNH 1997, GGU 439101-378-8, 33.0–105.6; LVR 1.8926; MI 6445.
- Fig. 9. *Isabelidinium* sp. 2 HNH 1997, GGU 439101-482-3, 40.7–100.2; LVR 1.8928; MI 6447.
- Fig. 10. *Isabelidinium* sp. 2 HNH 1997, GGU 439101-482-3, 50.5–112.8; LVR 1.8929; MI 6448.
- Fig. 11. *Isabelidinium* sp. 2 HNH 1997, GGU 439101-482-4, 38.8–101.8; LVR 1.8931; MI 6449.
- Fig. 12. *Isabelidinium* sp. 2 HNH 1997, GGU 439101-522-6, 22.8–111.8; LVR 1.8941; MI 6457.

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